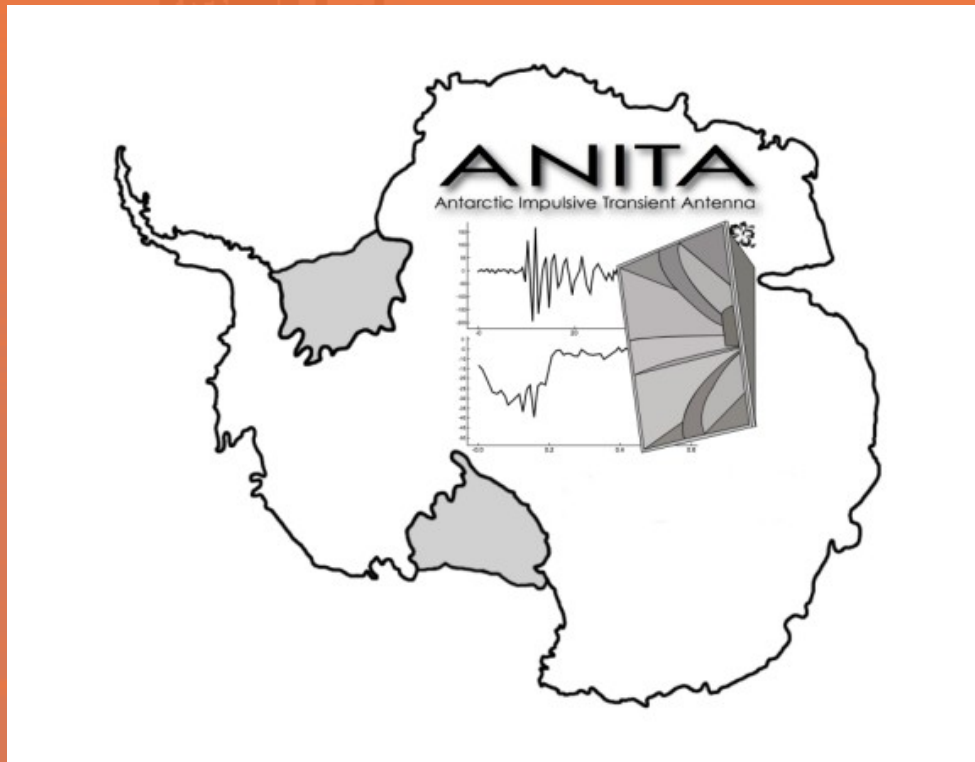


# Use of Radar Depth Sounding Data to Estimate Radio Frequency Attenuation in Greenlandic and Antarctic Ice



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University of Kansas  
for the ANITA Collaboration

**KU** THE UNIVERSITY OF  
**KANSAS**

# Importance

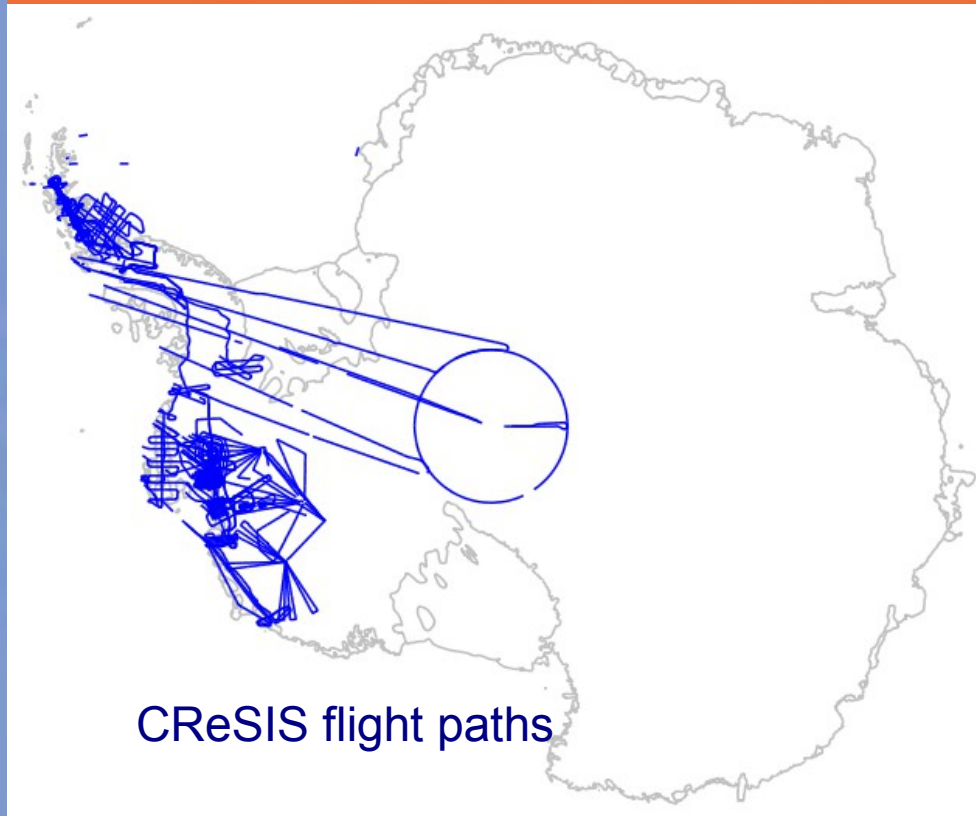
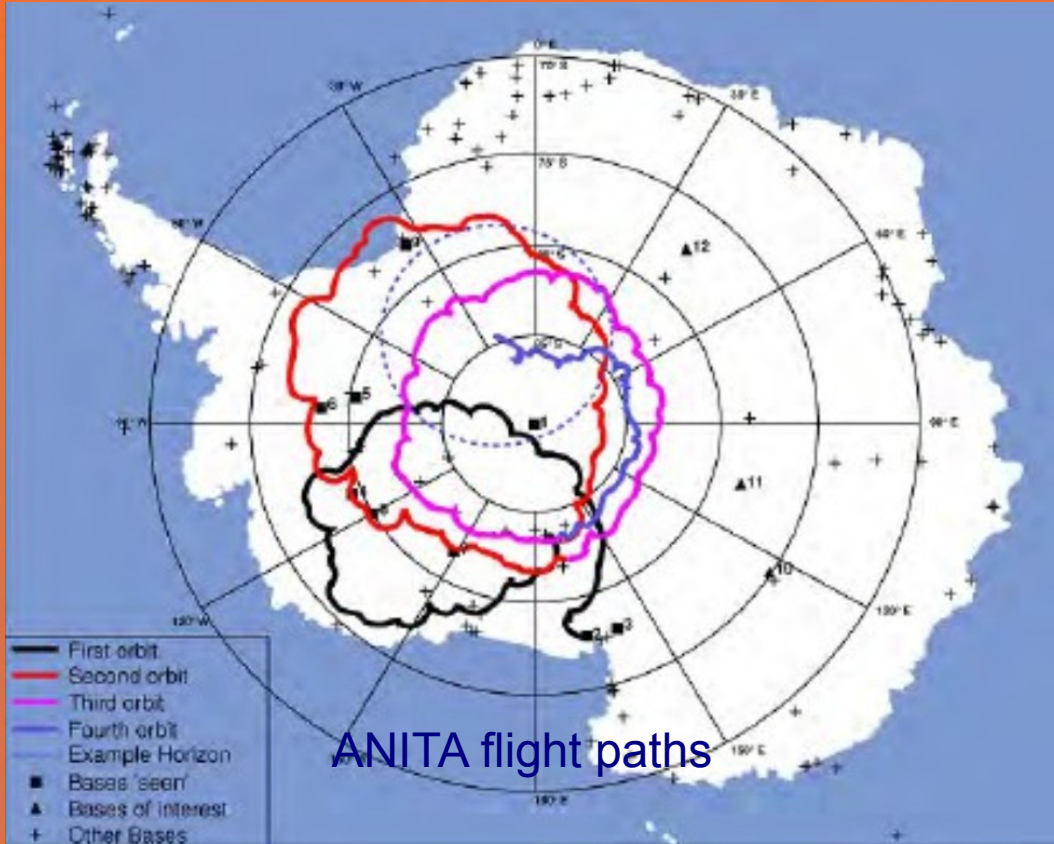
- ANITA is a suite of antennas that rides on a balloon “around the world,” looking at in-ice neutrino events
- As radio signals produced by the Askaryan effect from neutrino-ice interaction propagate through the ice they are attenuated (fall off by more than geometric amount)
- After accounting for attenuation, the neutrino energy can be estimated
  - UHE neutrinos are of most importance because of their relation to unexplained UHE CRs, and offer a better chance at directional reconstruction
- Needed for MC effective volume estimate



# Current State

- We currently have a model from Amir Javaid (University of Delaware), but continent-wide tests have not been done
- Center for Remote Sensing of Ice Sheets (CReSIS) has database of radar depth sounding (RDS) with good coverage in West Antarctica
  - Can potentially use these data to test our model
- Unfortunately for us, Greenland data is better (higher sampling rate) so use it for technique development





# Goals

Estimate depth-dependent attenuation length from RDS data



Match estimate from echogram close to a core with prediction from that core

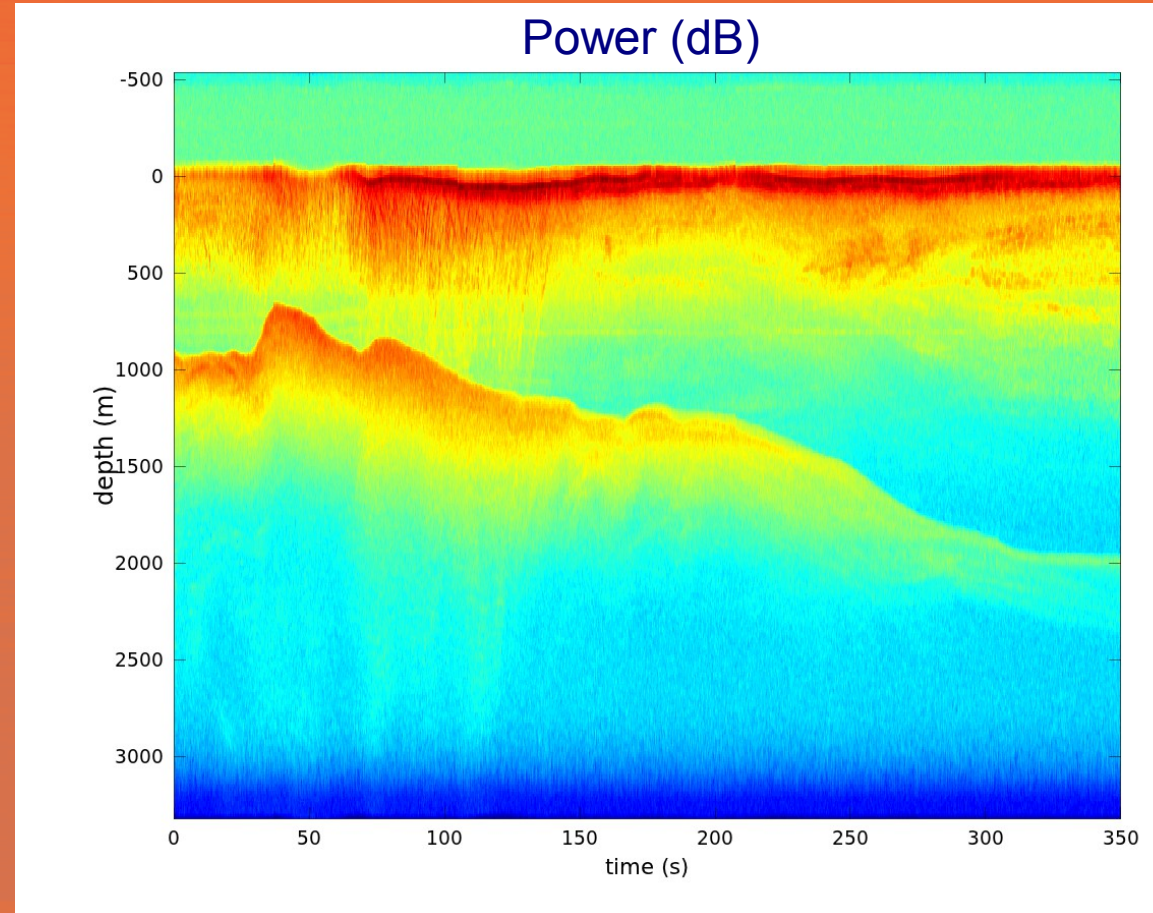


Determine if modeled attenuation length profiles are reasonable compared to estimates from RDS data

