

# Atmospheric temperature and water vapor profiles, transmittance calculations, and radiative transfer

Ralf Bennartz

Atmospheric and Oceanic Sciences Department &  
Cooperative Institute for Meteorological Satellite Studies

University of Wisconsin - Madison

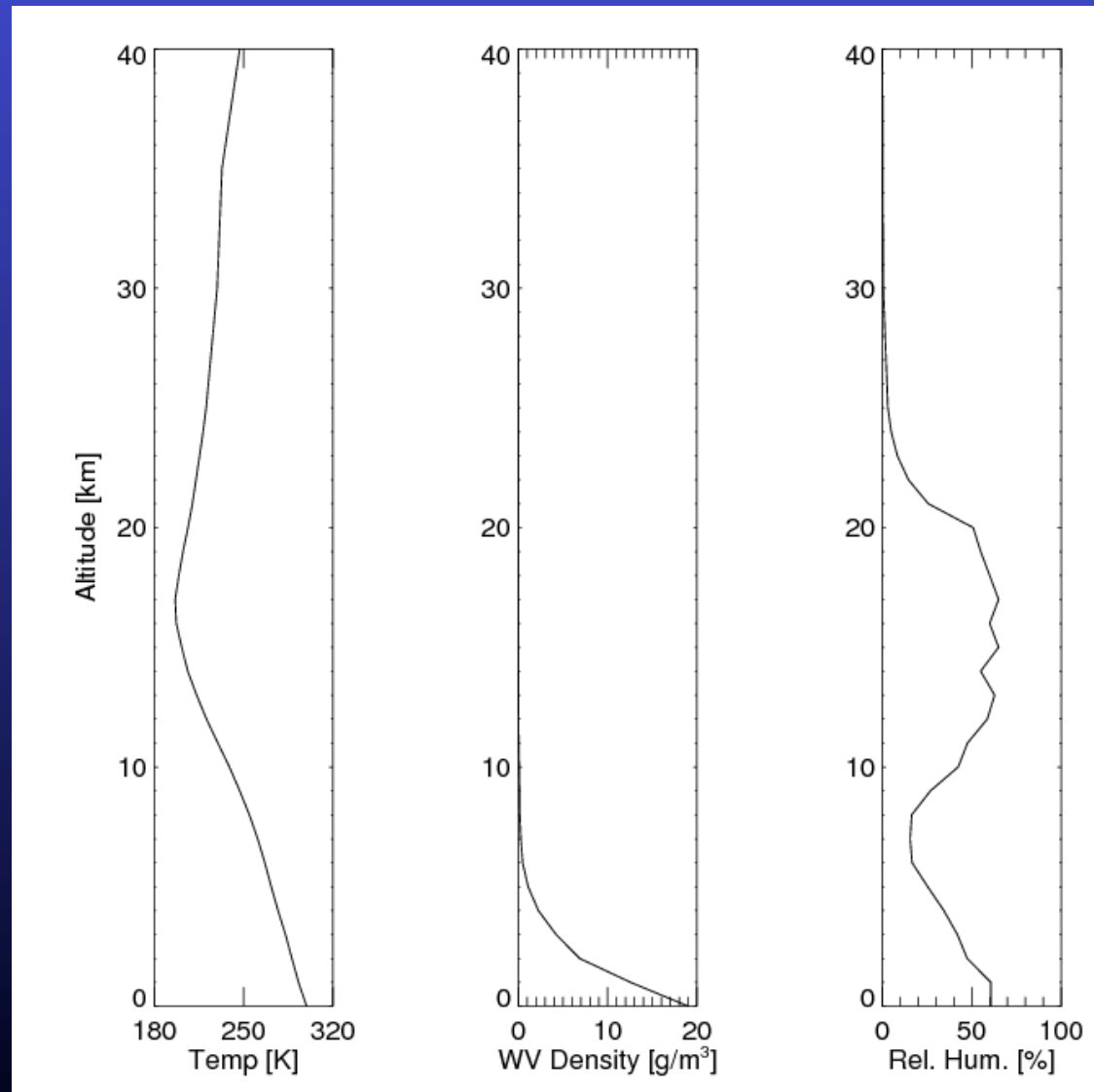
---

## Outline

- Introduction
  - Temperature and water vapor profiles
    - Radiosondes
    - Microwave radiometers
    - Global remote sensing measurements
    - Forecasts/analyses
  - Short-wave radiative transfer
    - Water vapor spectroscopy and retrievals
  - Conclusions
-

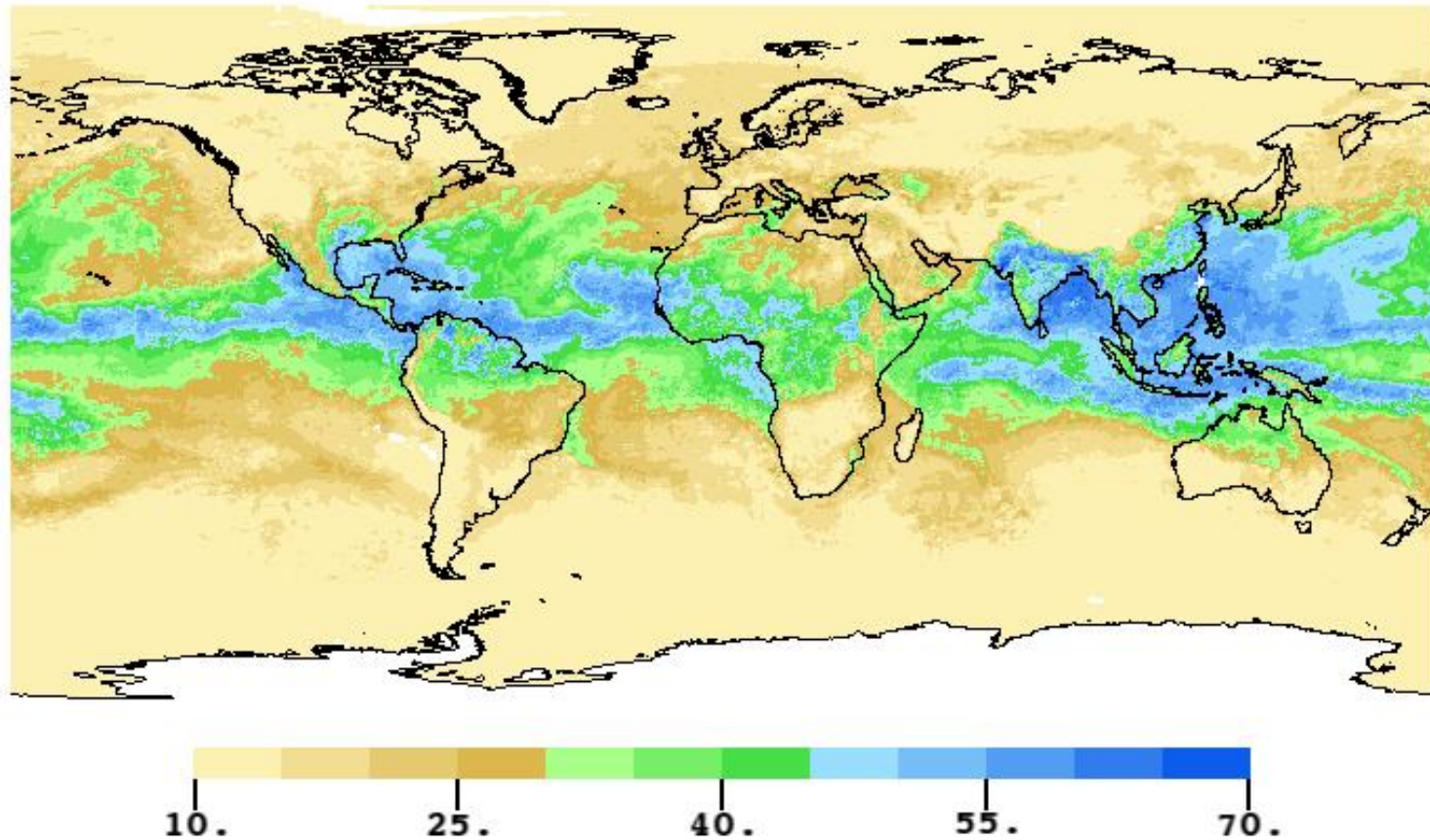
## Atmospheric temperature and moisture profiles

- Pressure scale height roughly 7.5 km
- Water vapor scale height 1-2 km



## Atmospheric temperature and moisture profiles

AIRS TOTAL PRECIPITABLE WATER VAPOR (millimeters) 20100907-20100909



## Observing water vapor and temperature profiles

- Local:
    - Radiosonde
    - Ground-based remote sensing
  - Global:
    - Satellite remote sensing
-

## Radiosonde

- Temperature, pressure, relative humidity as it ascends through the atmosphere
- High vertical resolution
- Small sampling volume
- Requires operator

