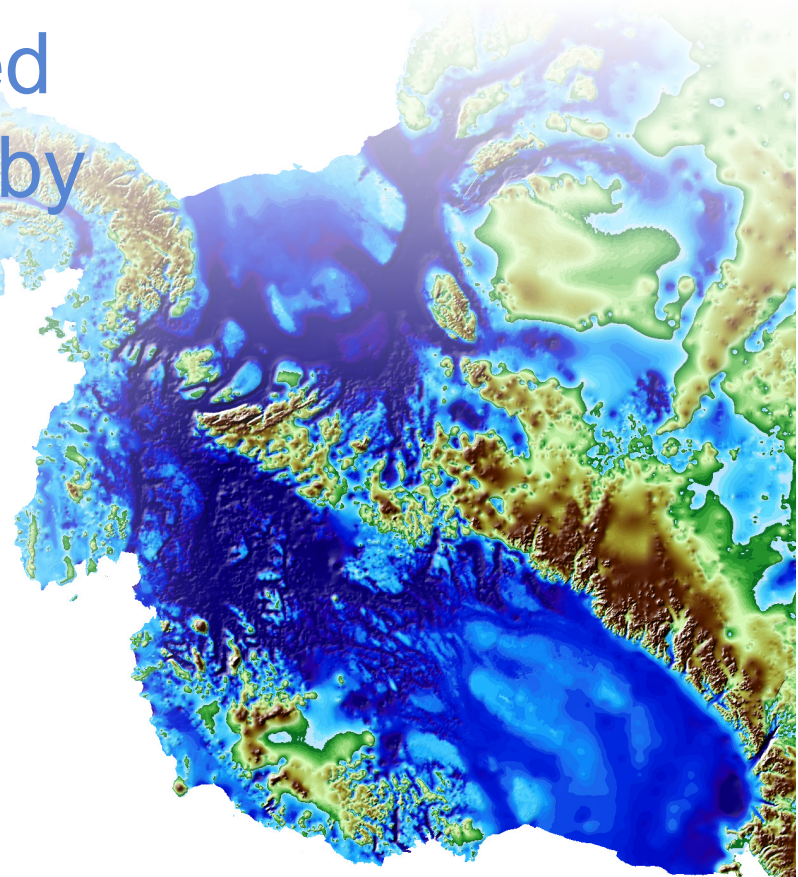




BedMachine: mapping the bed under the Antarctic ice sheet by combining sparse radar data and mass conservation

Mathieu MORLIGHEM

University of California Irvine,
Department of Earth System Science



Conservation of mass

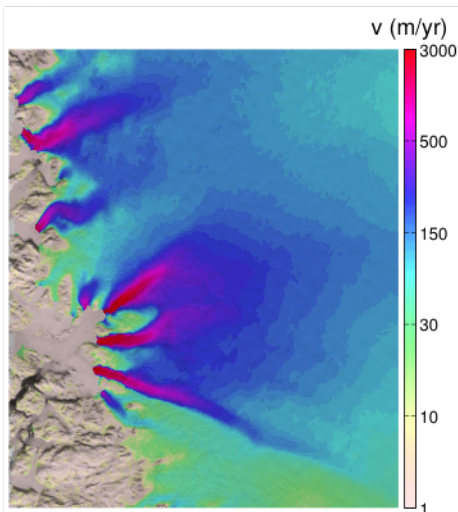
- Compressibility coefficient: $\chi_T = 1.2 \times 10^{-10} \text{ Pa}^{-1}$
 - Linear thermal expansion of ice: $5 \times 10^{-5} \text{ }^\circ\text{C}^{-1}$
- Density varies by less than 0.5% in Antarctica and Greenland
- Ice can be considered an incompressible fluid

$$\frac{D\rho}{Dt} + \rho \nabla \cdot \mathbf{v} = 0 \quad \rightarrow \quad \nabla \cdot \mathbf{v} = 0$$