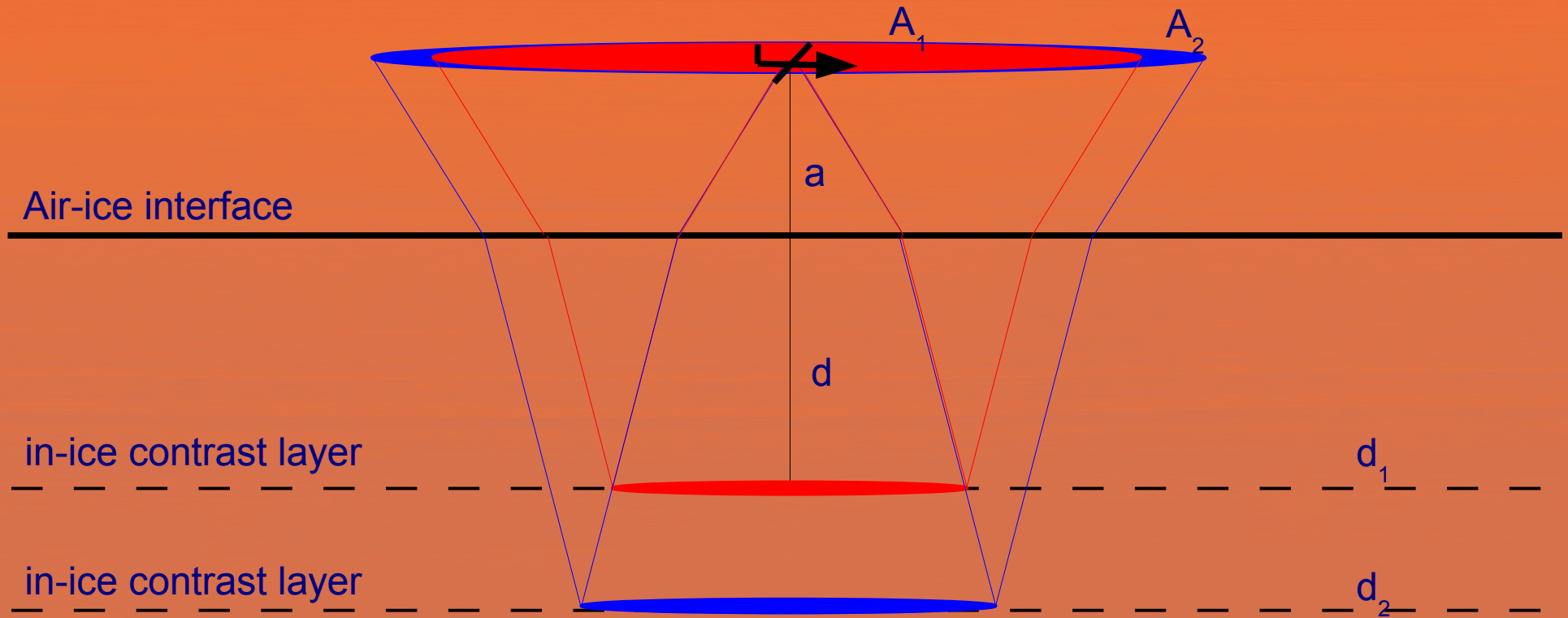


# Data – CReSIS MCoRDS

- Flyover w/ radar pulsing  $\sim 195\text{MHz}$  (BW  $\sim 10\text{MHz}$ )
- 2 channels
  - Low gain for surface bounce
  - High gain for bottom bounce
- Put through bounce detection algorithm and stitched together (accounting for gain mismatch)



# RDS cartoon



# Attenuation Length

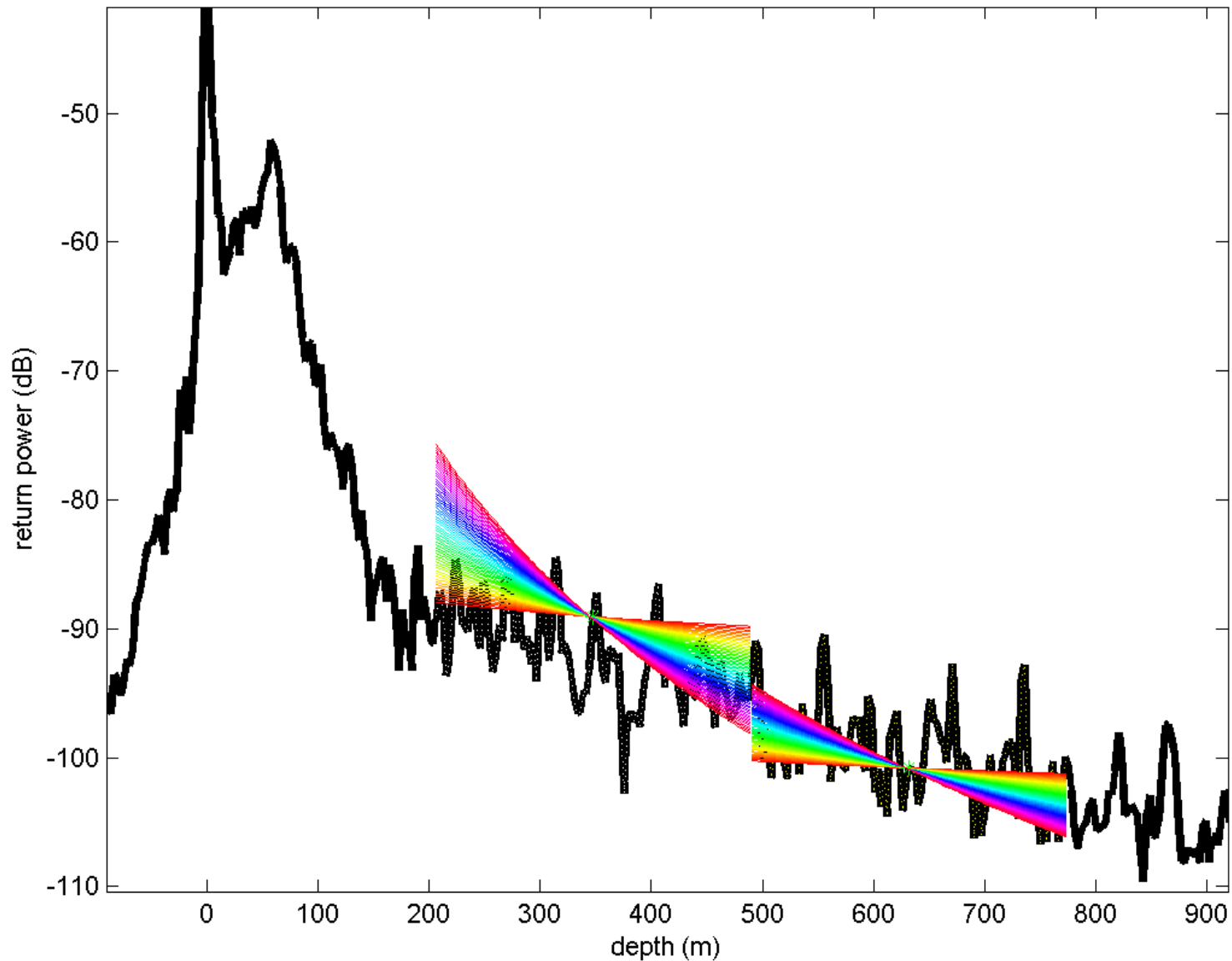
- Wish to solve for the attenuation length – the distance over which the amplitude decreases by a factor of  $e$
- The expected form of the data should be a power law multiplied by an exponential decay representing geometrical spreading and attenuation, respectively

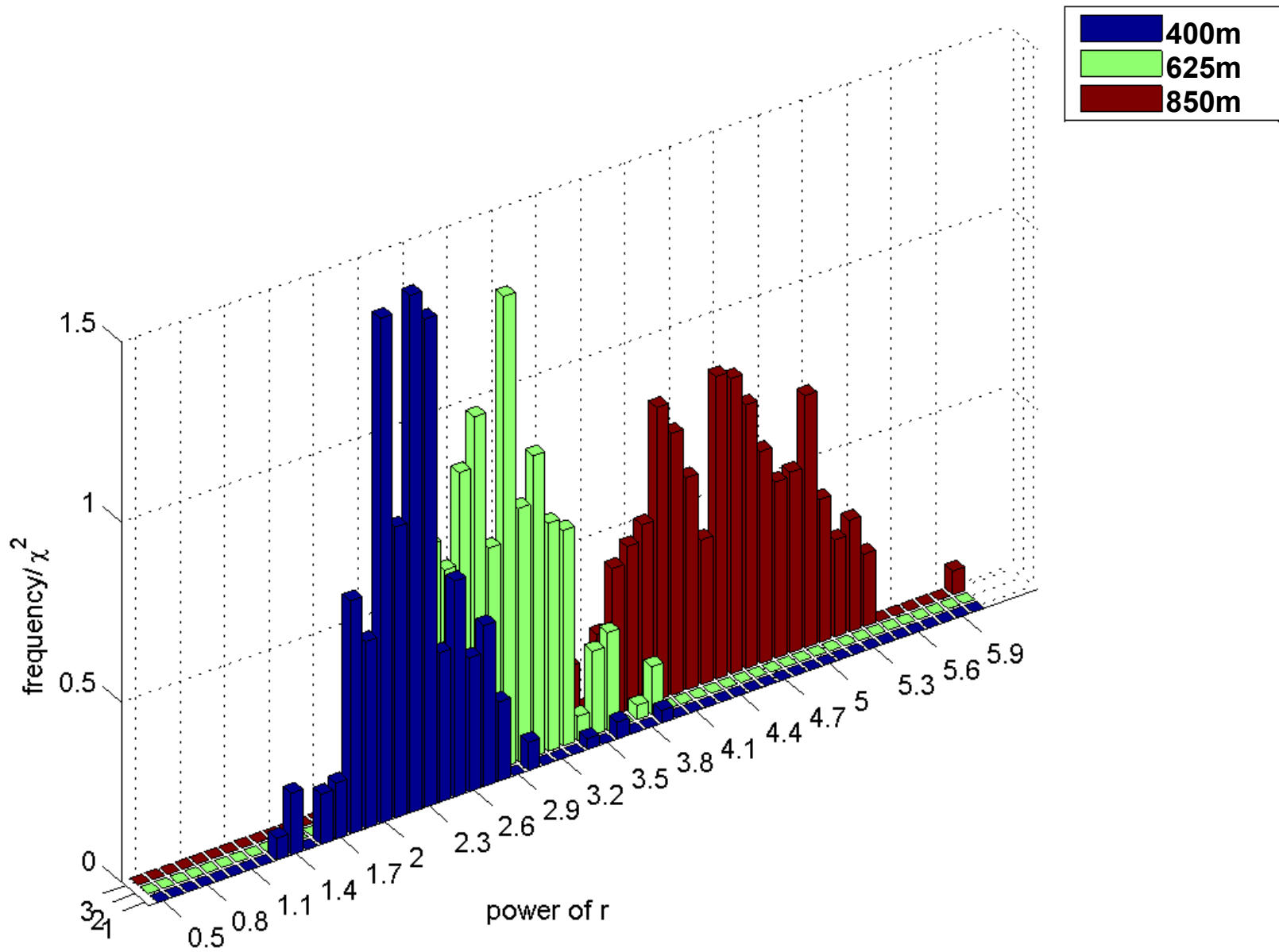
$$A \propto A_0 \frac{1}{(2a + 2d)^n} \exp\left(\frac{-2d}{L_\alpha}\right)$$





# Determining Appropriate Power Law



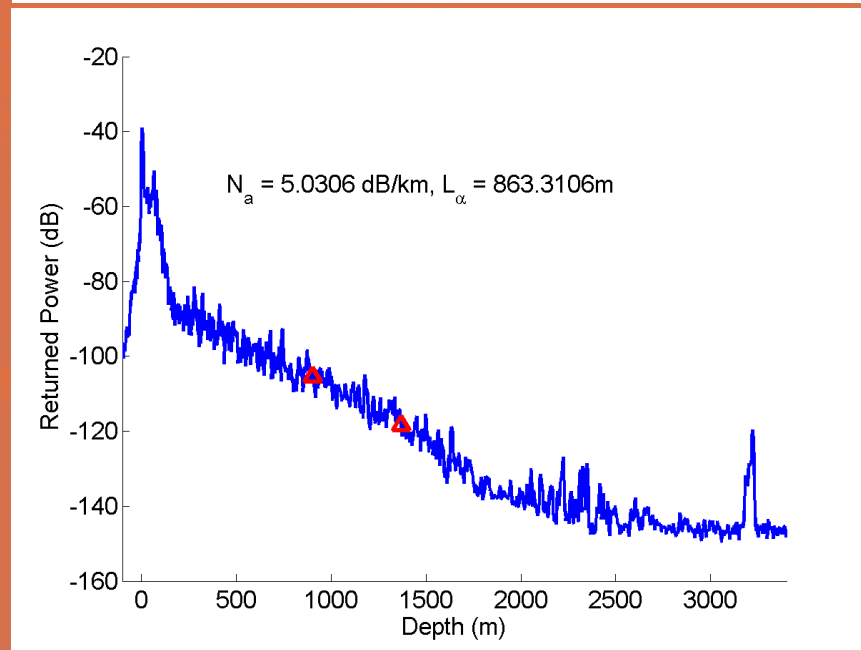
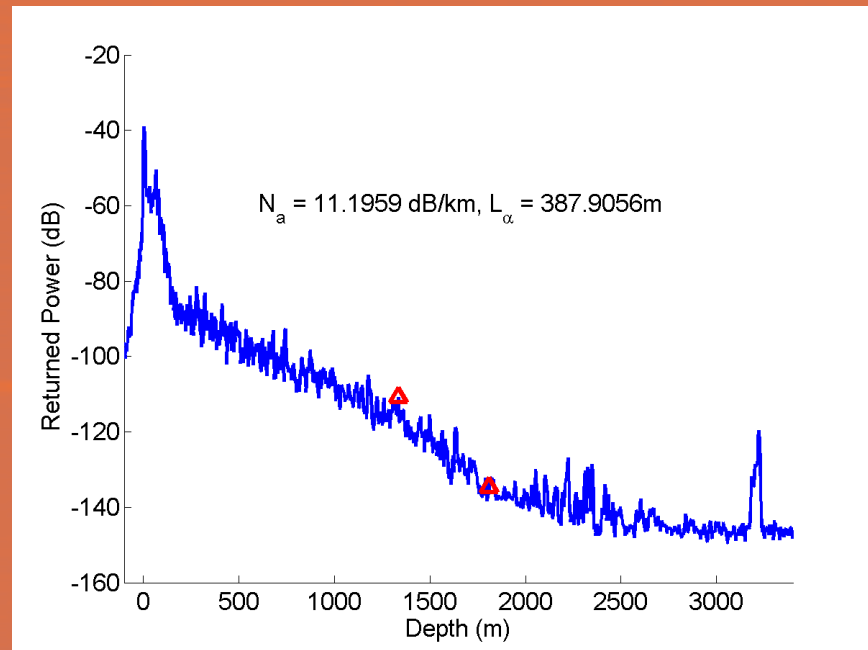
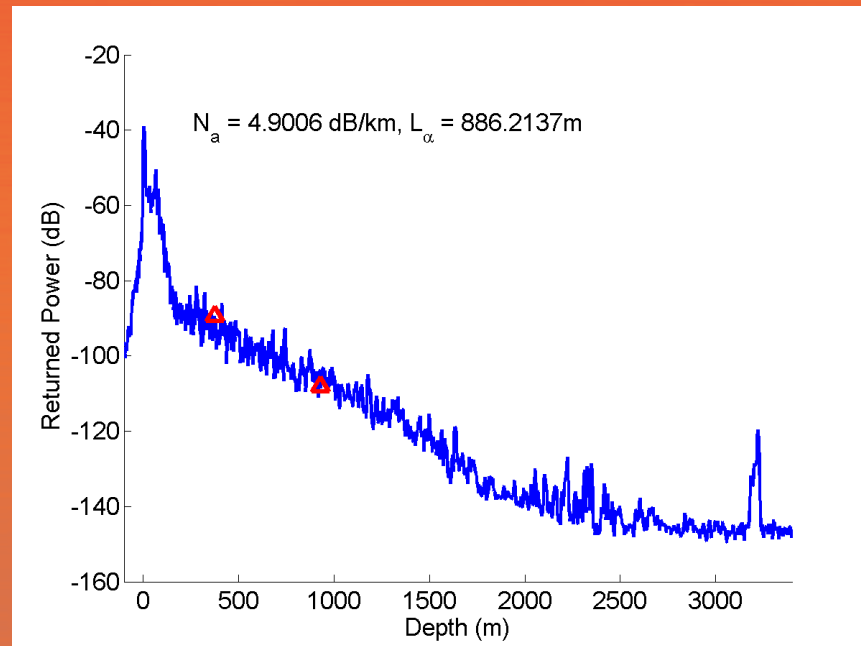
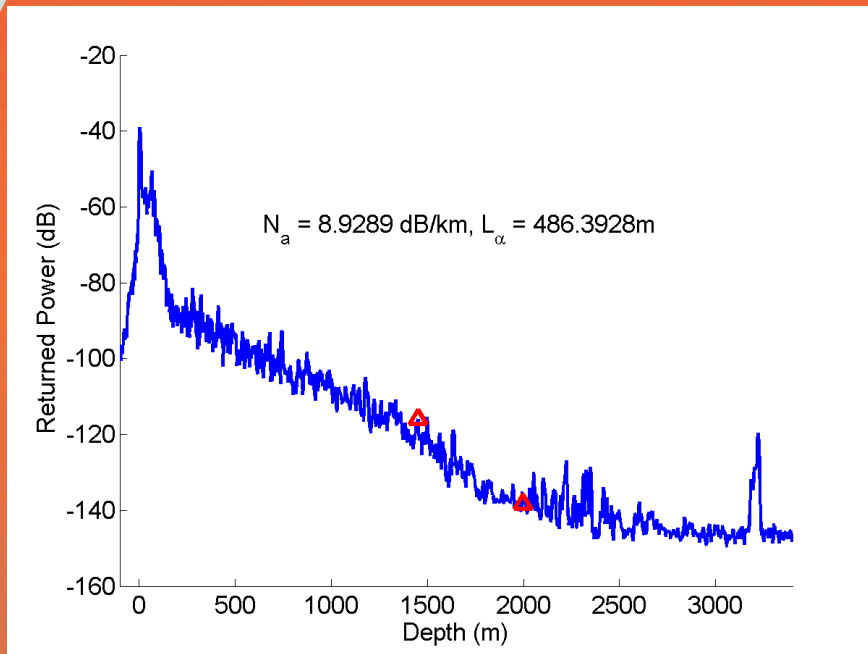