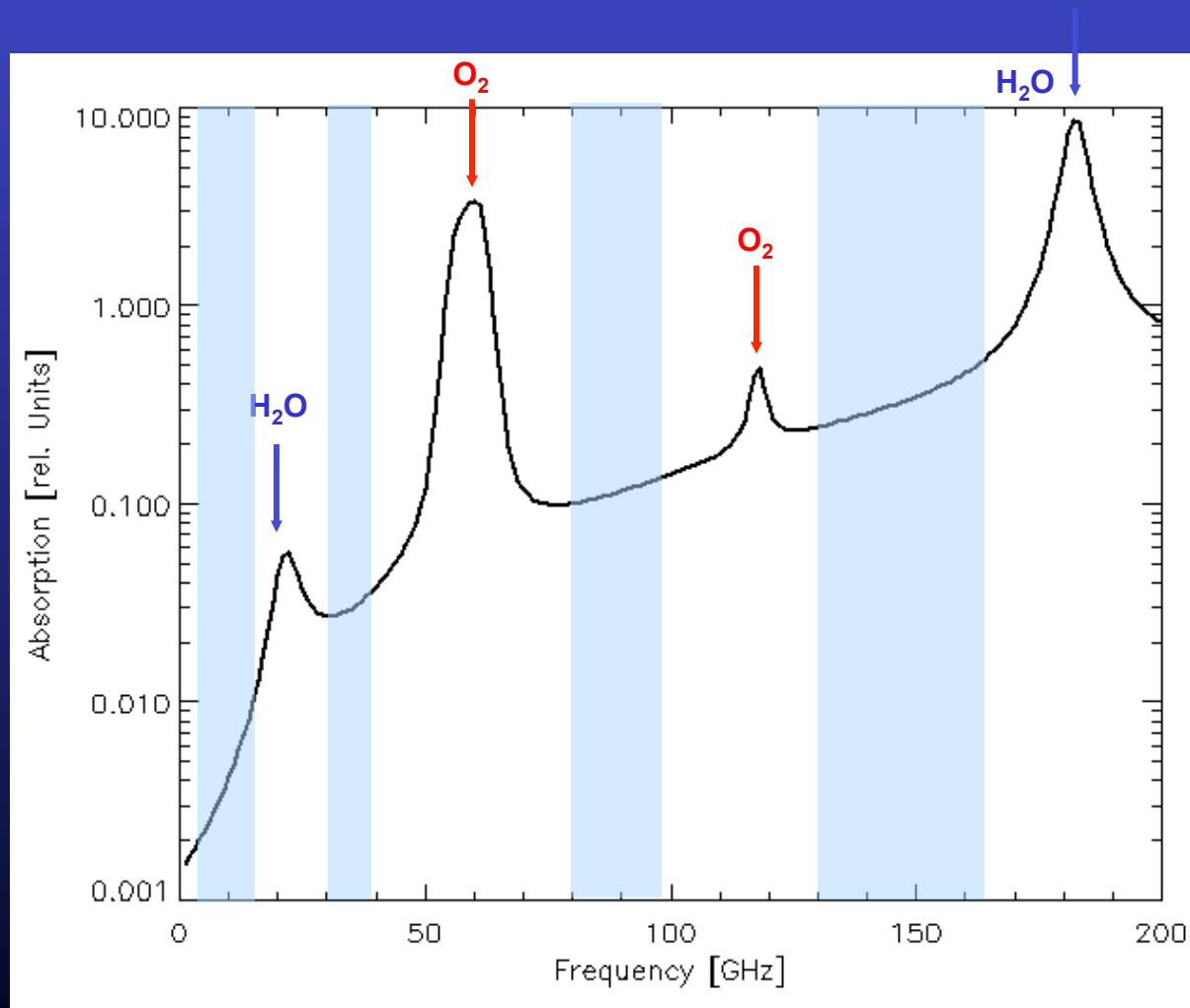


## Ground-based microwave radiometers

- MW-radiances in spectral range 10-200 GHz
  - Lower vertical resolution
  - Runs continuously
  - Runs automatically, if needed
  - High accuracy
- 



## Microwave spectral range



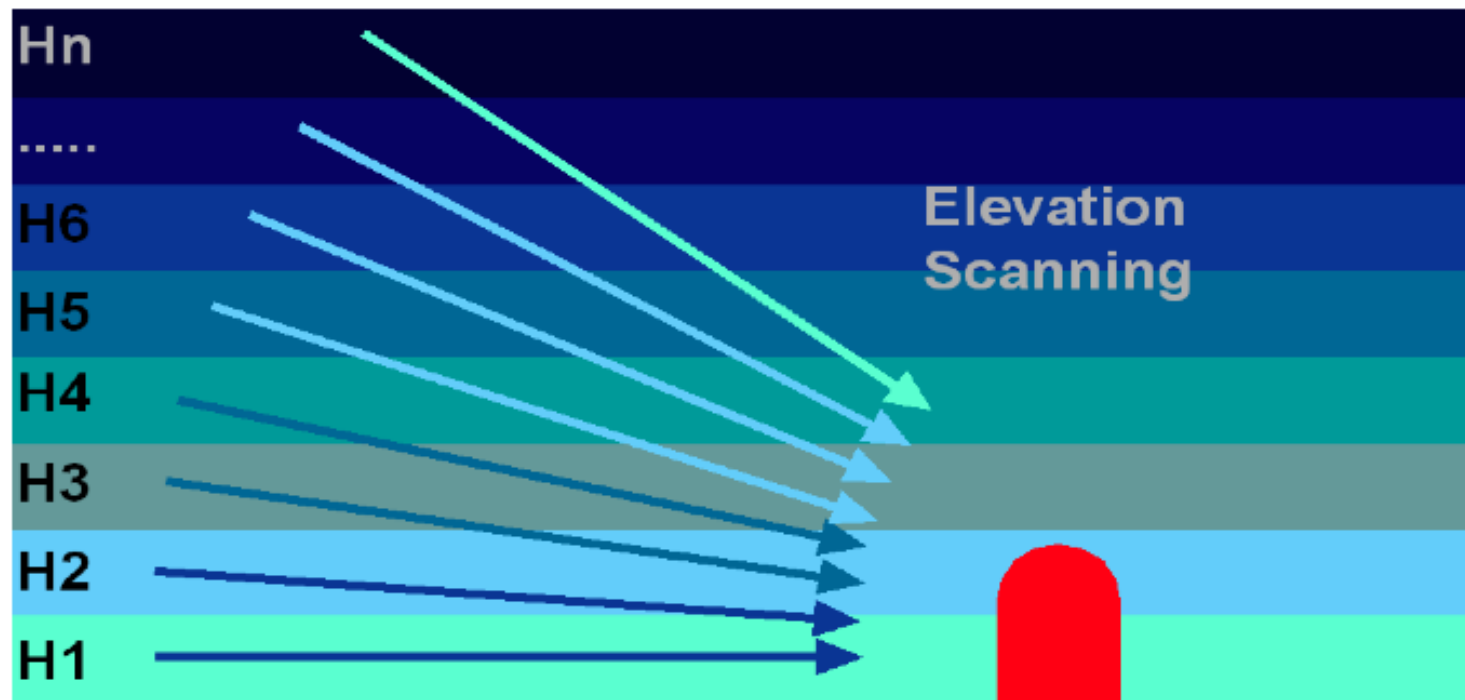


## Ground-based microwave radiometers

- MW-radiances in spectral range 10-200 GHz
- Lower vertical resolution
- Runs continuously
- Runs automatically, if needed