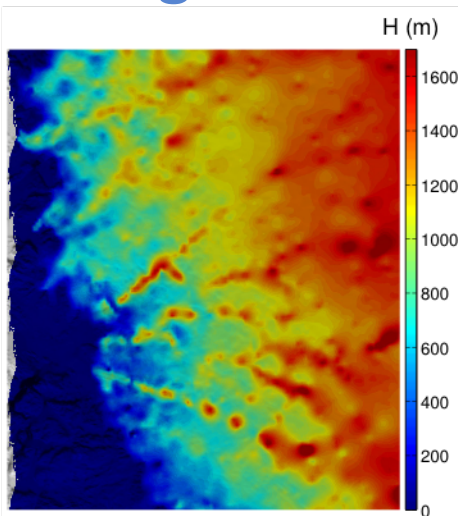
- Depth integrating between the bottom and the top surface of the ice:

$$\frac{\partial H}{\partial t} = - \underbrace{\nabla \cdot H \bar{\mathbf{v}}}_{\text{Flux divergence}} + \underbrace{\dot{M}_s}_{\text{SMB}} + \underbrace{\dot{M}_b}_{\text{Basal melt}}$$

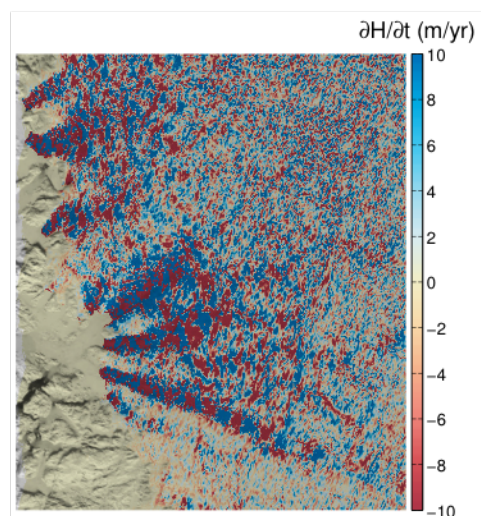
How about grounded ice?