# 2-point Method

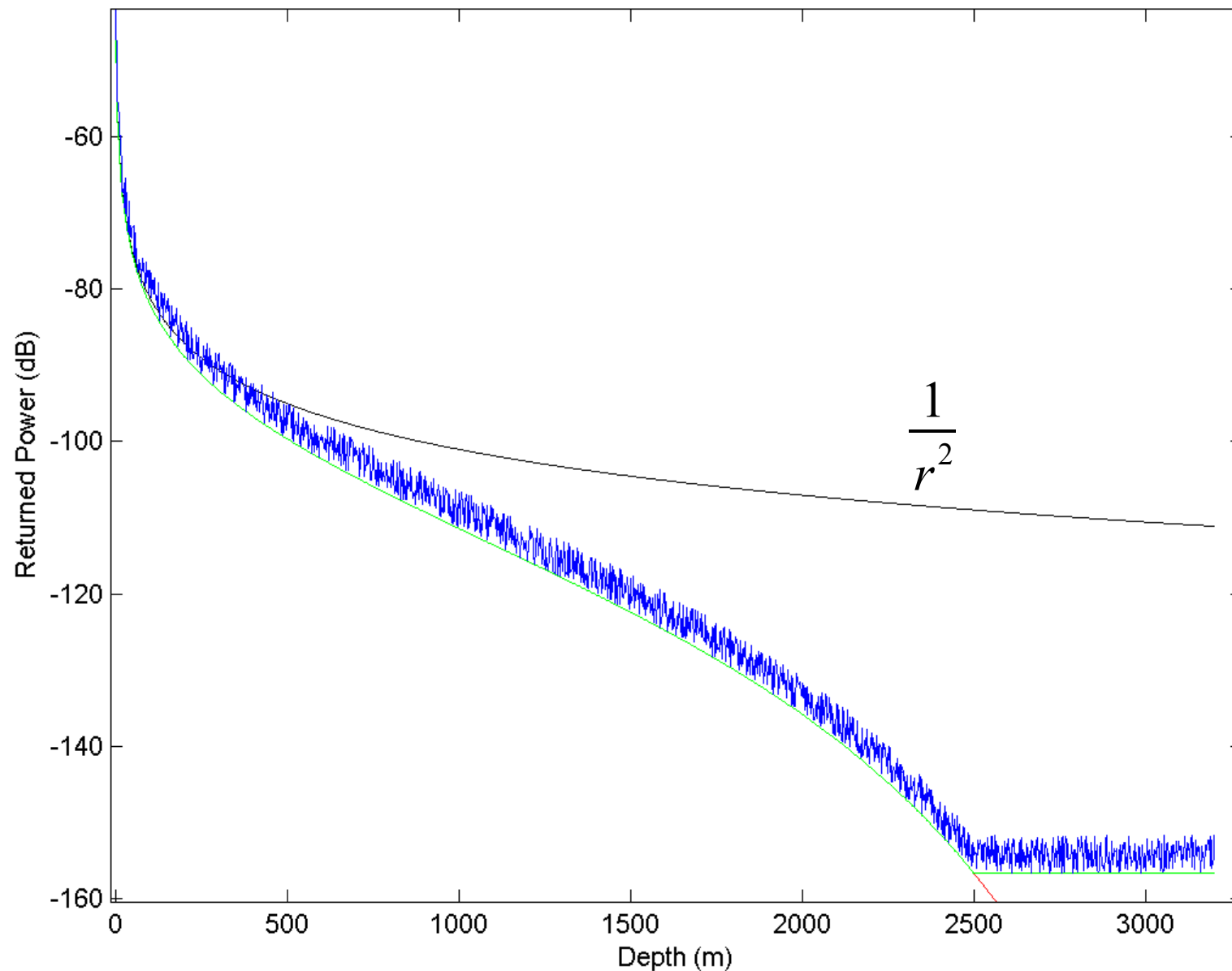
- Picks a random segment from a ping that is between 400-600m long at a random starting point (endpoint must be before thermal noise plateau)
- Calculates change in power due to  $1/r^2$ , whatever remains is attributed to attenuation
- $$N_a = \frac{1000}{2(r_2 - r_1)} \left( P_1 - P_2 + 20 \log \left( \frac{r_1 + a_1}{r_2 + a_2} \right) \right)$$
- Does this large number of times to diminish effects of “bad” pairs
- Statistical fit in log space is converted to attenuation length
- Assumptions:
  - The mechanism that produces the backscattering falls off as  $1/r$  in amplitude and is constant function of depth
    - Many layers with very small and approx equal reflection coefficients would satisfy this
  - Slowly varying attenuation length profile
- Does NOT assume any temperature information or have any fitted offset