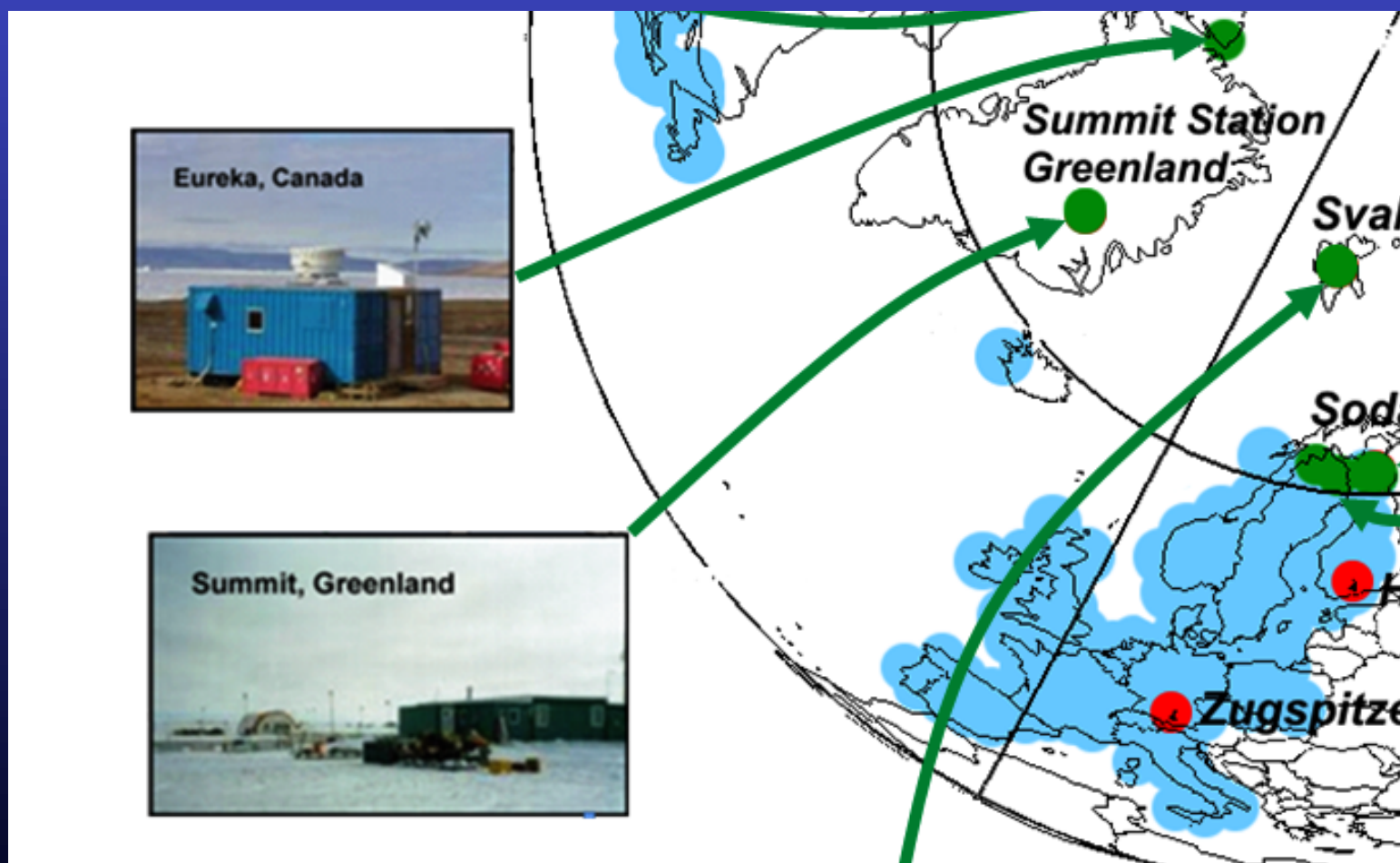


## Ground-based microwave radiometers

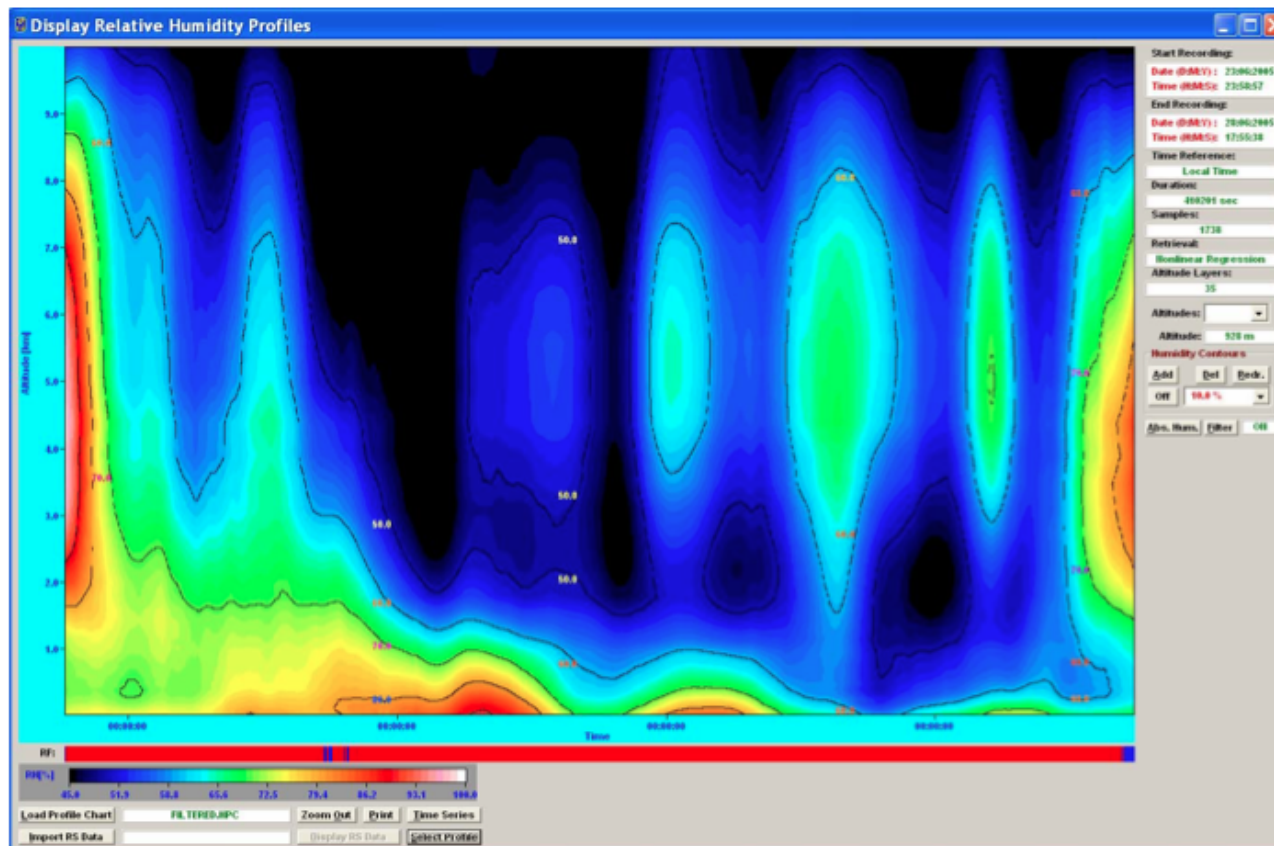


**Fig.2: Boundary layer scanning mode with different elevation angles.**

## Ground-based microwave radiometers



# Ground-based microwave radiometers



**Fig.11: Relative humidity map for the full troposphere (up to 10 km) computed from absolute humidity profiles and temperature profiles both measured in zenith mode.**

## Non-scattering radiative transfer theory (Schwarzschild 1906)

$$\frac{dI}{ds} = \underbrace{-\beta_A I}_{\text{Attenuation}} + \underbrace{\beta_A B}_{\text{Thermal Emission}}$$

$I$  : Radiance

$B$  : Planck-function

$\beta_A$  : Absorption cross-section per unit volume [ $\text{m}^2/\text{m}^3$ ]  
a.k.a volume absorption coefficient

---

## Non-scattering radiative transfer theory

$$I^\downarrow(0) = I_c t(0, \infty) + \frac{1}{\mu} \int_0^{TOA} B(z) \beta_A(z) t(0, z) dz$$

$$\tau(0, z) = \int_0^z \beta_A(z) dz \quad : \quad \text{Optical depth between } (0, z)$$

$$t(0, z) = e^{-\tau(0, z)/\mu} \quad : \quad \text{Transmission between } (0, z)$$

$$\mu = |\cos(\theta)| \quad : \quad \text{Cosine zenith angle}$$

---

## Non-scattering radiative transfer theory

$$I^\downarrow(0) = I_c t(0, \infty) + \frac{1}{\mu} \int_0^{TOA} B(z) \beta_A(z) t(0, z) dz$$

$$W^\downarrow(z) = \frac{1}{\mu} \beta_A(z) t(0, z) : \text{Emission weighting function}$$

$$\tau(0, z) = \int_0^z \beta_A(z) dz \quad : \quad \text{Optical depth between } (0, z)$$

$$t(0, z) = e^{-\tau(0, z)/\mu} \quad : \quad \text{Transmission between } (0, z)$$

$$\mu = |\cos(\theta)| \quad : \quad \text{Cosine zenith angle}$$

## Non-scattering radiative transfer theory

$$I^\downarrow(0) = I_c t(0, \infty) + \frac{1}{\mu} \int_0^{TOA} B(z) \beta_A(z) t(0, z) dz$$

$$W^\downarrow(z) = \frac{1}{\mu} \beta_A(z) t(0, z) : \text{Emission weighting function}$$

$$I^\downarrow(0) = \int_0^{TOA} B(z) W^\downarrow(z) dz$$

$$\tau(0, z) = \int_0^z \beta_A(z) dz \quad : \quad \text{Optical depth between } (0, z)$$

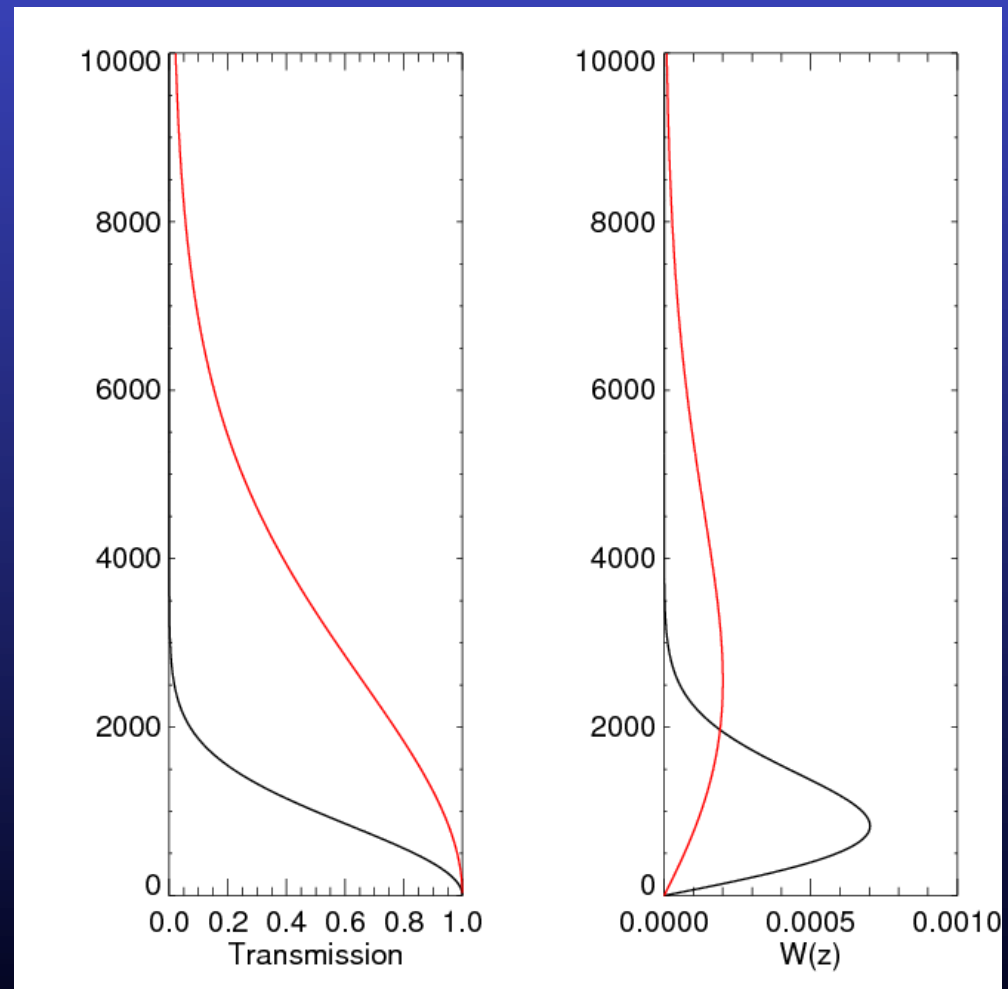
$$t(0, z) = e^{-\tau(0, z)/\mu} \quad : \quad \text{Transmission between } (0, z)$$