(a) InSAR ice velocity
Rignot and Mouginot 2012



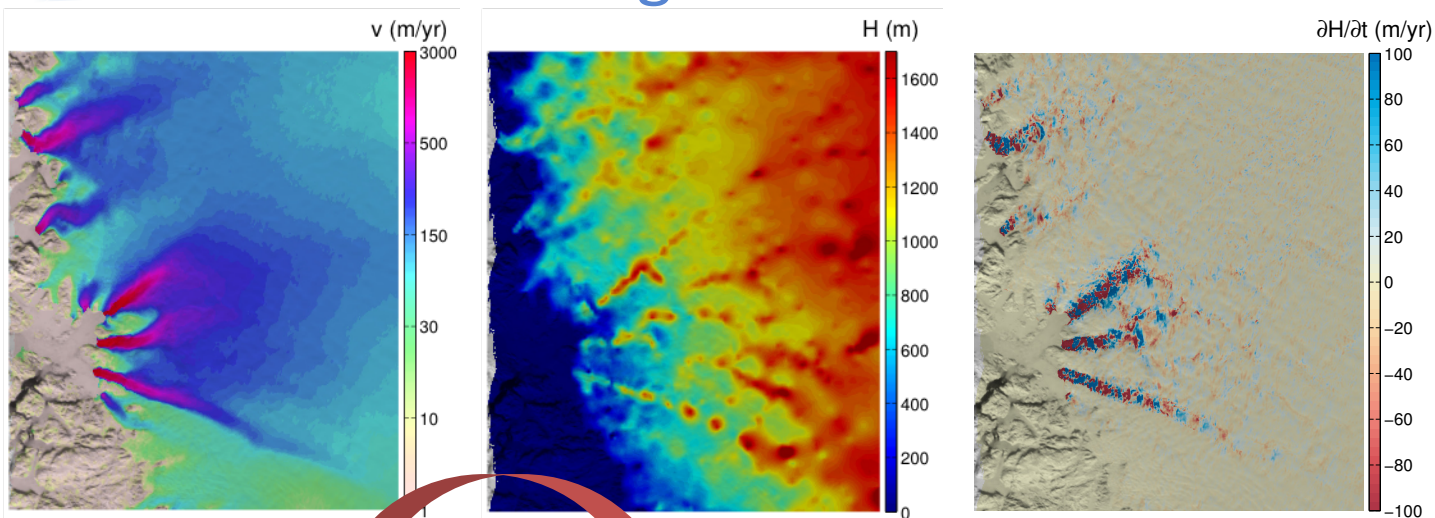
(b) Ice Thickness
From Kriging



(c) Calculated $\partial H/\partial t$

$$\frac{\partial H}{\partial t} = - \underbrace{\nabla \cdot H \bar{\mathbf{v}}}_{\text{Flux divergence}} + \underbrace{\dot{M}_s}_{\text{SMB}} + \underbrace{\dot{M}_b}_{\text{Basal melt}}$$

How about grounded ice?



(a) InSAR ice velocity
Rignot and Mouginot (2012)

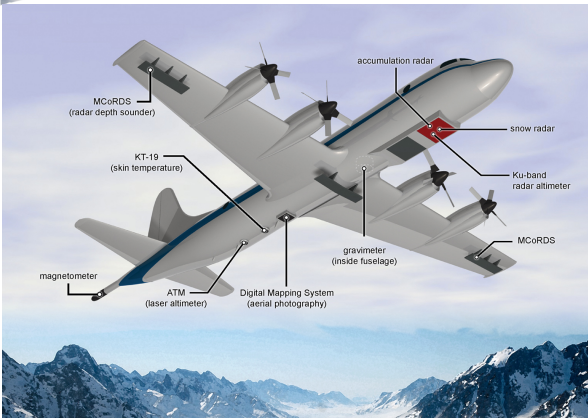
(b) Ice Thickness
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$$\frac{\partial H}{\partial t} = - \underbrace{\nabla \cdot H \bar{v}}_{\text{Flux divergence}} + \underbrace{\dot{M}_s}_{\text{SMB}} + \underbrace{\dot{M}_b}_{\text{Basal melt}}$$

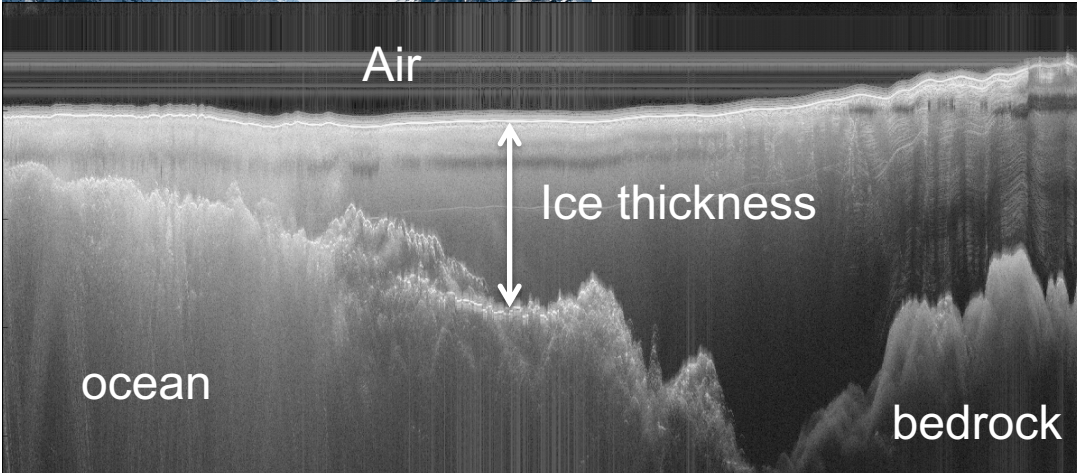
- Initialization of ice sheet model difficult
- Strong need for “physics-derived” bedrock elevation maps