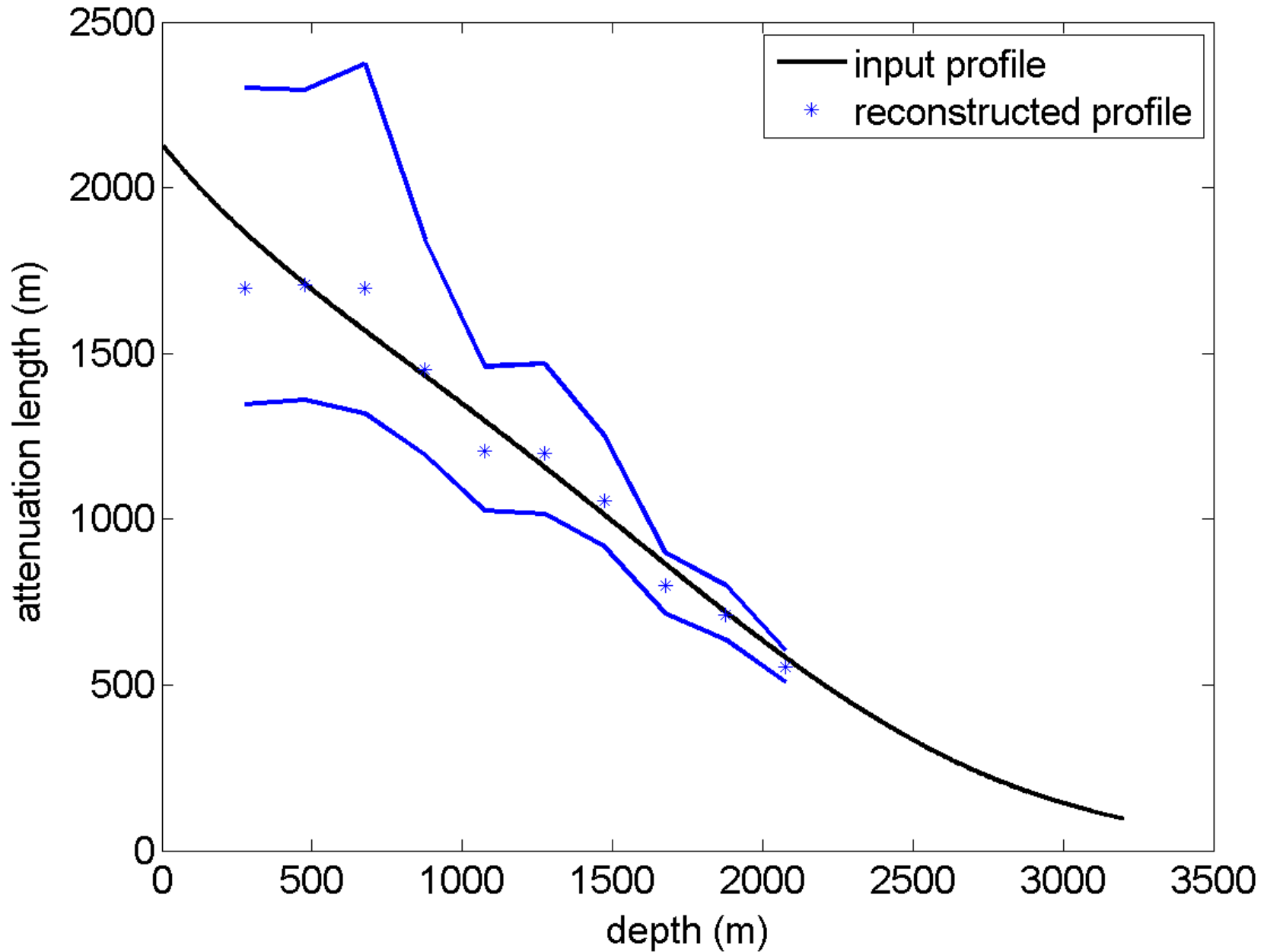
# Examples



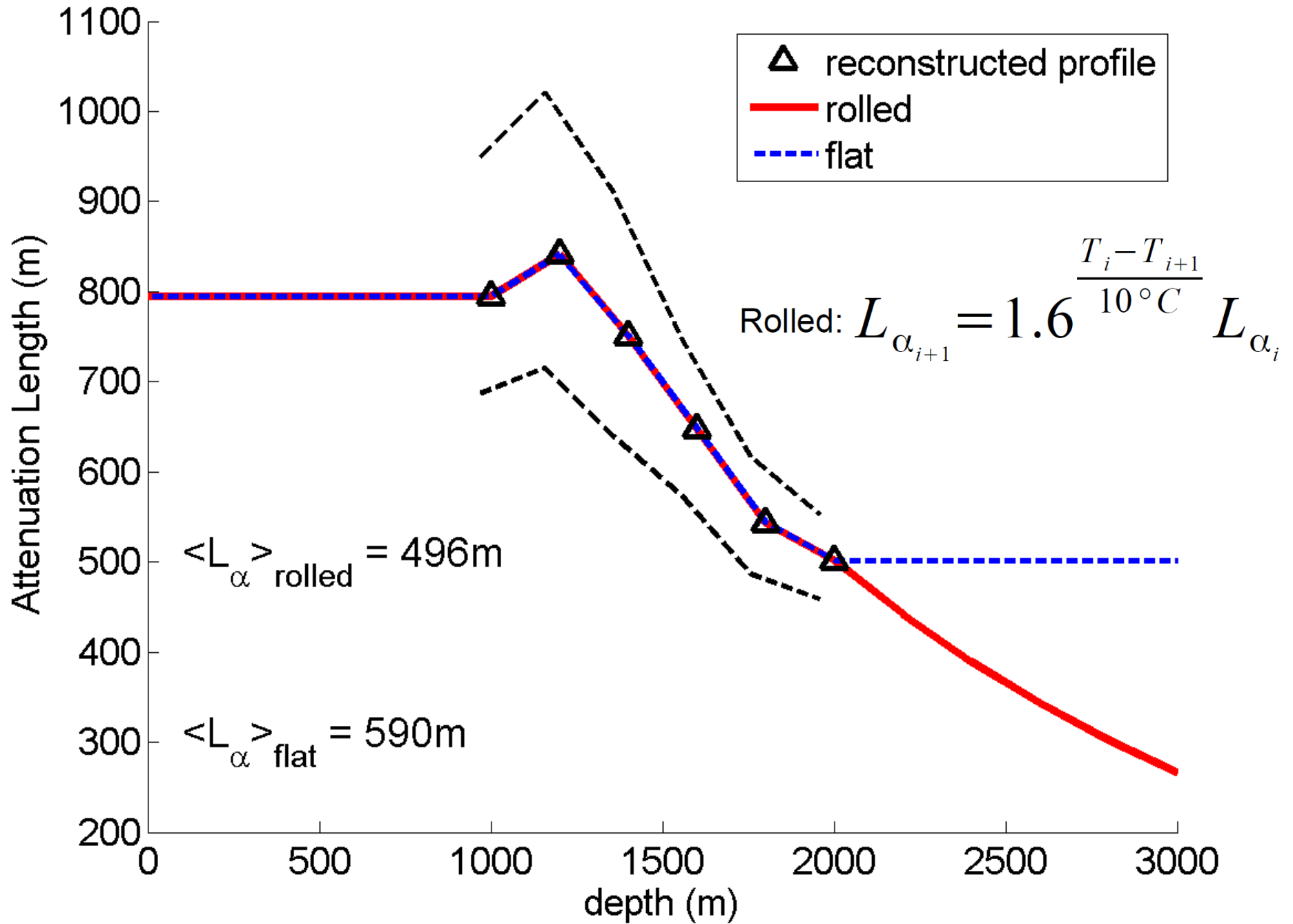
# Synthetic Example

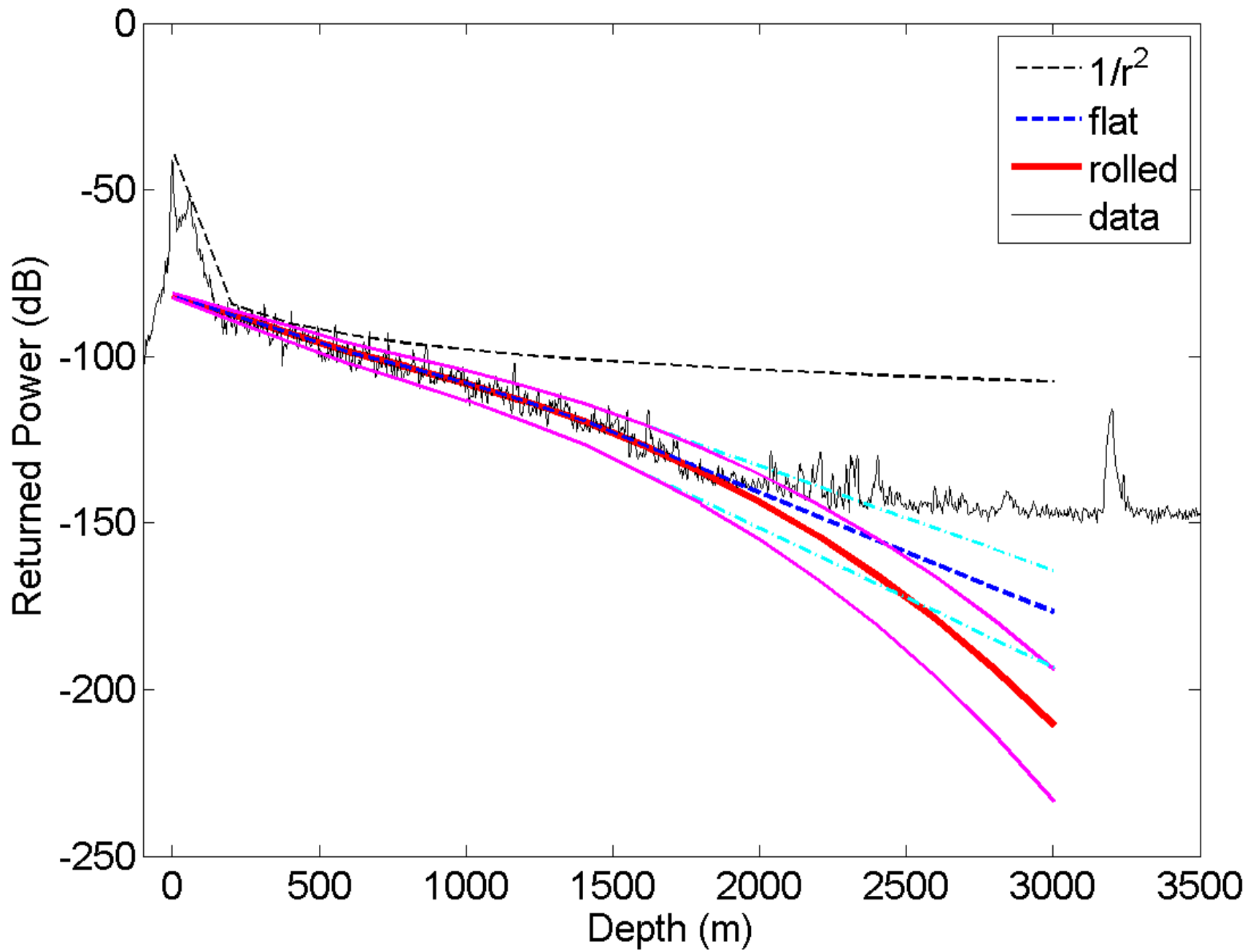


# Fit to Synthetic Data



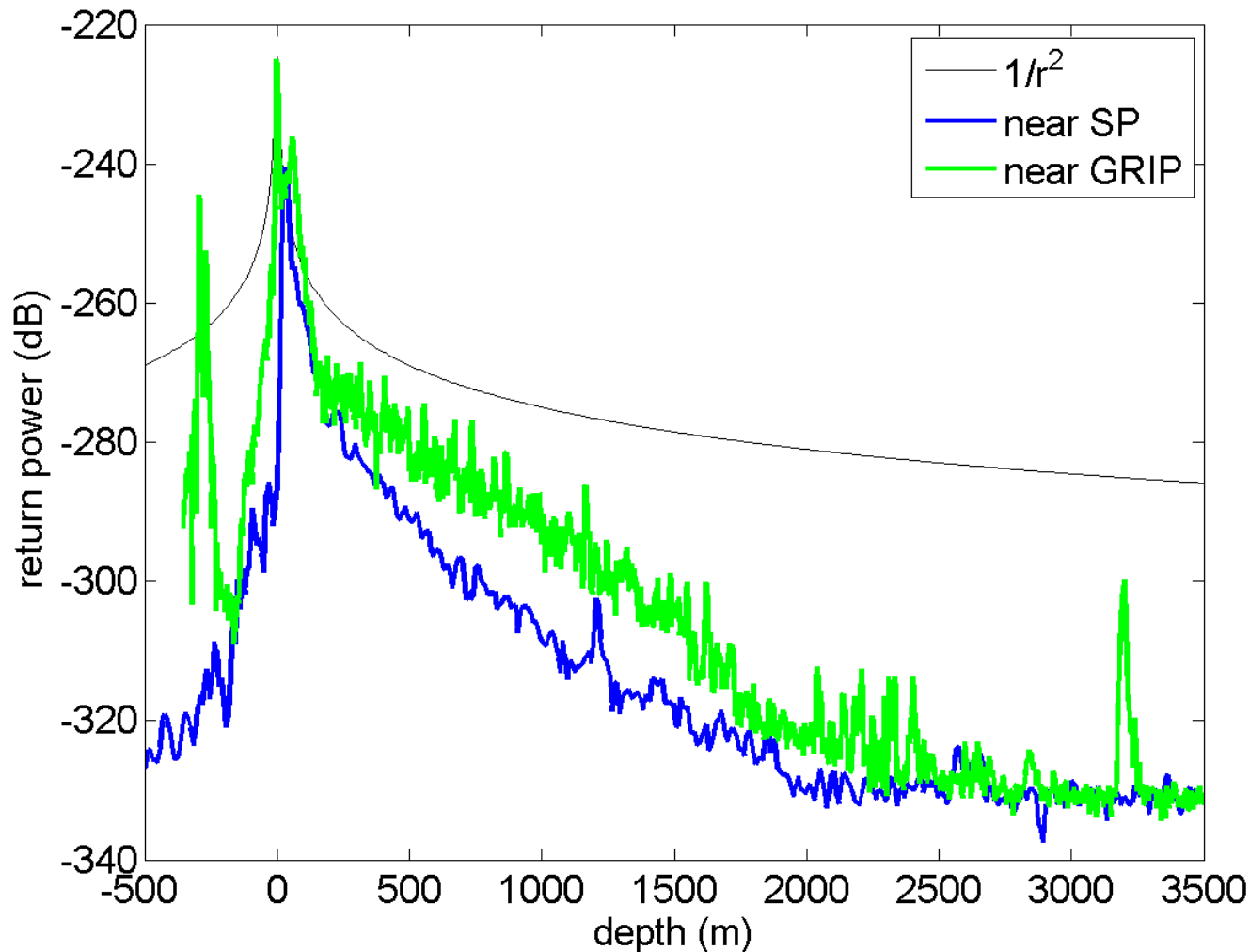
# 190 MHz





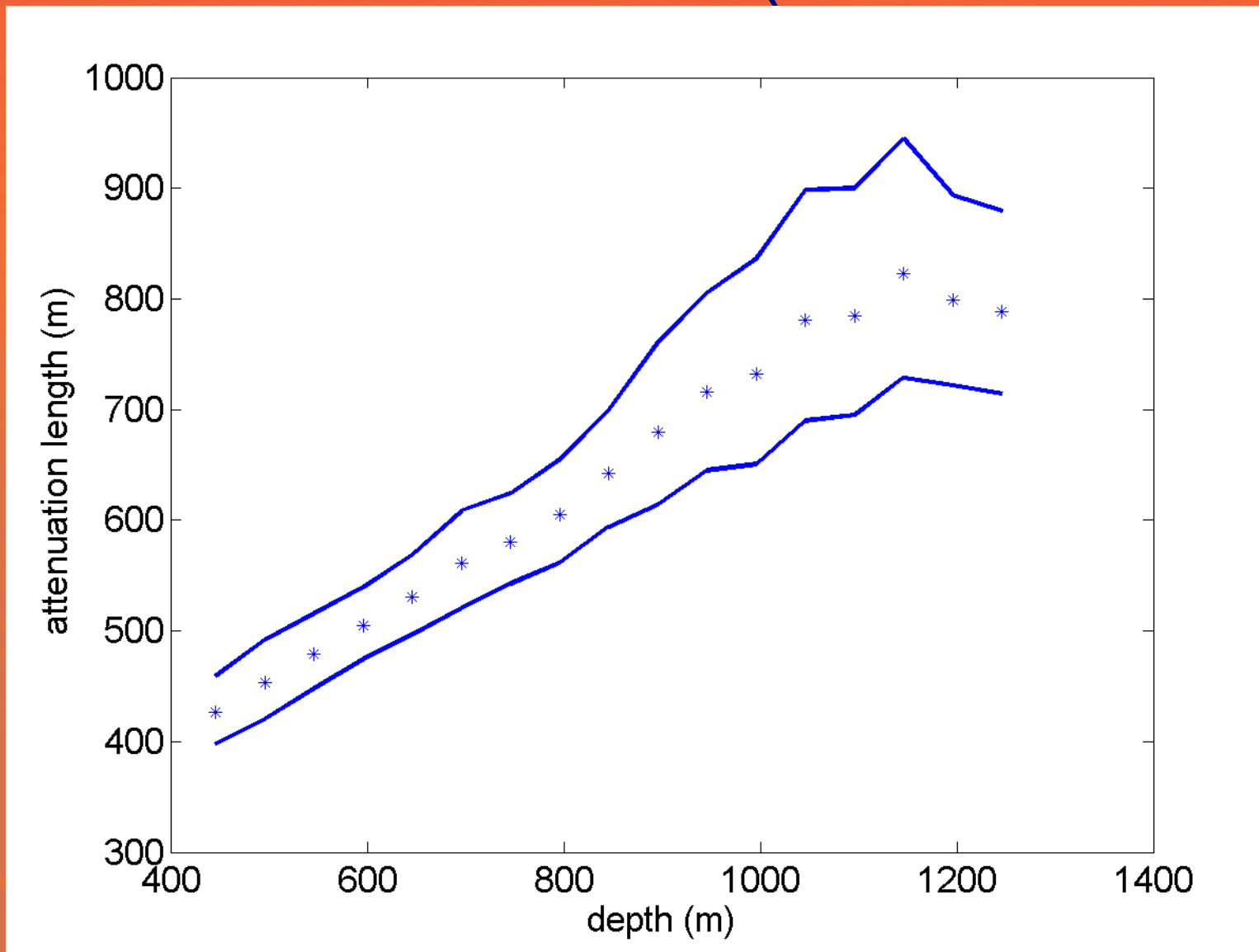


# Comparison to Antarctic Data



- Antarctic and Greenlandic returns look very different (supposedly same DAQ)
- Speculation: constant reflector assumption that works in Greenland is invalid