$$\mu = |\cos(\theta)| \quad : \quad \text{Cosine zenith angle}$$



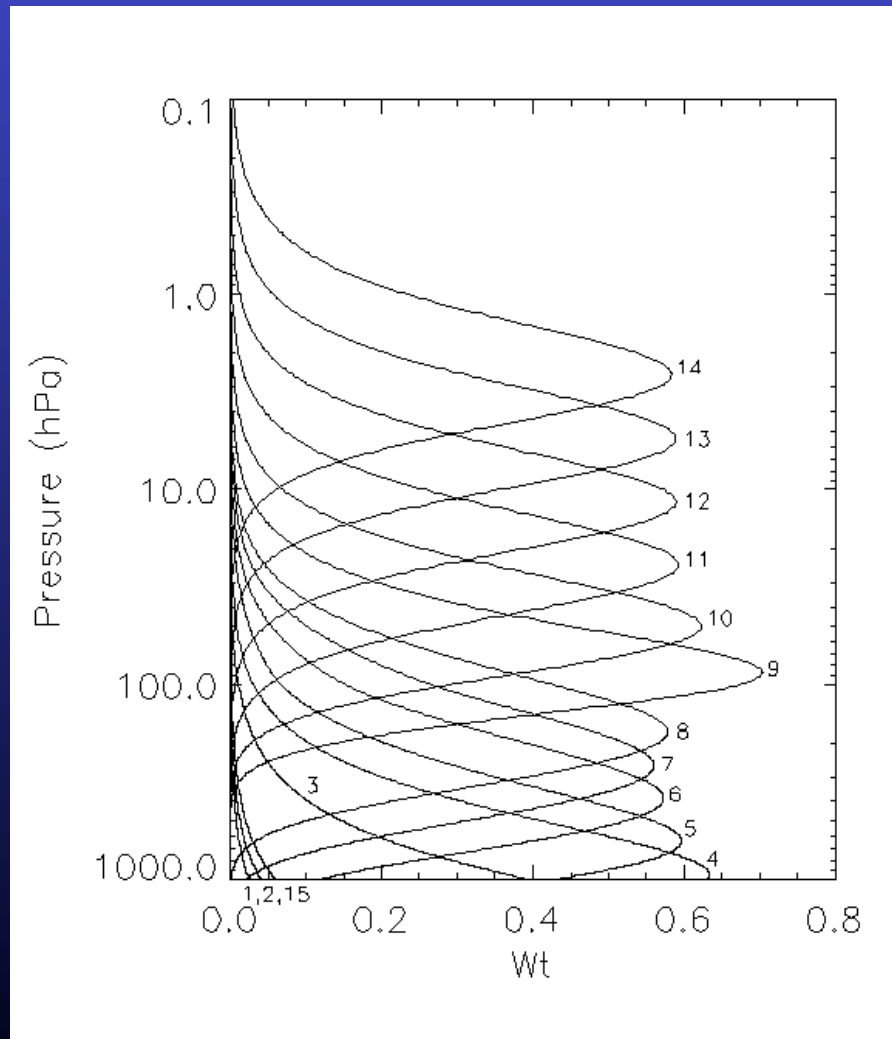
## Schematic example of weighting functions

- Black: More strongly absorbing wavelength
- Red: less strongly absorbing wavelength



## Actual microwave weighting function example

- E.g. 15 Channel instrument observing  $O_2$ -absorption band at 50-60 GHz



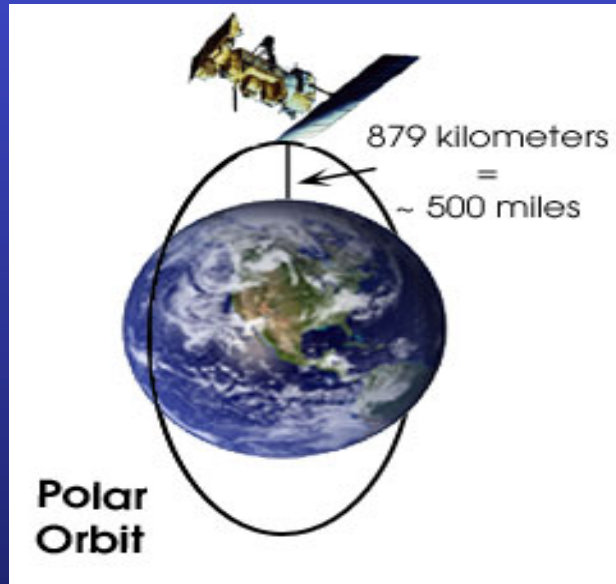
# Observing water vapor and temperature profiles

- Local:
    - Radiosonde
    - Ground-based remote sensing
  - Global:
    - Satellite remote sensing
-

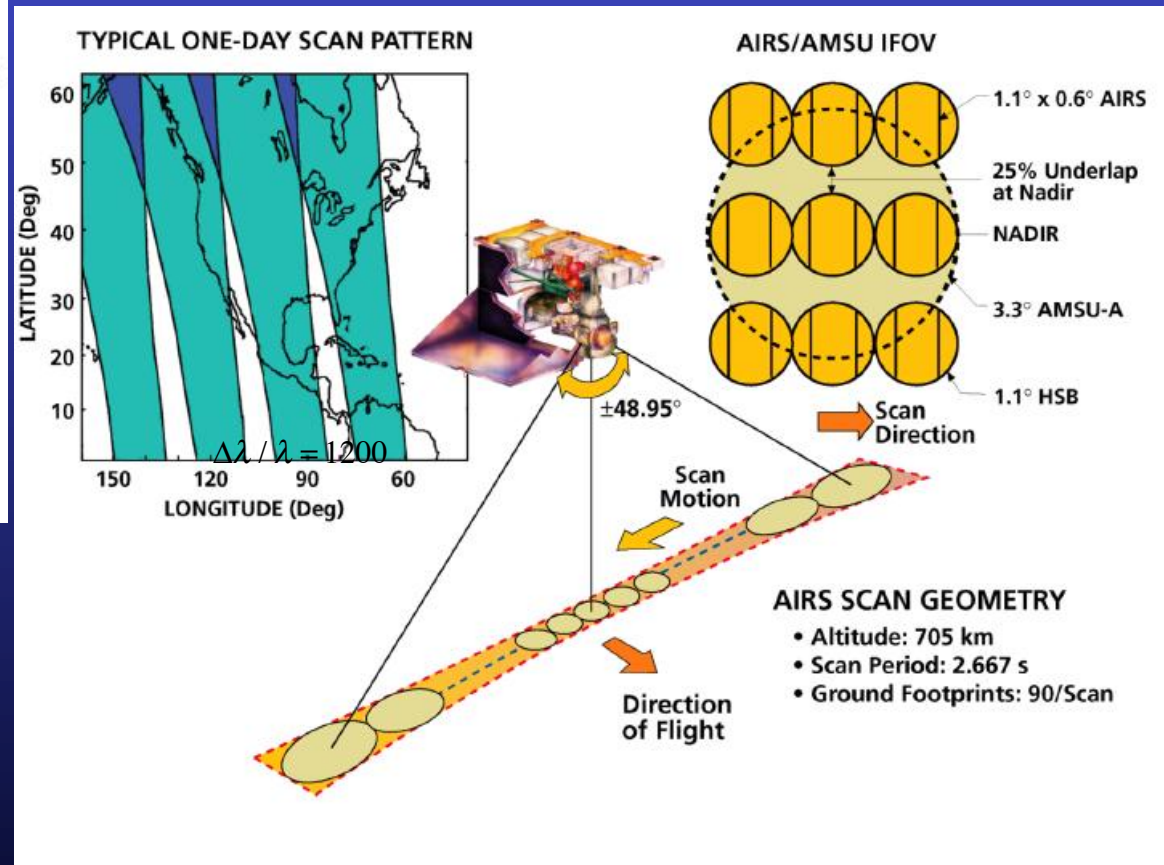
# Observing water vapor and temperature profiles

- Local:
    - Radiosonde
    - Ground-based remote sensing
  - Global:
    - Satellite remote sensing
-

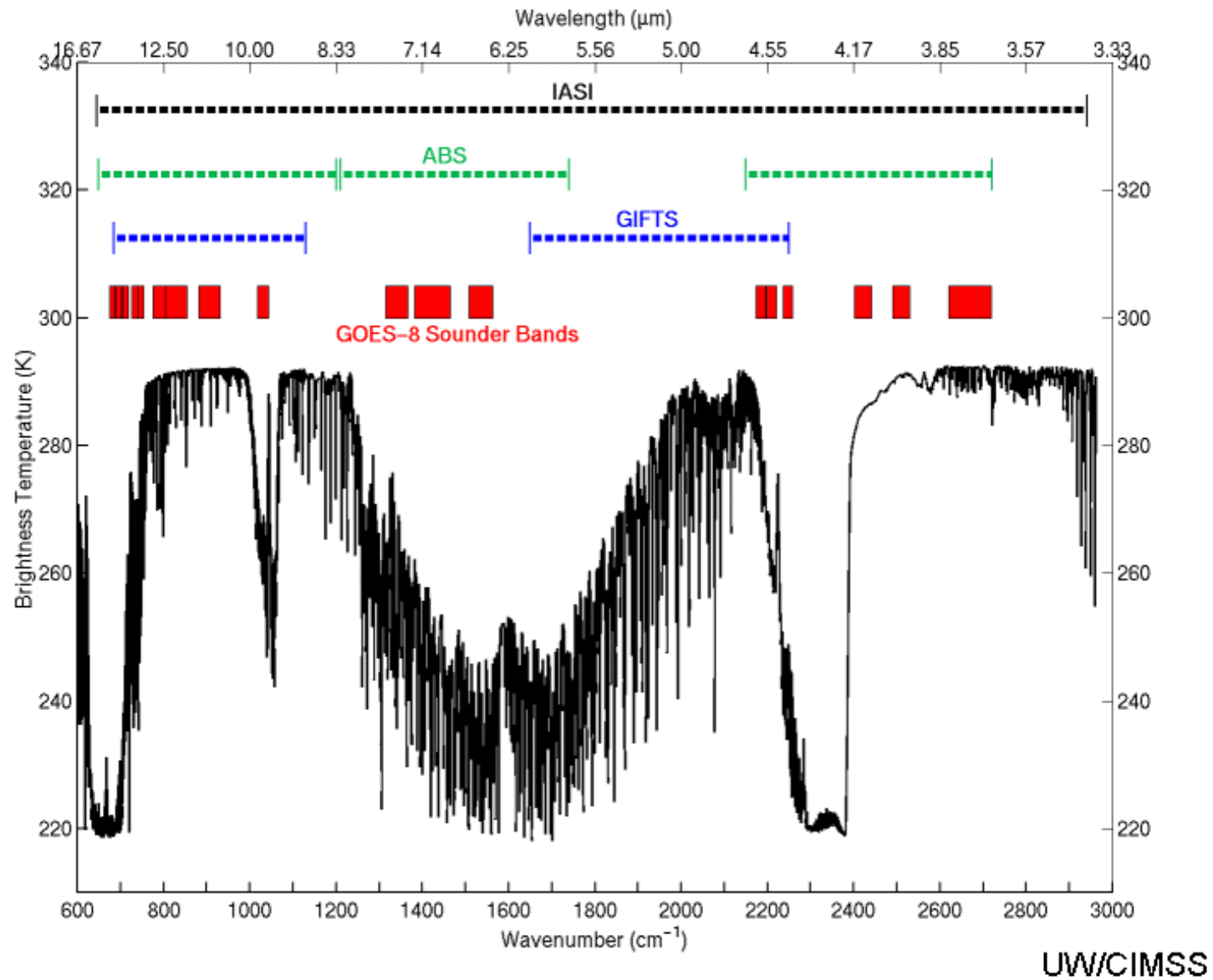
## Example: Atmospheric InfraRed Sounder (AIRS)



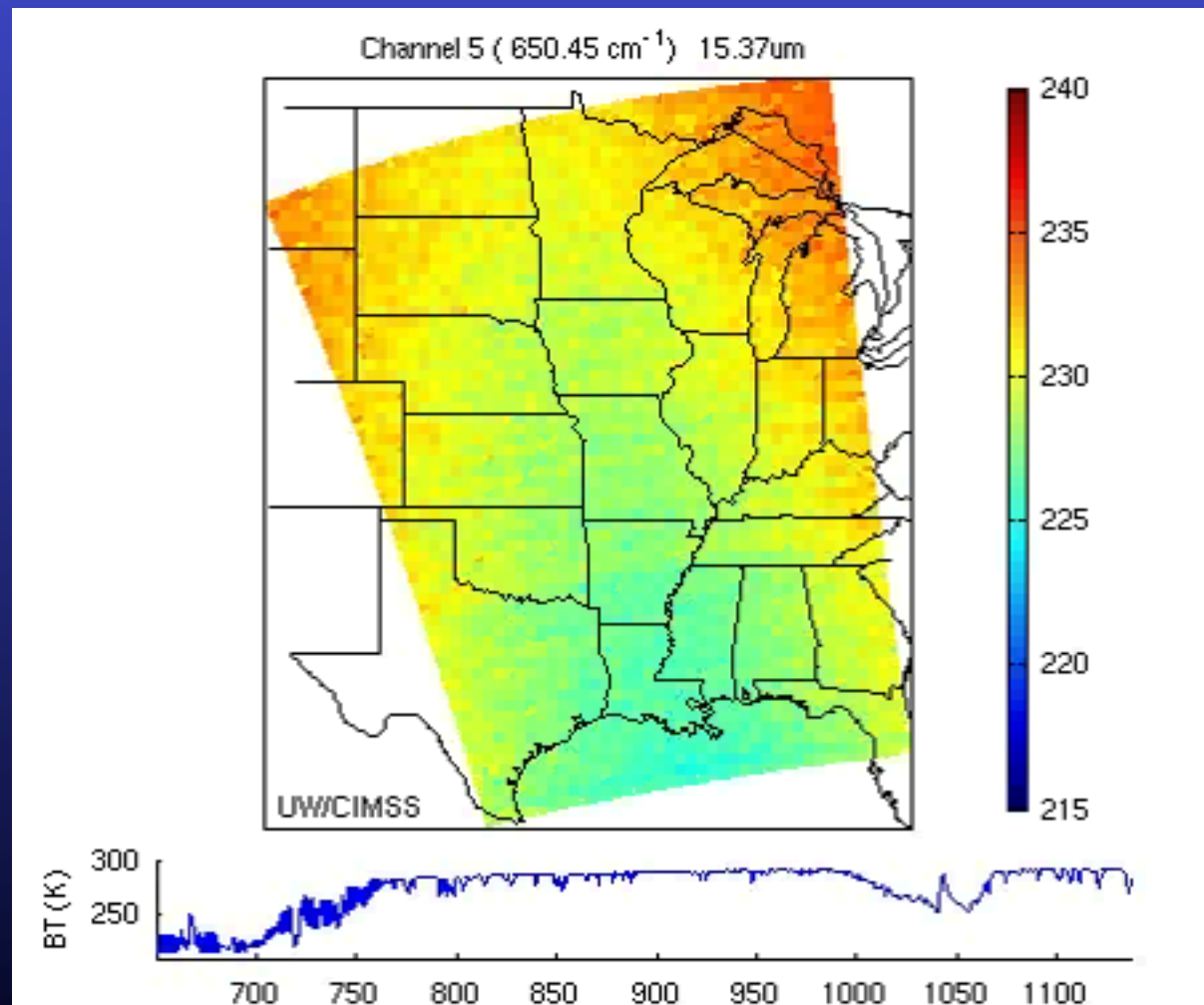
- 2378 channels
- 3.7-15.4 micron
- $dL/L=1200$



## Example: Atmospheric InfraRed Sounder (AIRS)



## Example: Atmospheric InfraRed Sounder (AIRS)

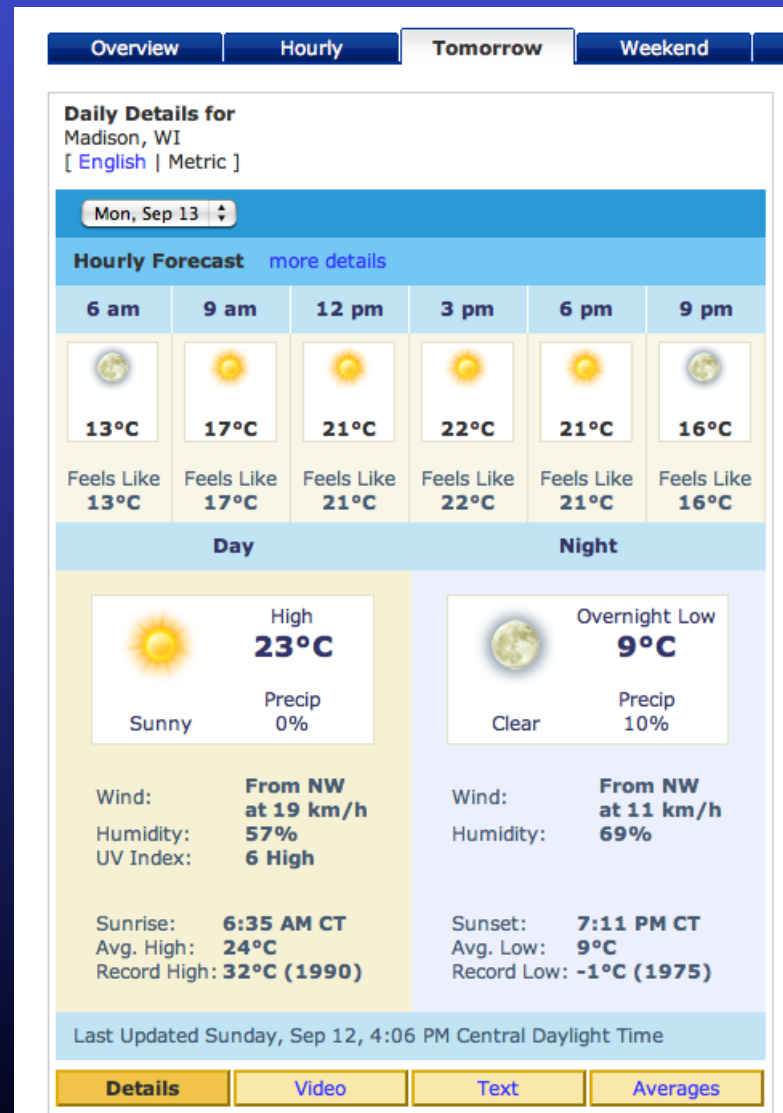


## Example: Atmospheric InfraRed Sounder (AIRS)

- Provides data since mid 2002
  - Temperature and water vapor profiles available globally about once every 12hours
  - Spatial resolution about 15 km horizontally
  - Accuracy T :  $\pm 1\text{K}$  @  $\text{dz}=1\text{km}$
  - Accuracy RH :  $\pm 10\%$  @  $\text{dz}=1\text{km}$
  - Other, similar instruments out there. Data continuity high priority for NOAA and EUMETSAT
-