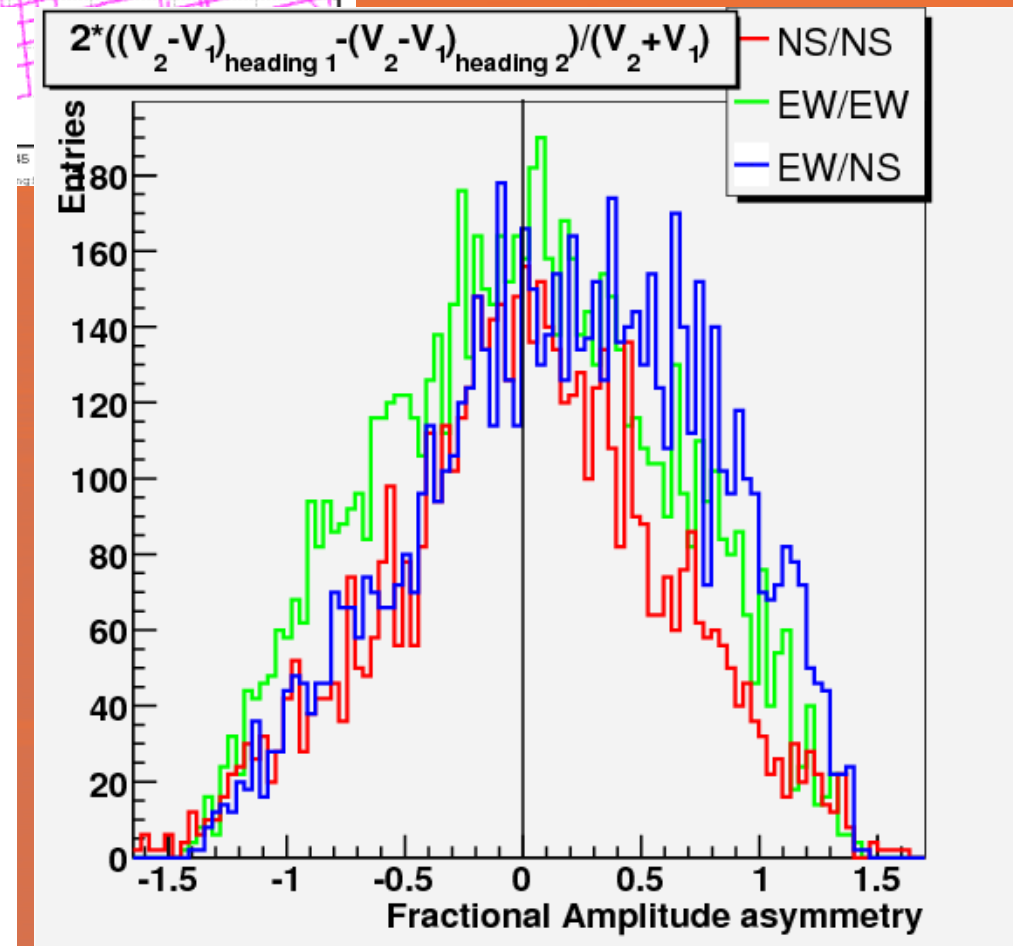
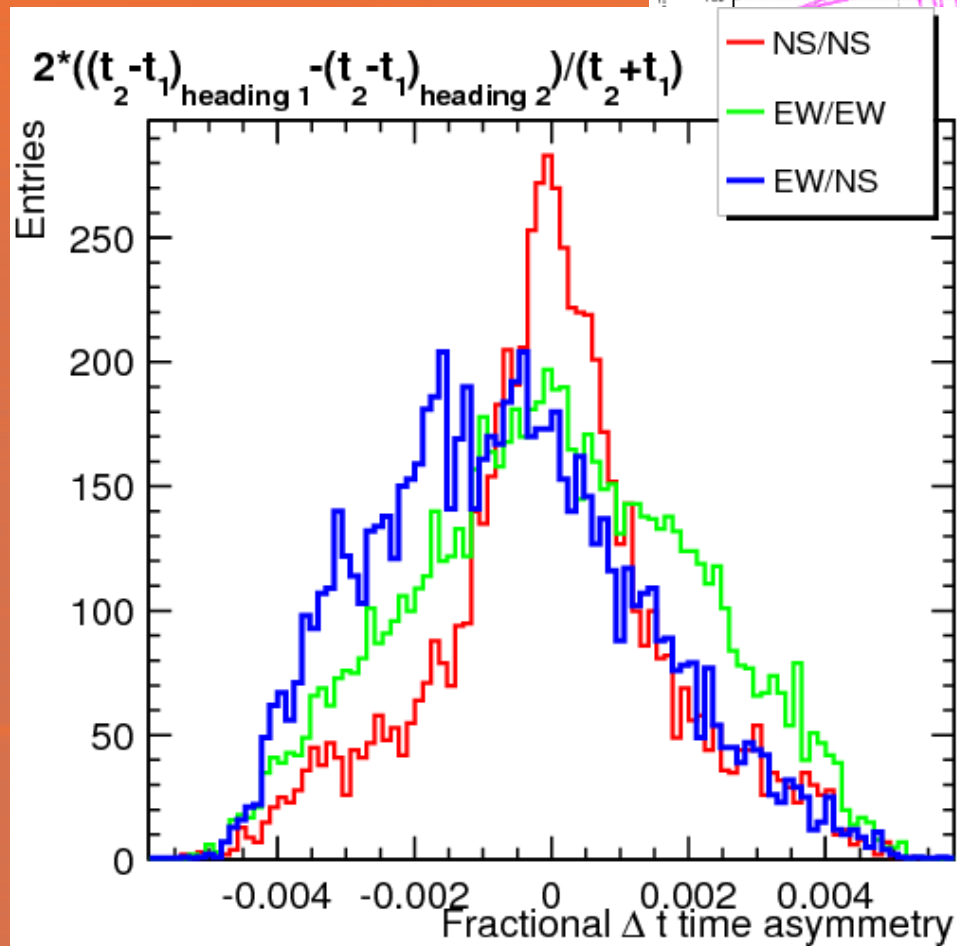
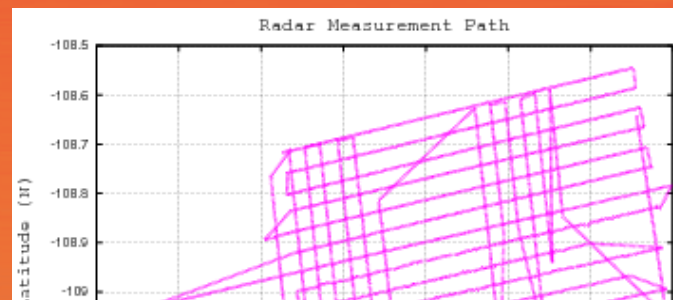
# Implied South Pole Profile (without correction)



Inverted from what we would expect



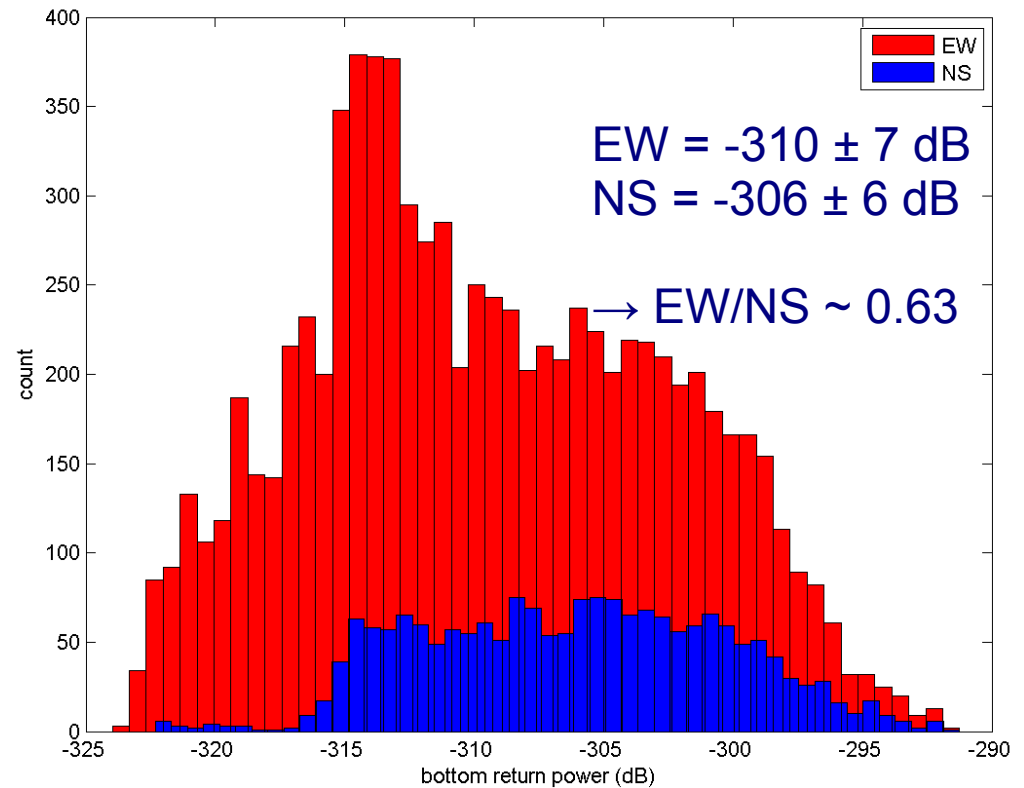
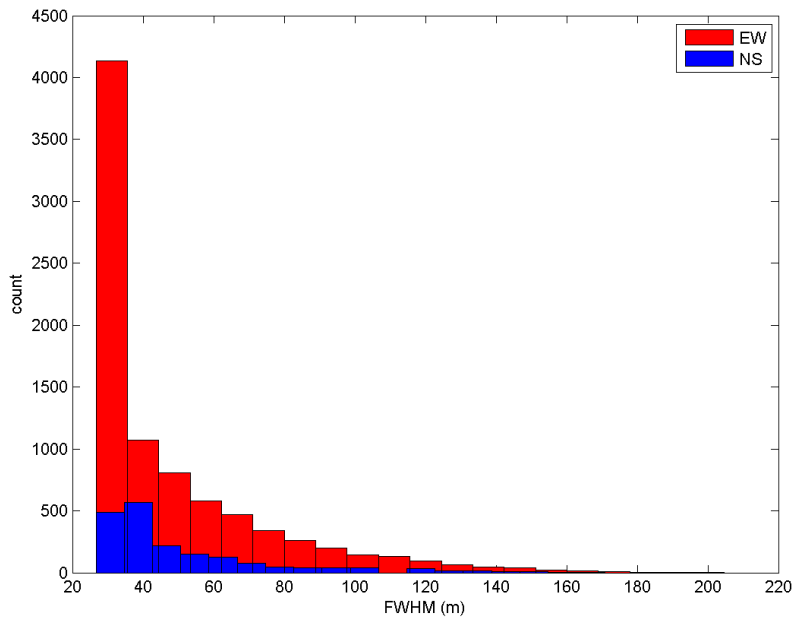
# Birefringence (Thwaites)



Analysis and figures by Dave Besson



# Near SP



No strong evidence of reflection broadening, but majority of data is on-axis (NS or EW)

Approximately the same voltage asymmetry is observed, but points are not co-located



- Conclusions:
  - Good fit with synthetic data (even with large noise values)
  - Hard to compare with GRIP profile (temperature-only attenuation profile from data is a bad fit, Greenland potentially has larger chemical effect)
  - Sampled 2-point method is fairly robust and does not require record by record tuning as did some previous methods
    - Minimum-maximum segment length and vertical bin number responsible for most change
- Need to:
  - Find reliable borehole record with nearby CReSIS data to truly verify method as well as tune those “universal” variables
    - Investigate correcting Antarctic data
  - Compare to values given by Amir's model
  - Other ice properties from CReSIS? More with birefringence, also radar bands of 550-900 MHz, 2-6 GHz, and 13-17 GHz



# Thank You Questions?

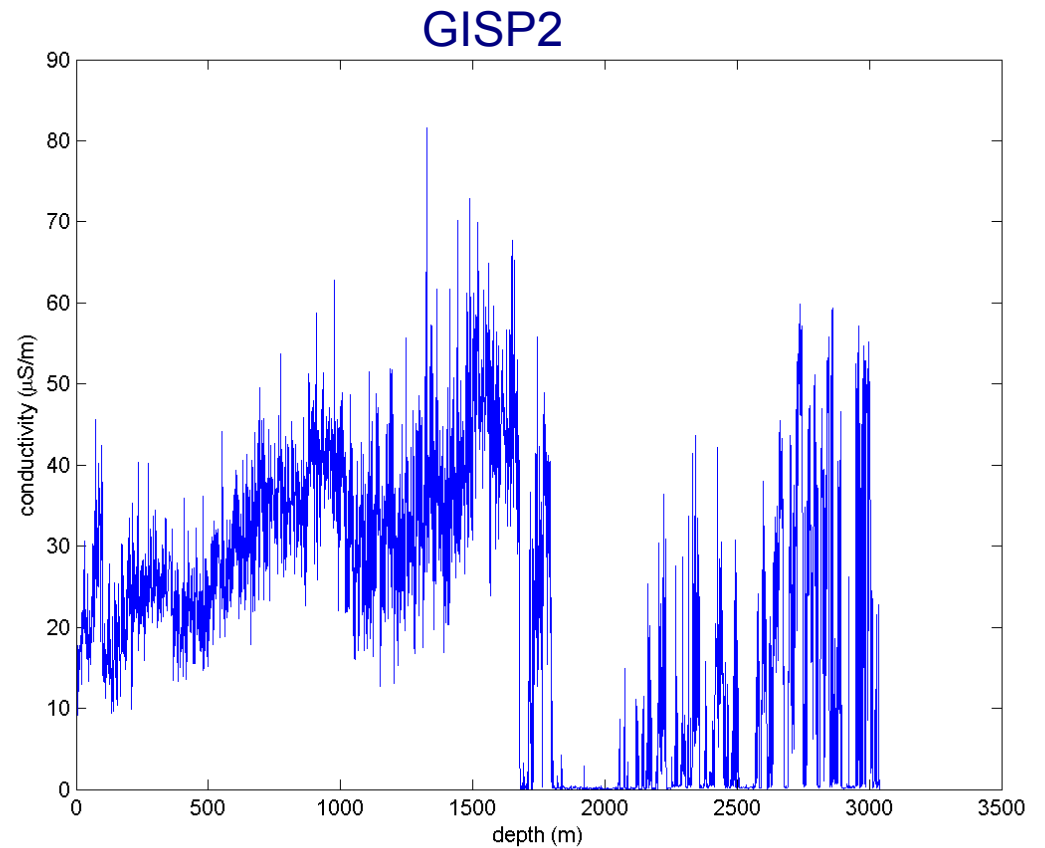
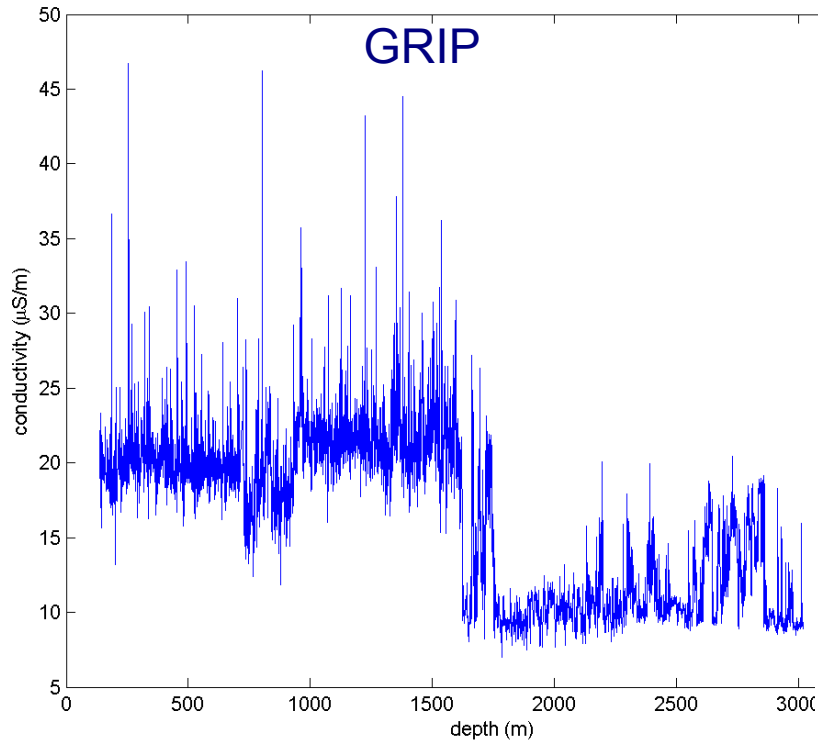
- Funded by NASA
- Sources
  - CReSIS ([data.cresis.ku.edu](http://data.cresis.ku.edu))
  - John Paden for RDS answers
  - Amir Javaid PhD dissertation, U. of Delaware
  - Joseph A MacGregor “Development and Applications of a Radar-Attenuation Model for Polar Ice Sheets”, PhD Dissertation, 2008, U. of Washington