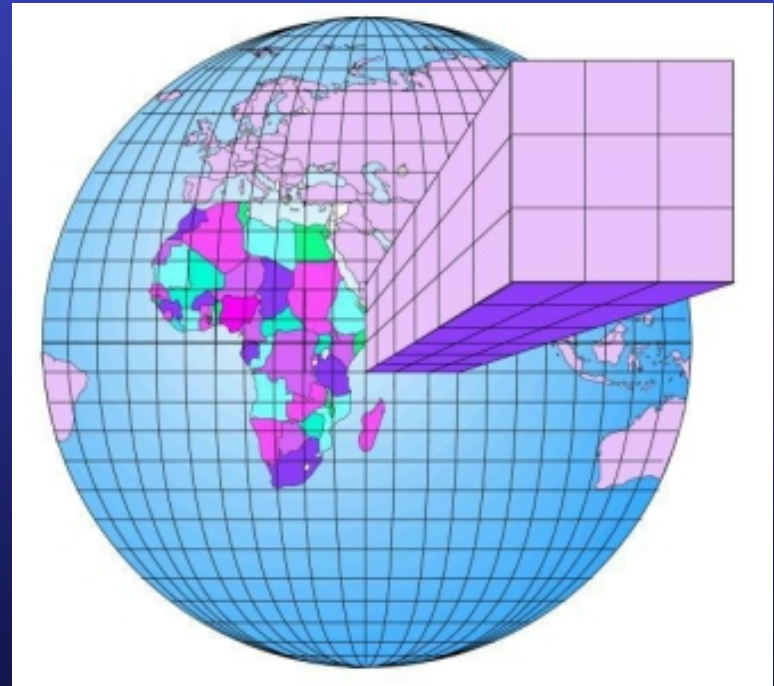


# Numerical Weather Prediction Models



## Numerical Weather Prediction Models

- Integrates Navier Stokes equations forward in time on a rotating spherical grid.
- Variables:  $T$ ,  $p$ ,  $q$ ,  $(u,v,w)$
- Examples:
  - NCEP's Global Forecasting System (GFS, GDAS)
  - ECMWF's Integrated Forecasting System (IFS)



## Numerical Weather Prediction Models

- Initial conditions: Every single grid box needs to be initialized with values of  $T, q, (u, v, w)$  before the equation can be integrated forward (i.e. a 'forecast' can be made).



## Numerical Weather Prediction Models

- Initial conditions: Every single grid box needs to be initialized with values of  $T, q, (u, v, w)$  before the equation can be integrated forward (i.e. a 'forecast' can be made).
  - The initial state of the model is called **analysis**.
-

## Numerical Weather Prediction Models

- Initial conditions: Every single grid box needs to be initialized with values of  $T, q, (u, v, w)$  before the equation can be integrated forward (i.e. a 'forecast' can be made).
  - The initial state of the model is called **analysis**.
  - However, measurements are not available everywhere (in fact most everywhere there are no measurements).
-

## Numerical Weather Prediction Models

- Initial conditions: Every single grid box needs to be initialized with values of  $T, q, (u, v, w)$  before the equation can be integrated forward (i.e. a 'forecast' can be made).
  - The initial state of the model is called **analysis**.
  - However, measurements are not available everywhere (in fact most everywhere there are no measurements).
  - How to incorporate the sparse observations into an optimal analysis
-

## Numerical Weather Prediction Models

- Initial conditions: Every single grid box needs to be initialized with values of  $T, q, (u, v, w)$  before the equation can be integrated forward (i.e. a 'forecast' can be made).
  - The initial state of the model is called **analysis**.
  - However, measurements are not available everywhere (in fact most everywhere there are no measurements).
  - How to incorporate the sparse observations into an optimal analysis → **Data assimilation**.
-

## Data Assimilation

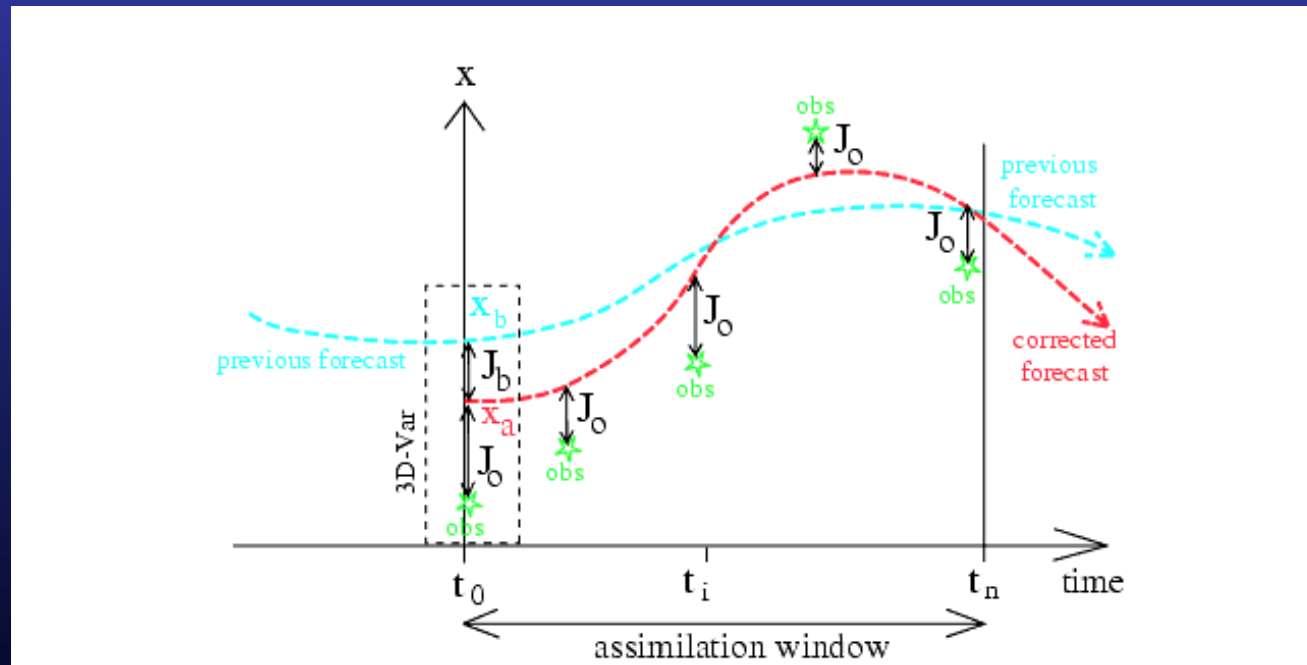
- Bayesian approach to obtain best estimate of state of the atmosphere given **forecast** and a **set of new observations**.





## Data Assimilation

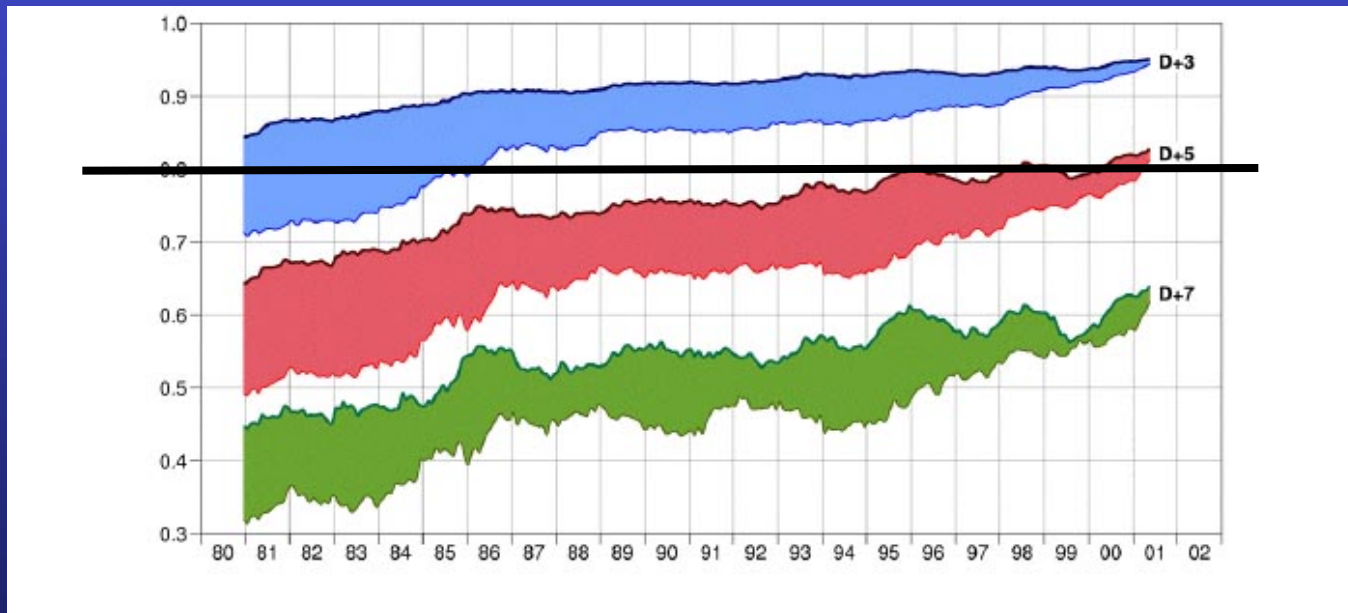
- Bayesian approach to obtain best estimate of state of the atmosphere given **forecast** and a **set of new observations**.



## Numerical Weather Prediction Models

- **Forecast at time t:** Best estimate of the state of the atmosphere at some future point in time t.
  - **Analysis at time t:** Best estimate of the state of the atmosphere point in time t based on **forecast + sparse observations**.
  - Both, forecast and analysis provide estimates of T,p,q everywhere on the globe.
-

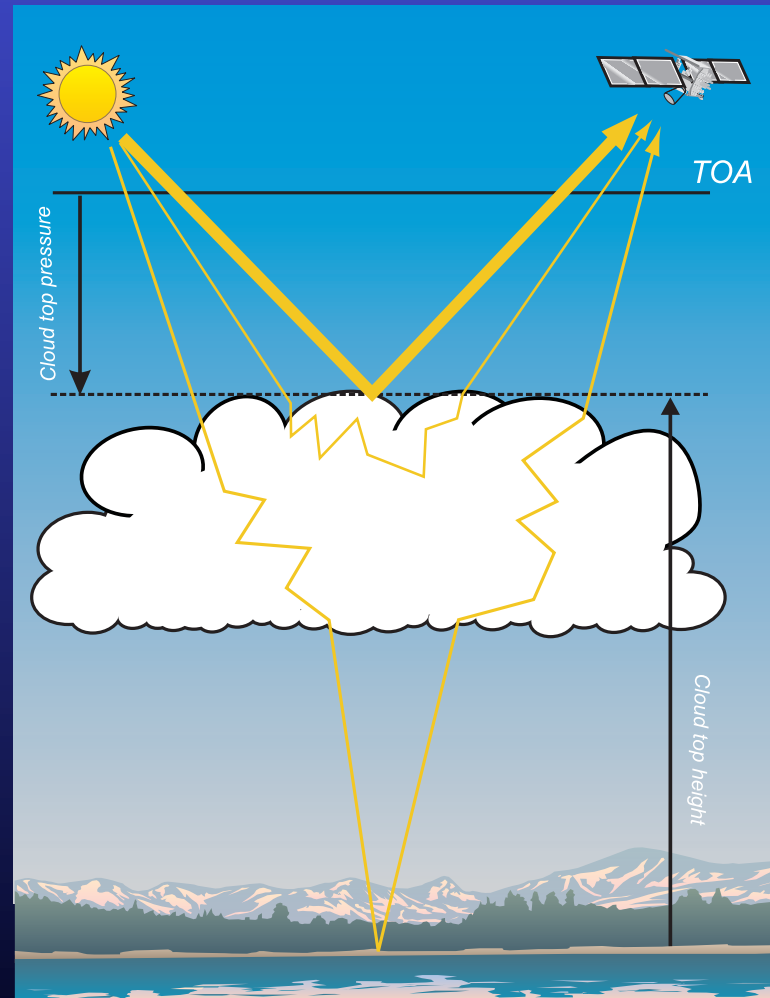
## NWP Forecast Skill



- 0: No skill, 1: perfect skill (based on anomaly correlations)
  - 5 day forecast 2001 better than 3 day forecast 1980
  - Improvements due to satellite (better initial conditions) and better computer technology.
-

## Short-wave (300-3000 nm) radiative transfer

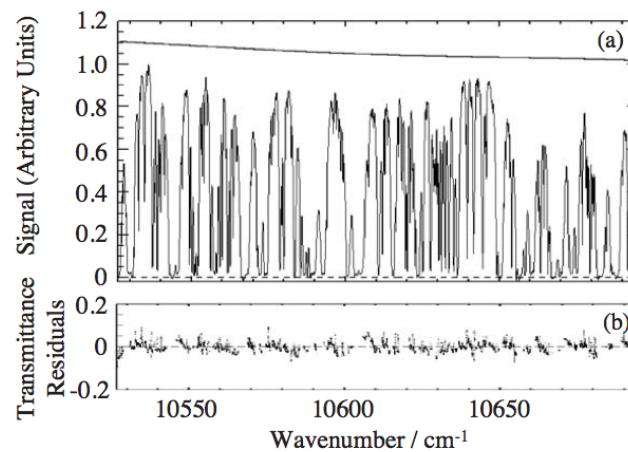
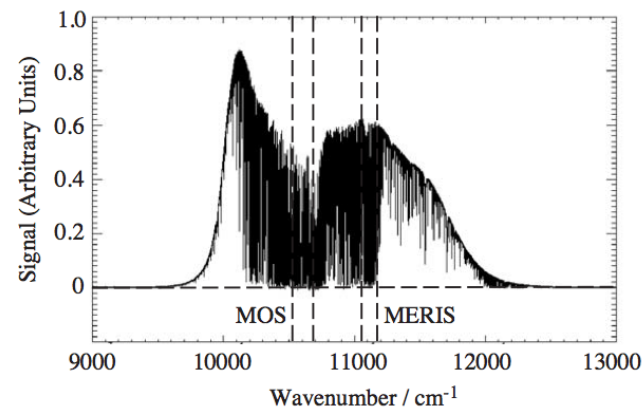
- Scattering become important, RT gets way more complicated than infrared
- How to determine aerosol scattering properties  
→ **Eloranta talk**
- Formal radiative transfer solutions and absorption/scattering interaction



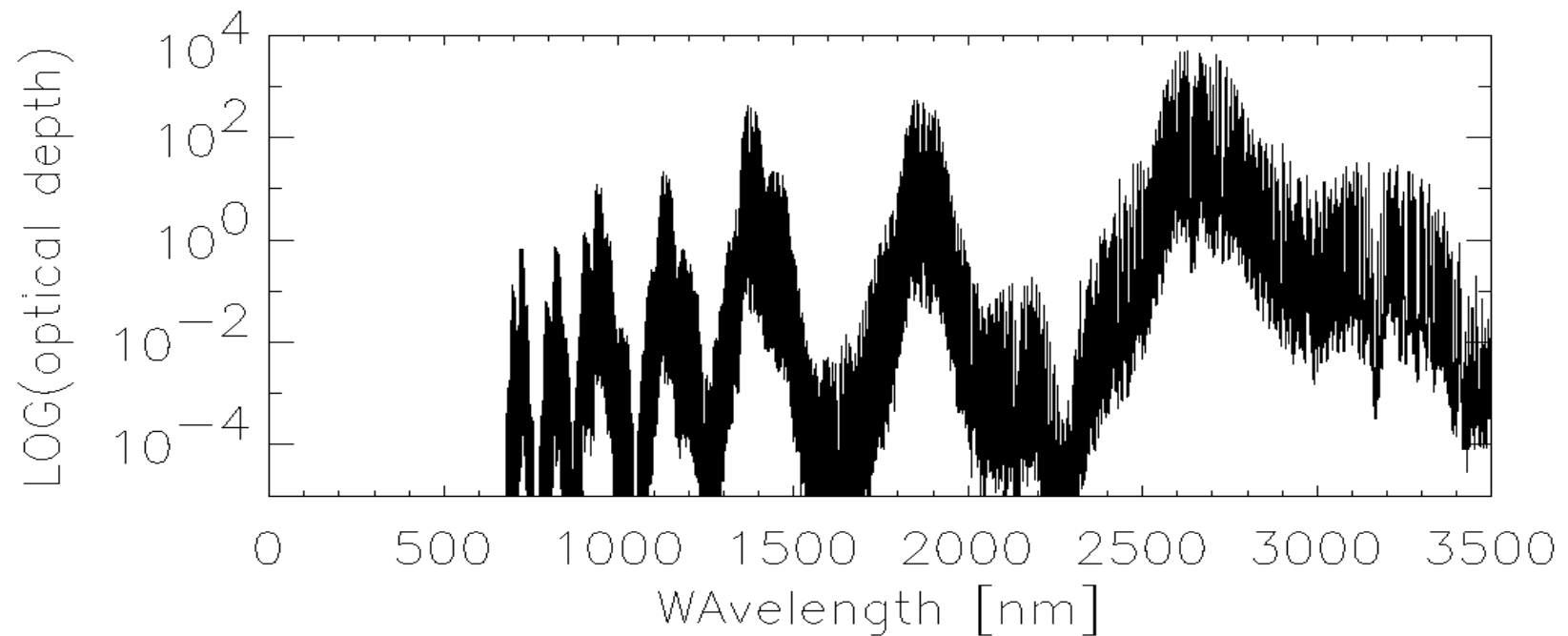
## Example: Solar radiative transfer example water vapor spectroscopy in the water vapor rst-band (945 nm)

*P. Albert et al. / Journal of Quantitative Spectroscopy & Radiative Transfer 84 (2004) 181–193*

187



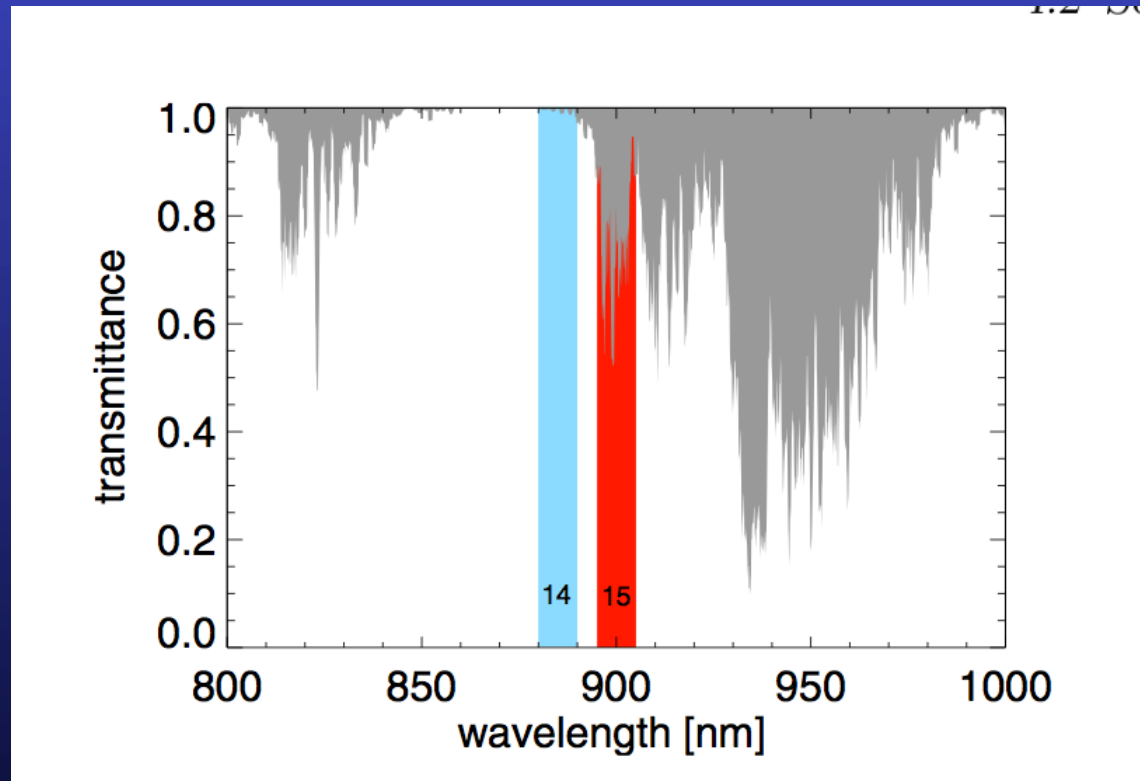
## Broad band water vapor absorption in the NIR