# Extras



# Conductivity Near GRIP



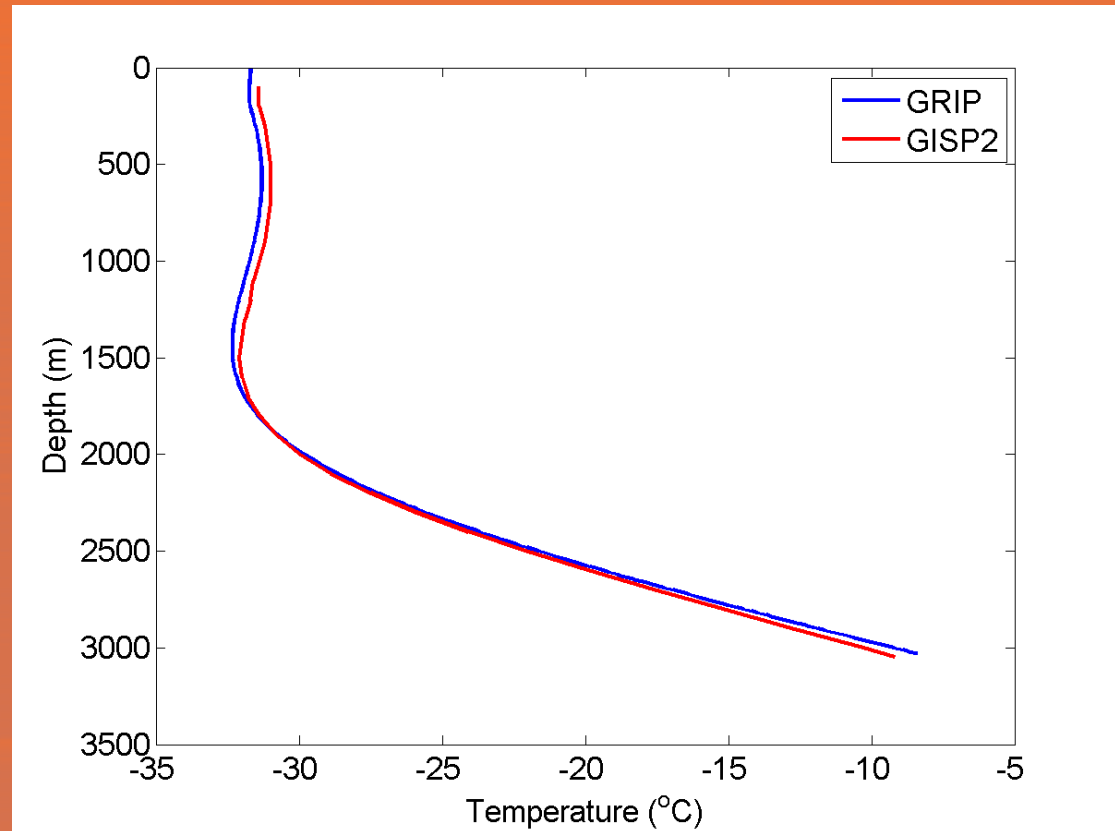


# Comparison of GRIP with GISP2

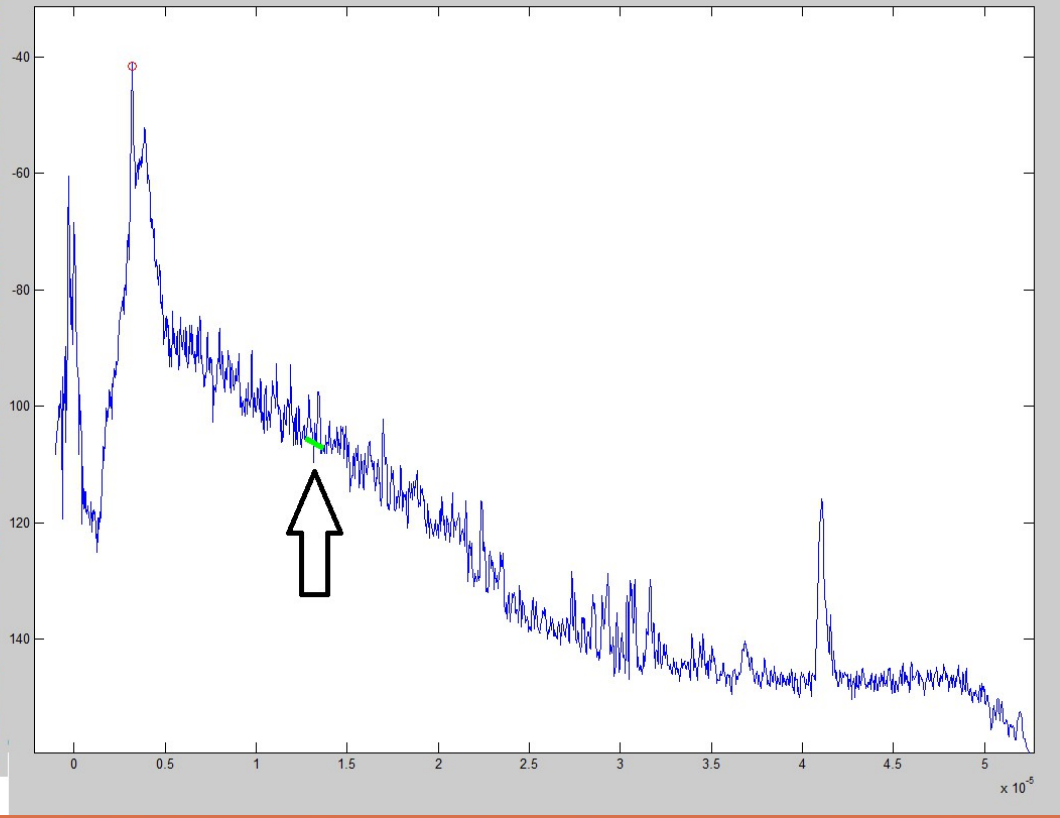
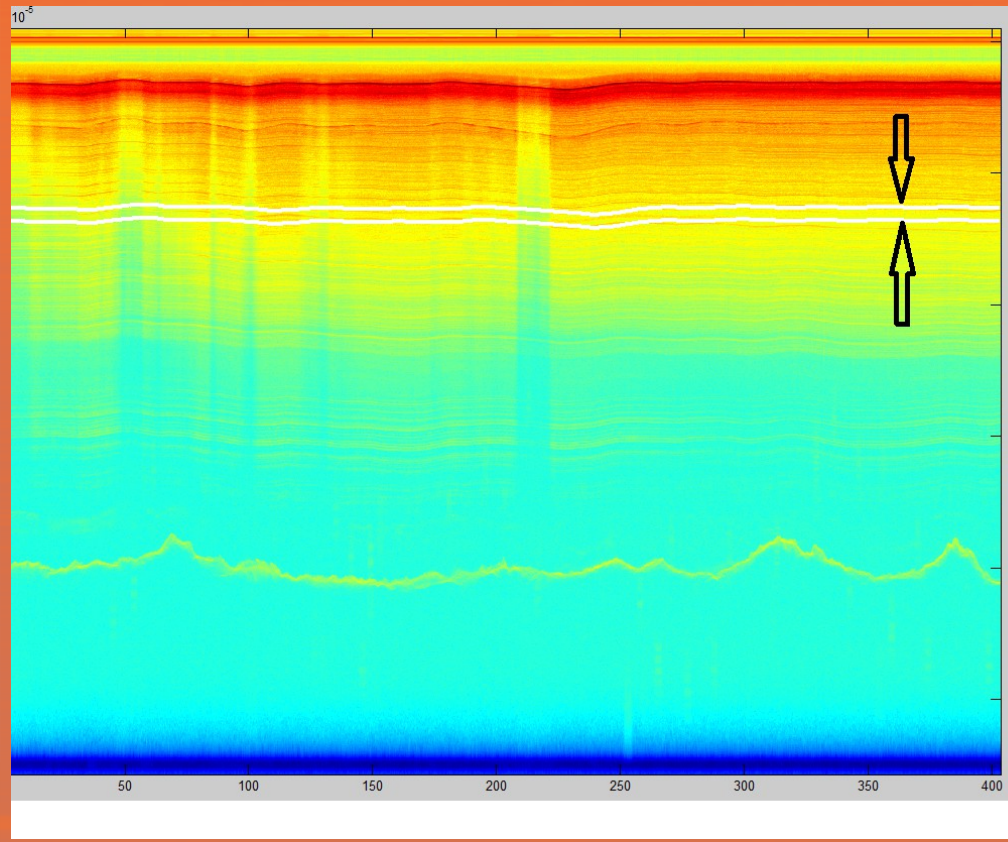
“Several temperature logs and fixed-depth temperature experiments were conducted in the 3-km borehole (GISP-D) at the GISP2 site during 1994, 1995, and 1996. These experiments were conducted using the U.S. Geological Survey's temperature logging system (Clow et al., 1996; Clow and Gundestrup, 1997). The sample spacing for the primary temperature logs is about 20 cm while the accuracy of the temperature measurements is estimated to be 4.5 mK.

However, the temperature logs must be deconvolved to account for the finite response time of the temperature probe, corrected for the thermal disturbance caused by the drilling process, and corrected for the effects of fluid convection within the borehole. As of this time (SEP 1996), we are still developing the fluid convection and drilling disturbance corrections.

The following "uncorrected" data were obtained on 14 June, 1995. Although the quality is insufficient for borehole thermometry studies, it is probably adequate for many other studies (e.g. constraining ice dynamics models). The fully processed data will become available at a later date.” GISP README, ca. 1996



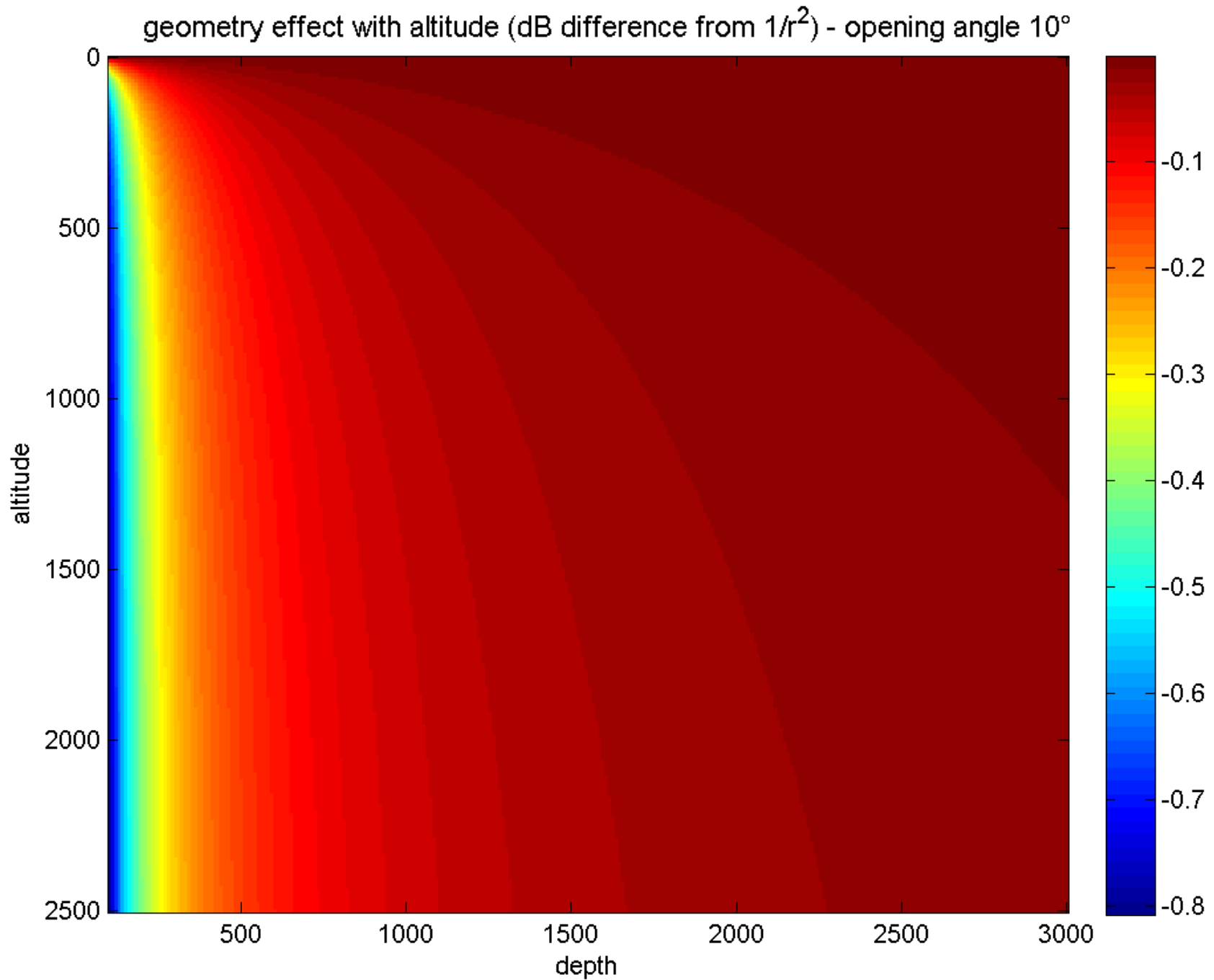
# Quick note <sup>190 MHz</sup> – stitch time