Mid-latitude summer atmosphere

---

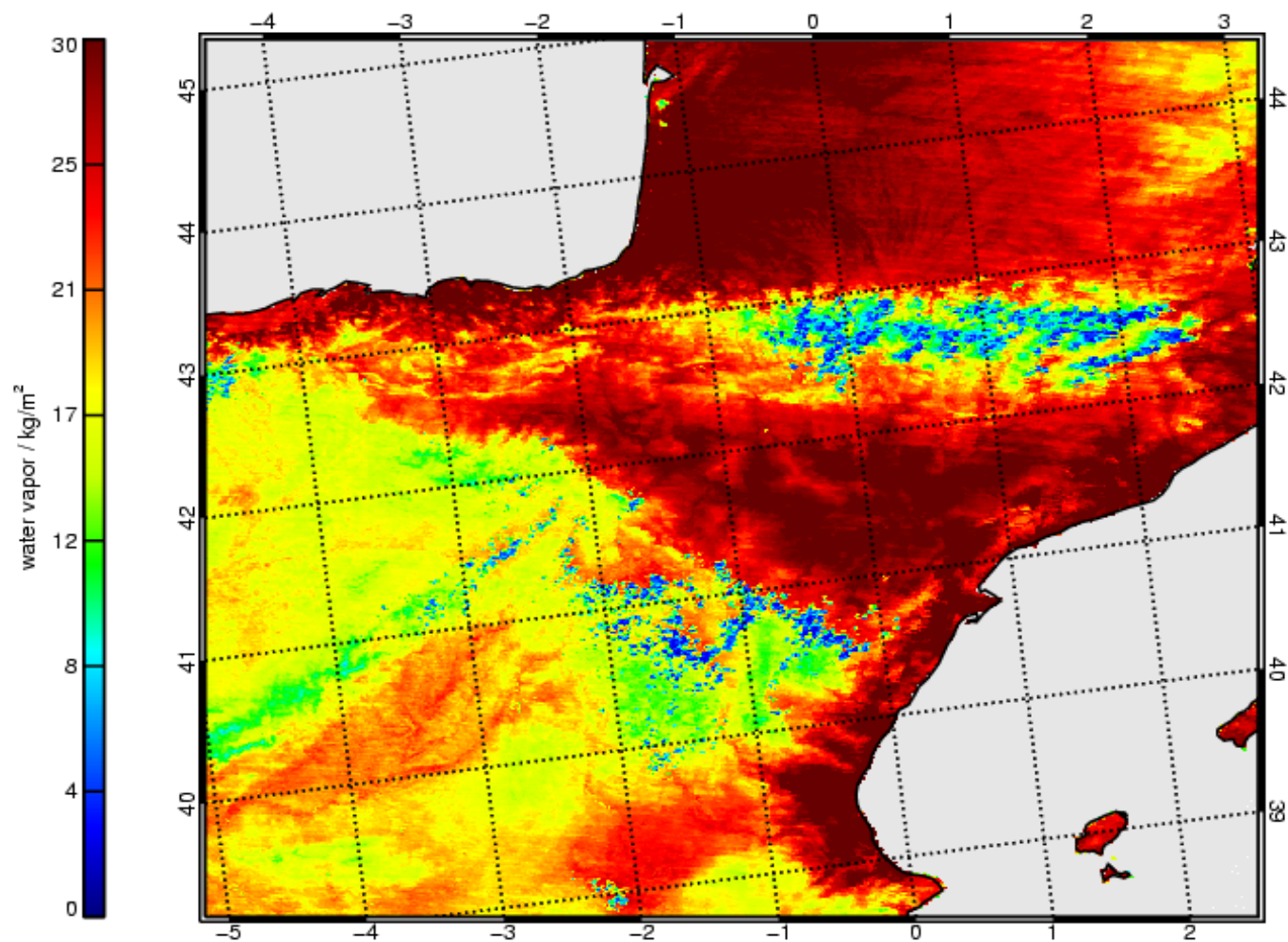
## Differential absorption water vapor spectroscopy



Operationally used on satellites/aircrafts since 2000. (E.g. Bennartz and Fischer 2001, Rem. Sens. Env.)

---

## Near infrared MODIS water vapor – example





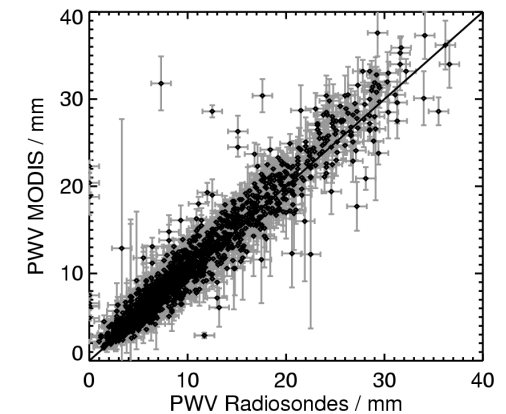
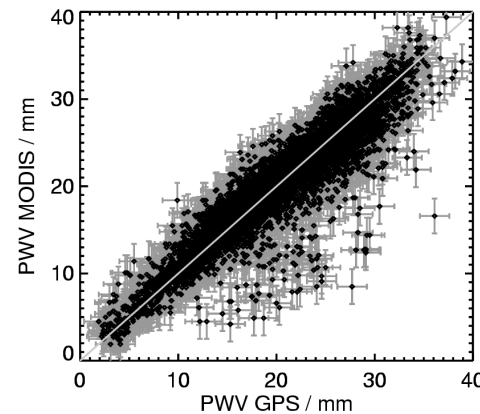
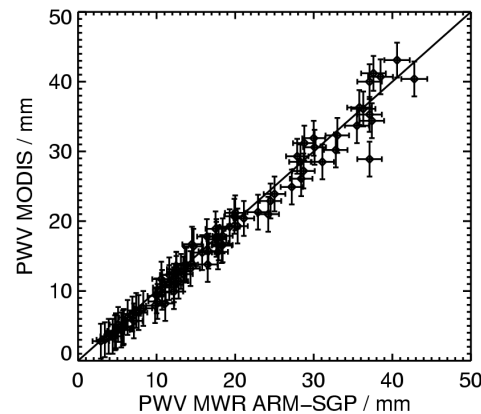
# Validation of Near infrared water vapor estimate over land (Accuracy: 1.5-2 kg/m<sup>2</sup> (about 10%))

Microwave

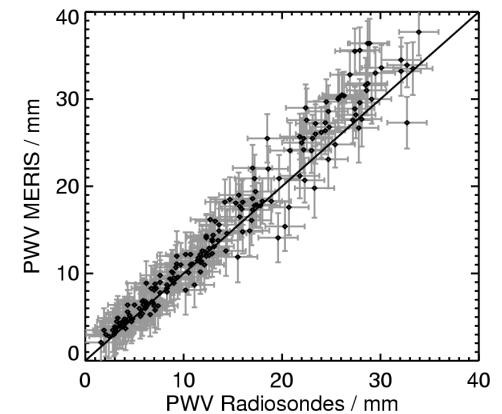
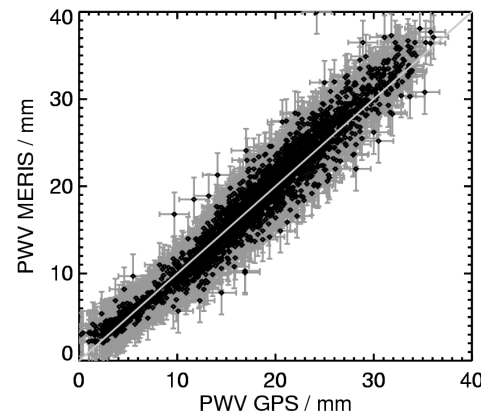
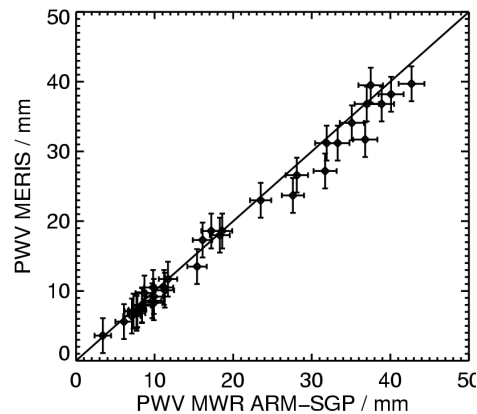
GPS

Radiosonde

MODIS



MERIS



(Bennartz & Fischer, 2001; Albert et al., 2005)

## Summary

- Infrared and microwave spectrometers provide excellent opportunity to observe temperature and humidity.
  - NWP models (forecasts/analyses) provide good estimates but are less reliable in data sparse regions (southern versus northern hemisphere).
  - Ground-based observations of temperature and water vapor → microwave radiometer (stable, accurate, cheap (30K), runs automatically).
  - Various radiative transfer tools, models etc.. (Monte-Carlo, adding-and-doubling, photon pathlength...)
-