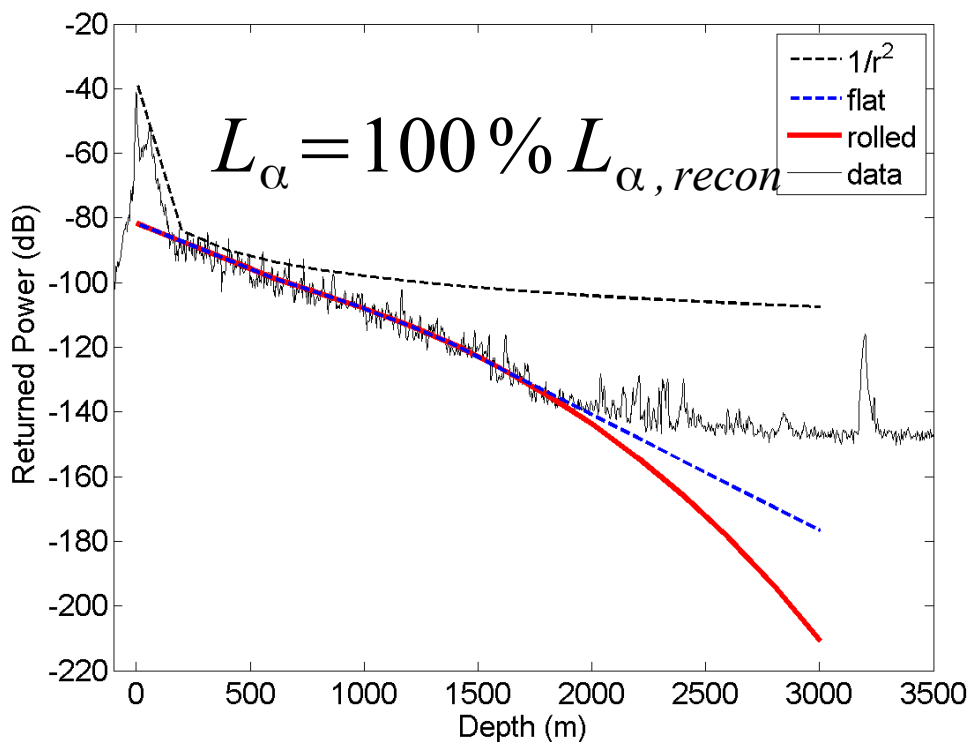
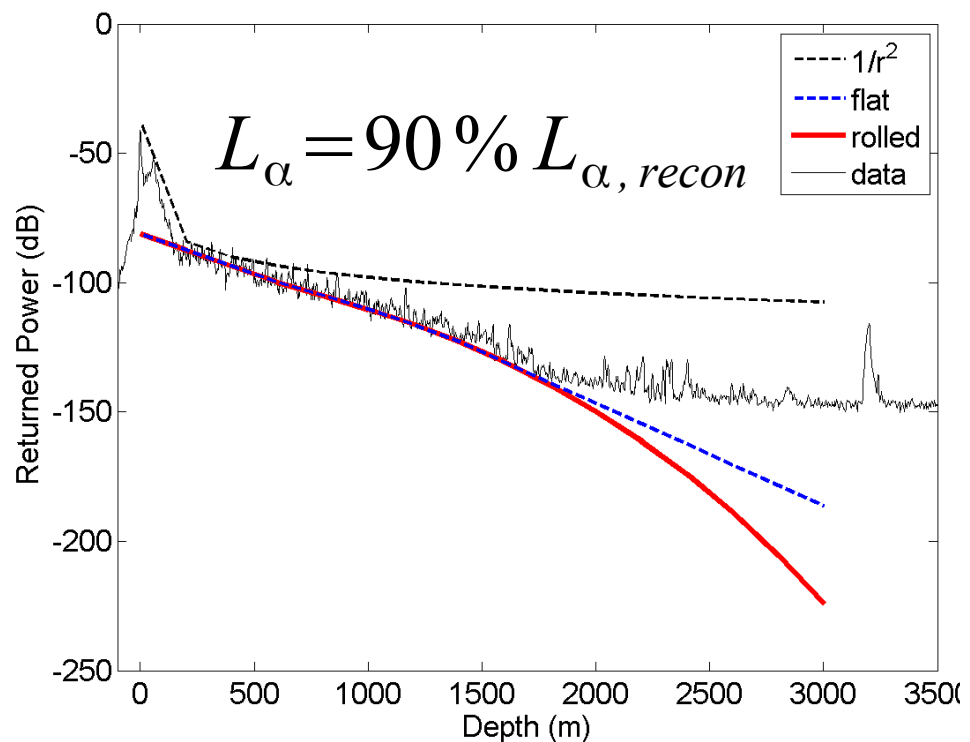
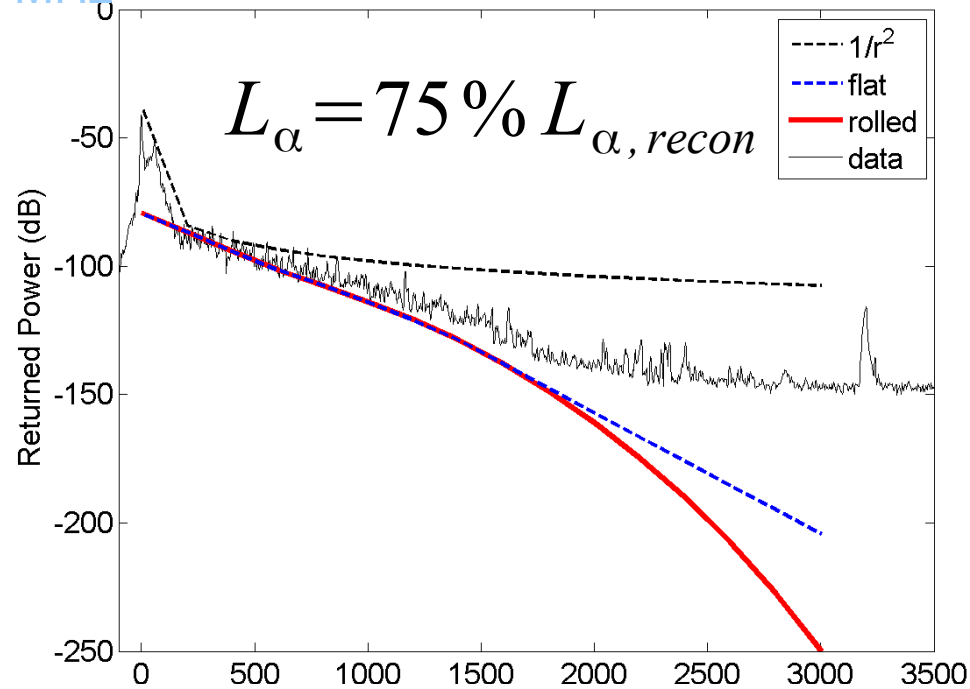
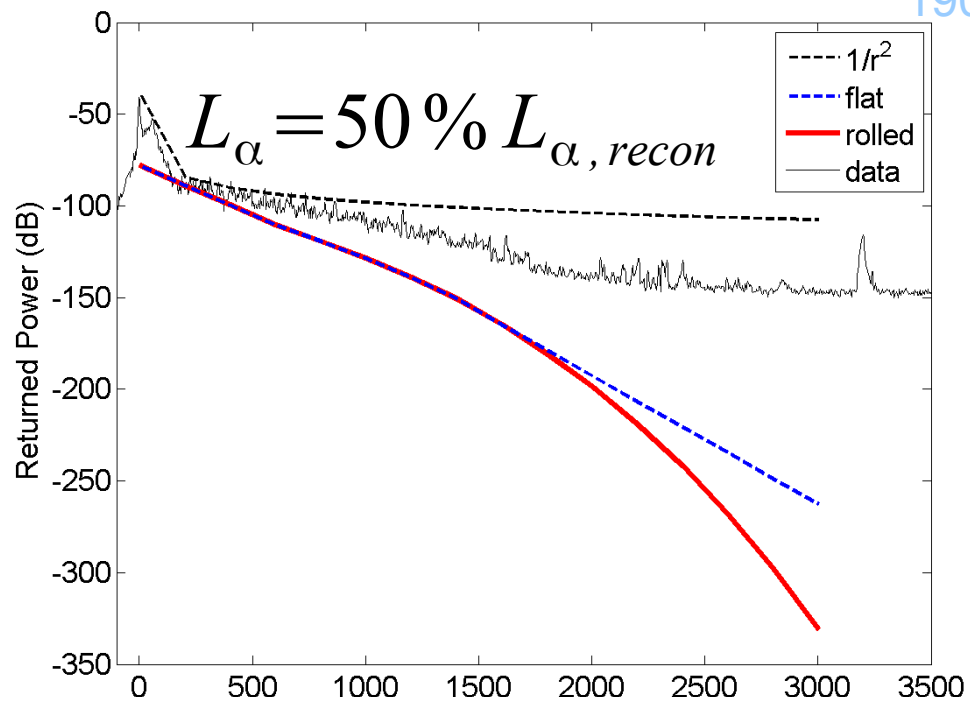


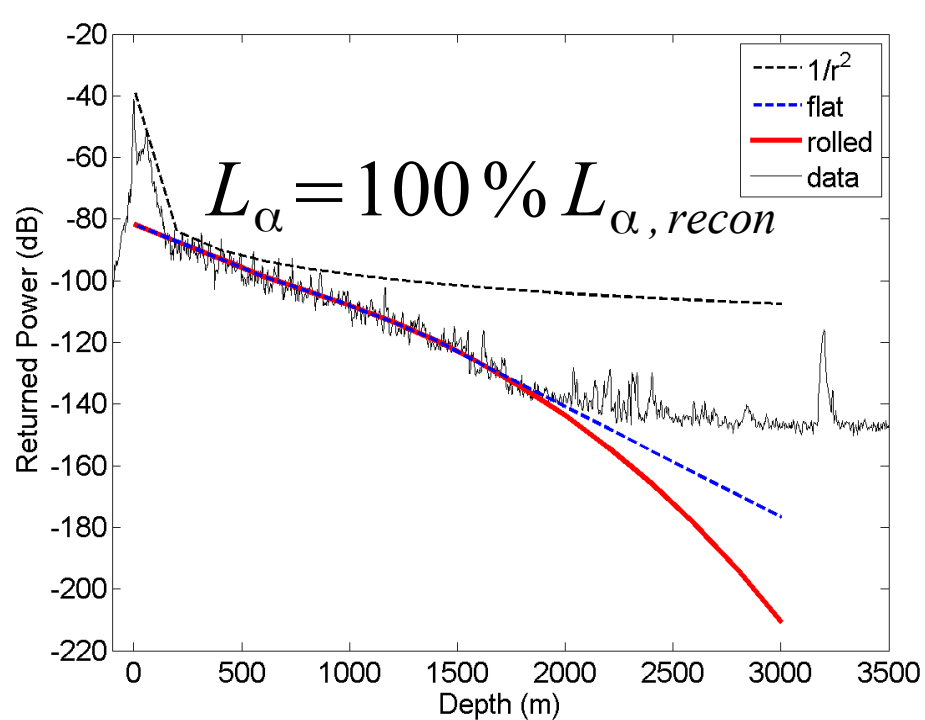
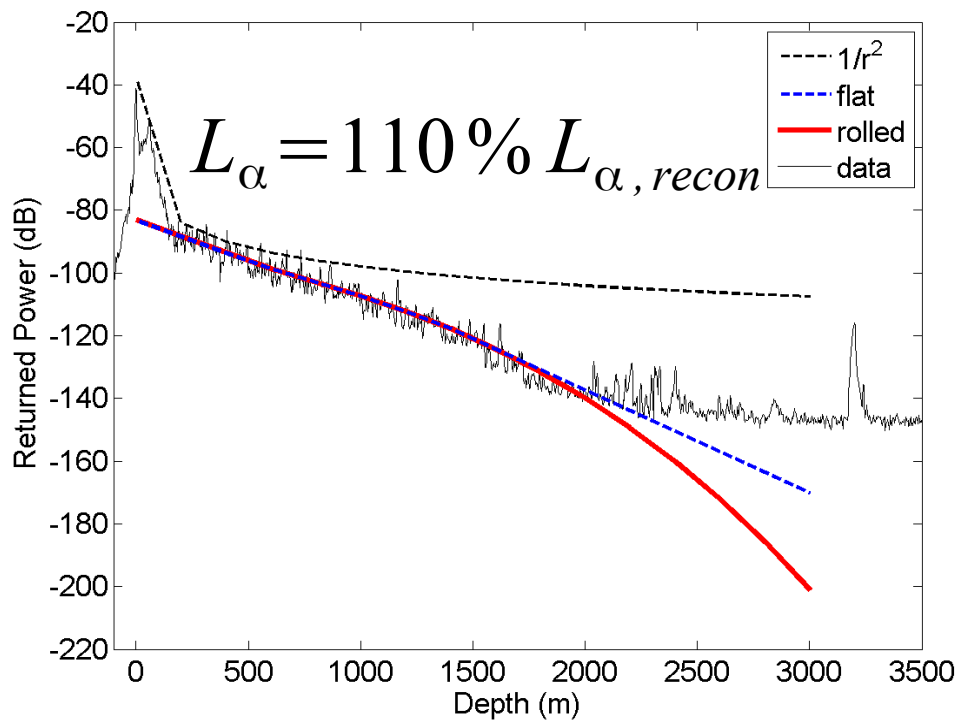
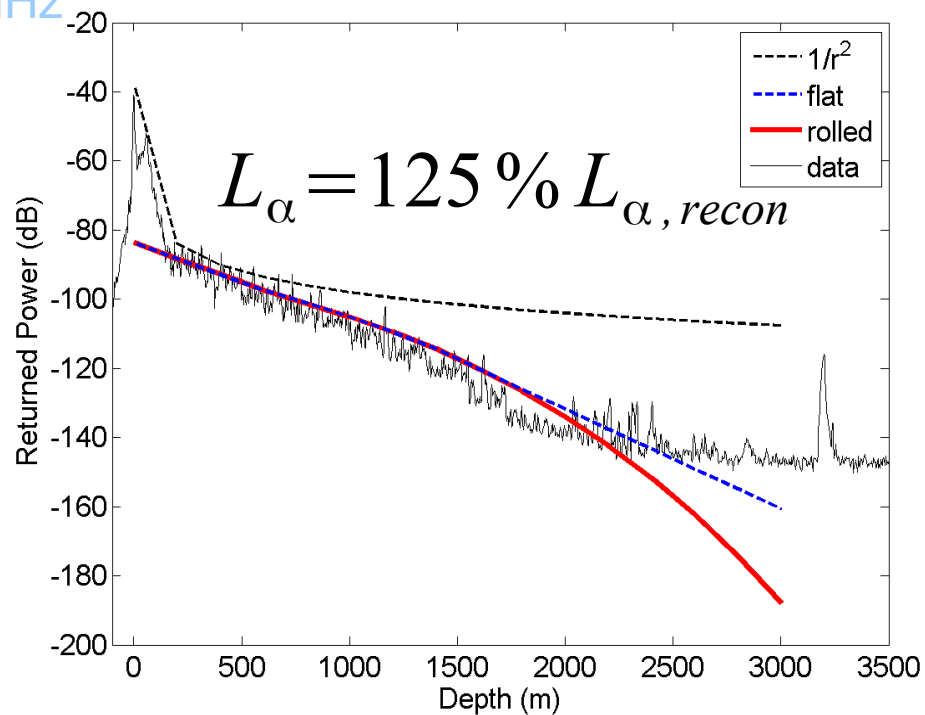
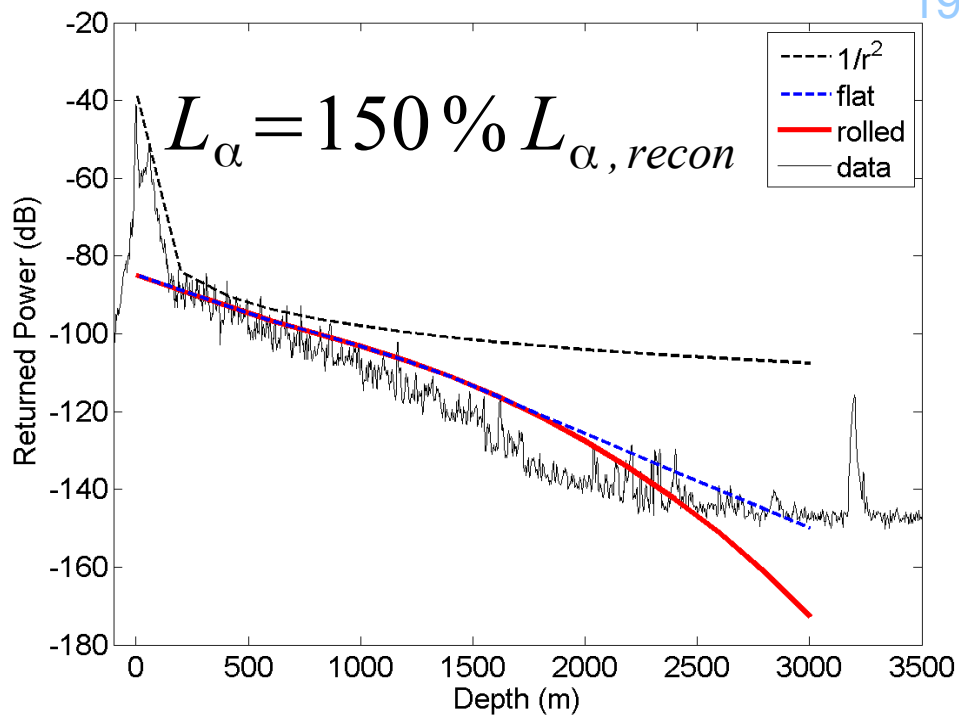
# Constraints on the Method

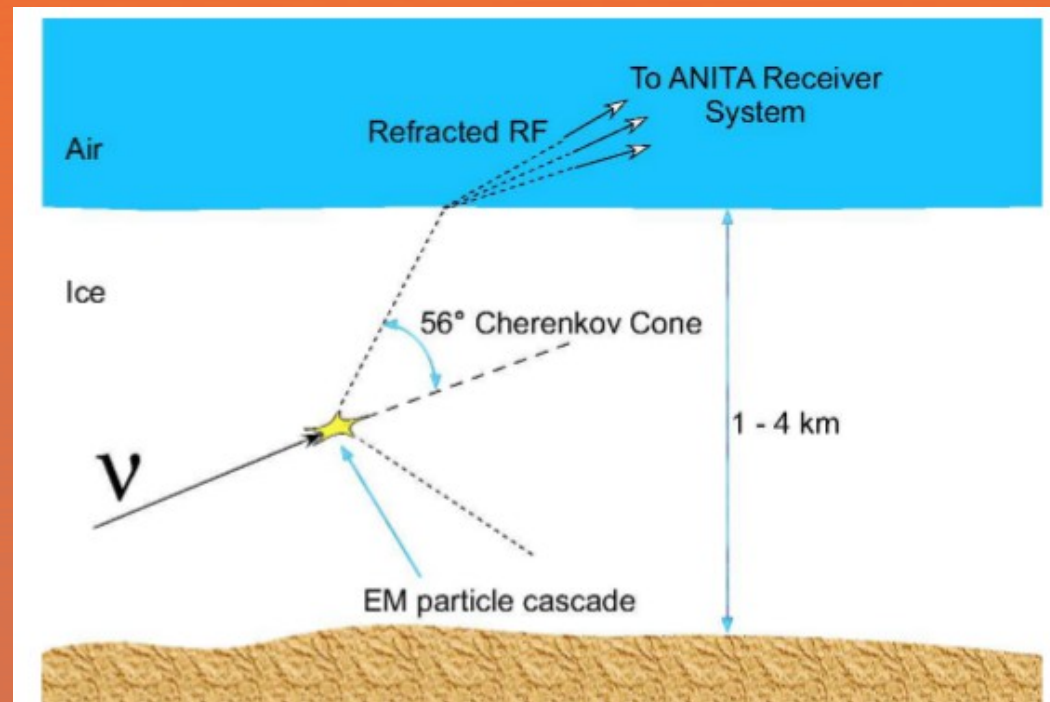
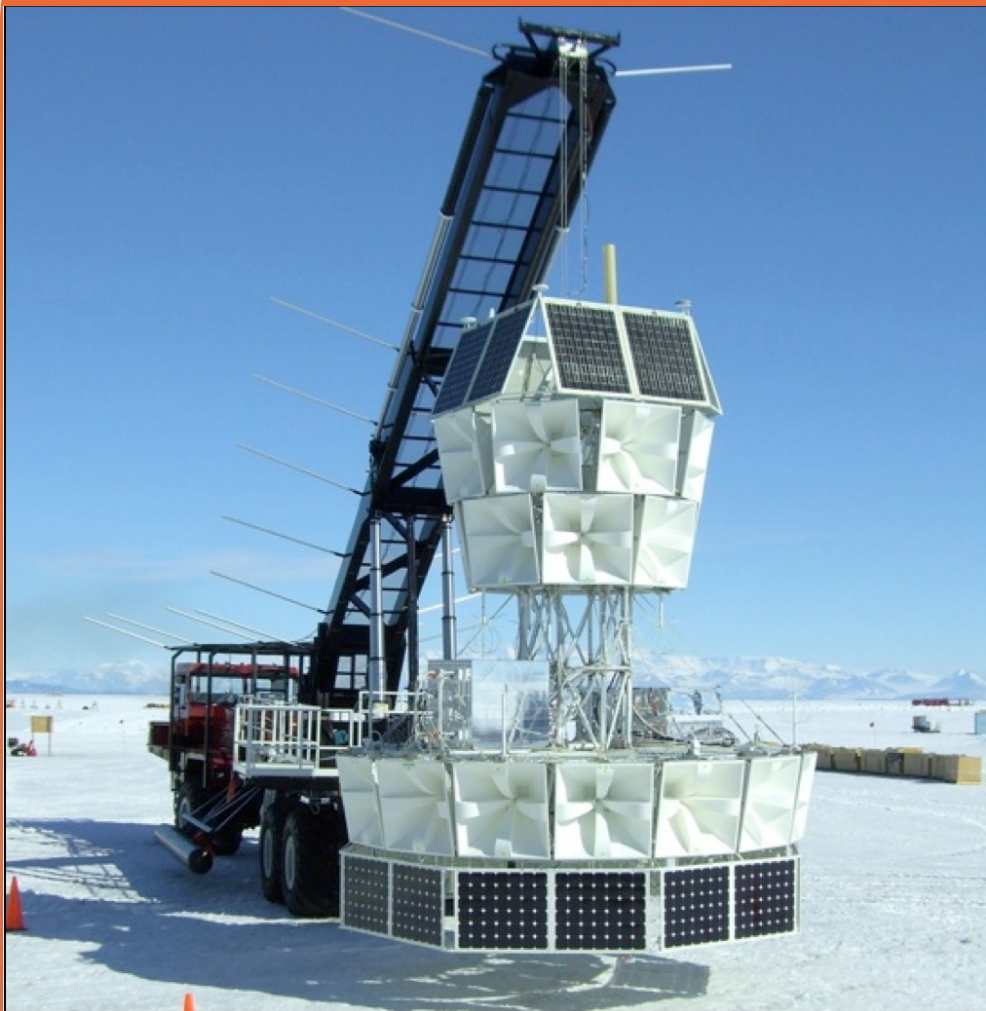
- Needs to compare data that is proximal to have same chemistry and bottom depth (within same echogram ~50km)
- For various DAQ related reasons, more straightforward approaches to measuring the attenuation (e.g. surface and basal reflection strength) are not possible
- What's left? The continuum signal from each individual ping
  - Not Rayleigh scattering – cross section implies -240dB
  - Small dielectric contrasts could do this (bi-annual layers)



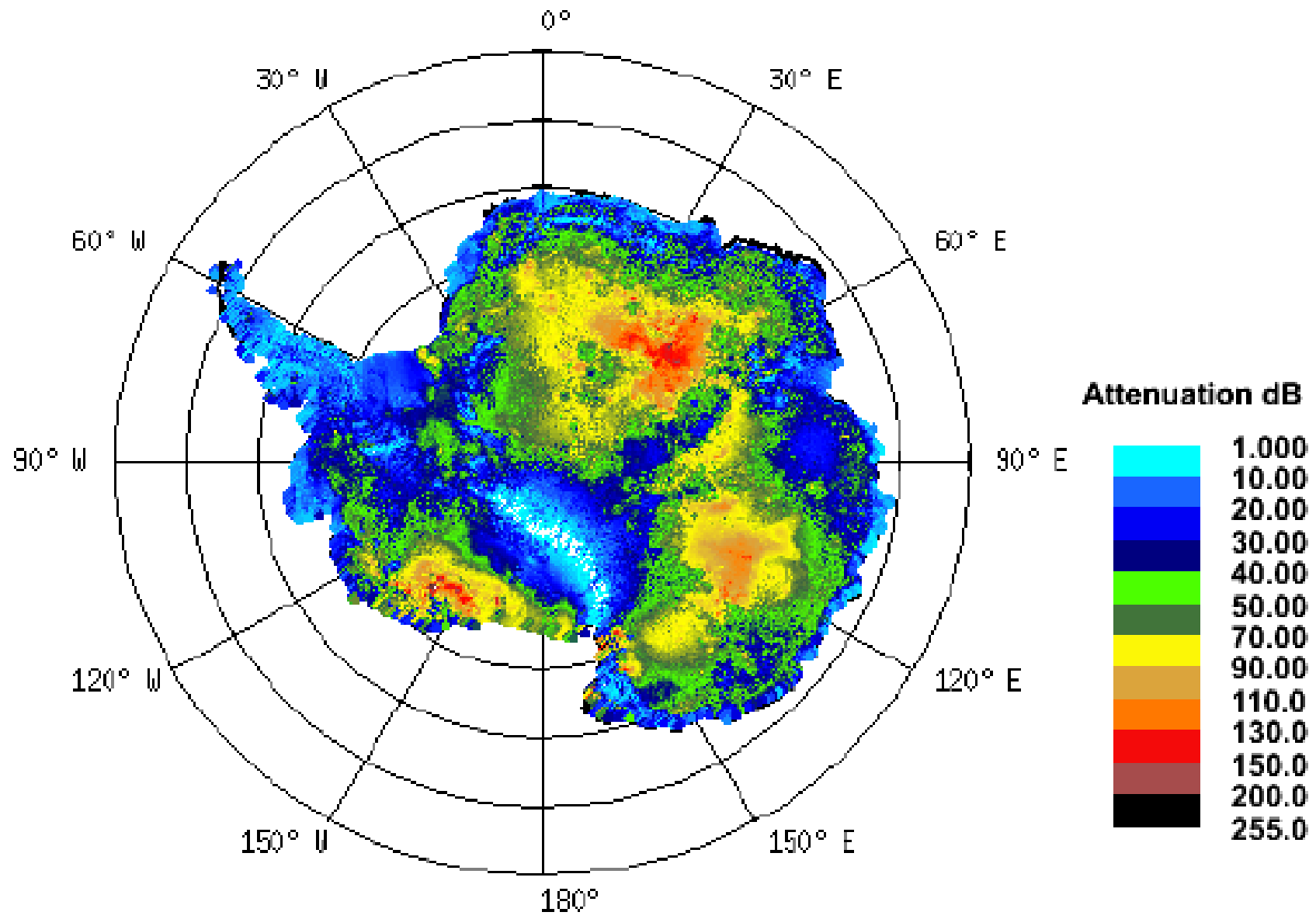








# Total Two-Way Attenuation



Based off Amir Javaid model