Ice penetrating radar

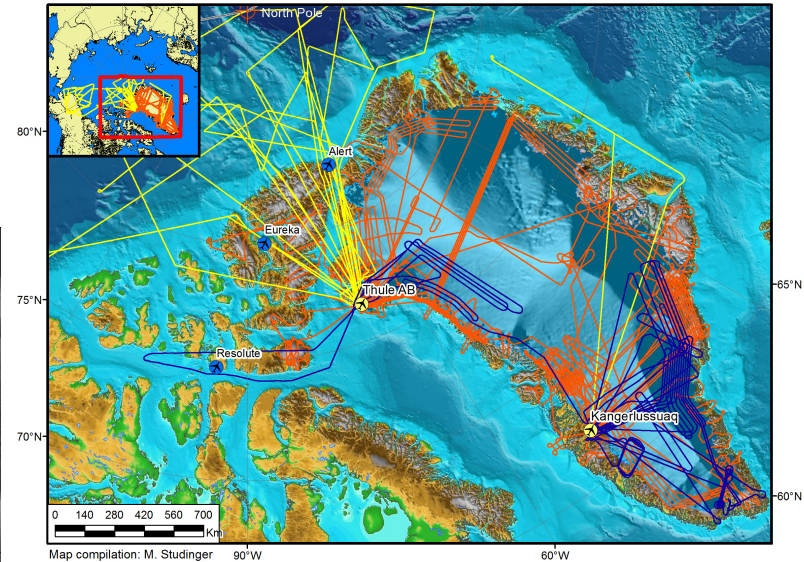


NASA P-3

Radar echogram



Flight lines OIB Arctic 2012

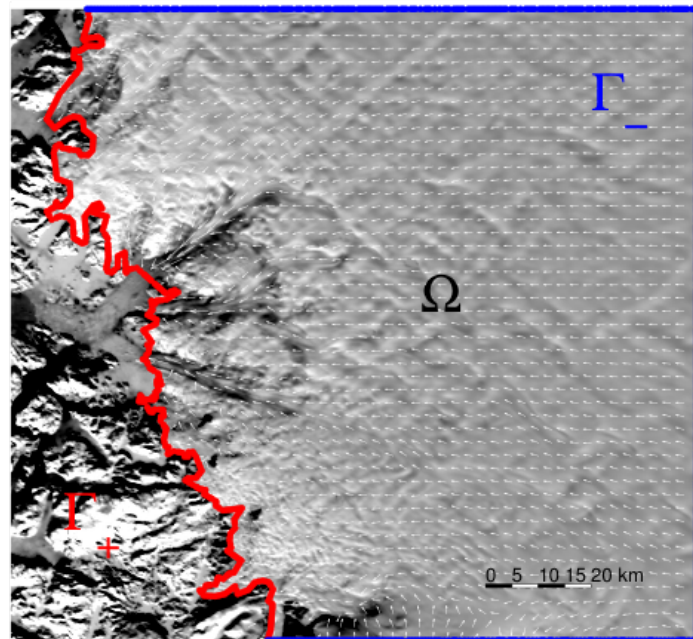


Mass Conservation Approach

Find H such that:

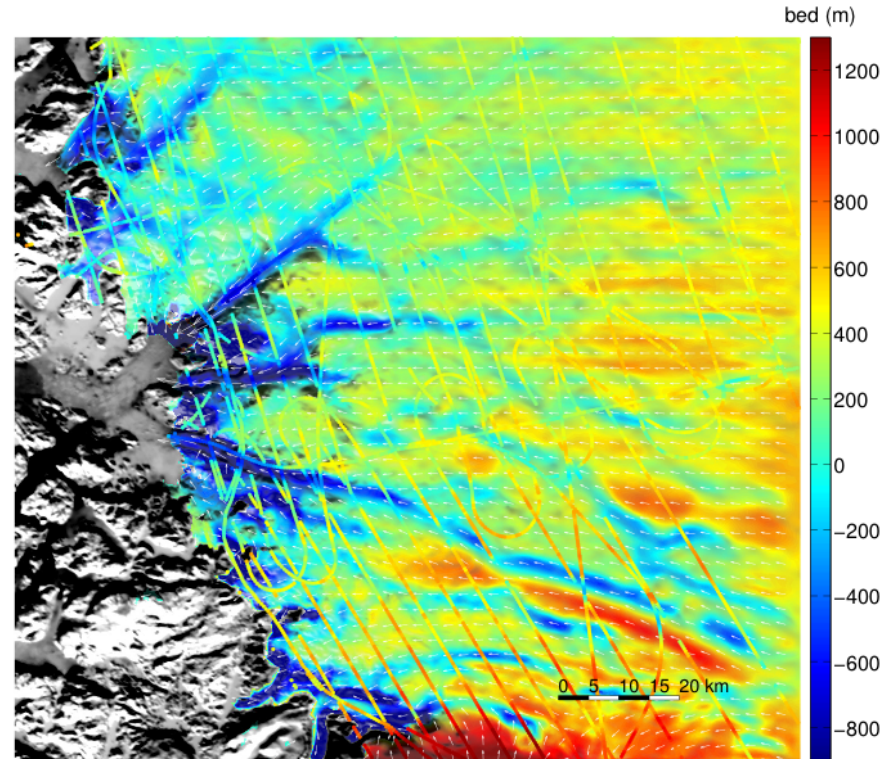
$$\left\{ \begin{array}{l} \nabla \cdot H \bar{\mathbf{v}} = \underbrace{\dot{M}_s + \dot{M}_b}_{\dot{a}} - \frac{\partial H}{\partial t} \quad \text{in } \Omega \\ H = H_{obs} \quad \text{on } \Gamma_- \end{array} \right.$$

- $\partial H / \partial t$ from ICESat/ATM
- SMB from regional climate model
- H_{obs} from OIB flight line
- Velocity \approx InSAR surface velocities or modeled



Mass Conservation Approach

- Thickness only constrained on the inflow boundary
- We want the *modeled* thickness to match the measurements



PDE-Constrained optimization

Minimize the cost function:

$$\mathcal{J}(\bar{\mathbf{v}}, \dot{a}) = \int_{\text{Tracks}} \frac{1}{2} (H - H_{obs})^2 dl$$

With the constraint:

$$\begin{cases} \nabla \cdot H \bar{\mathbf{v}} & = & \dot{a} & \text{in } \Omega \\ H & = & H_{obs} & \text{on } \Gamma_- \end{cases}$$

Controls:

$$\bar{\mathbf{v}} \in [0.95 (\mathbf{v}_{obs} - 50) \quad \mathbf{v}_{obs} + 50] \text{ m/yr}$$

$$\dot{a} = \dot{a}_{obs} \pm 1 \text{ m/yr}$$

Variational data assimilation

Lagrangian:

$$\mathcal{L}(H, \lambda, \bar{\mathbf{v}}, \dot{a}) = J(H) + \int_{\Omega} \lambda (\nabla \cdot H \bar{\mathbf{v}} - \dot{a}) d\Omega$$

The adjoint state is defined by:

$$\forall \varphi \in V_H \quad \int_T \varphi (H - H_{obs}) dx - \int_{\Omega} \varphi \bar{\mathbf{v}} \cdot \nabla \lambda d\Omega + \int_{\Gamma_+} \varphi \lambda \bar{\mathbf{v}} \cdot \mathbf{n} dS = 0$$

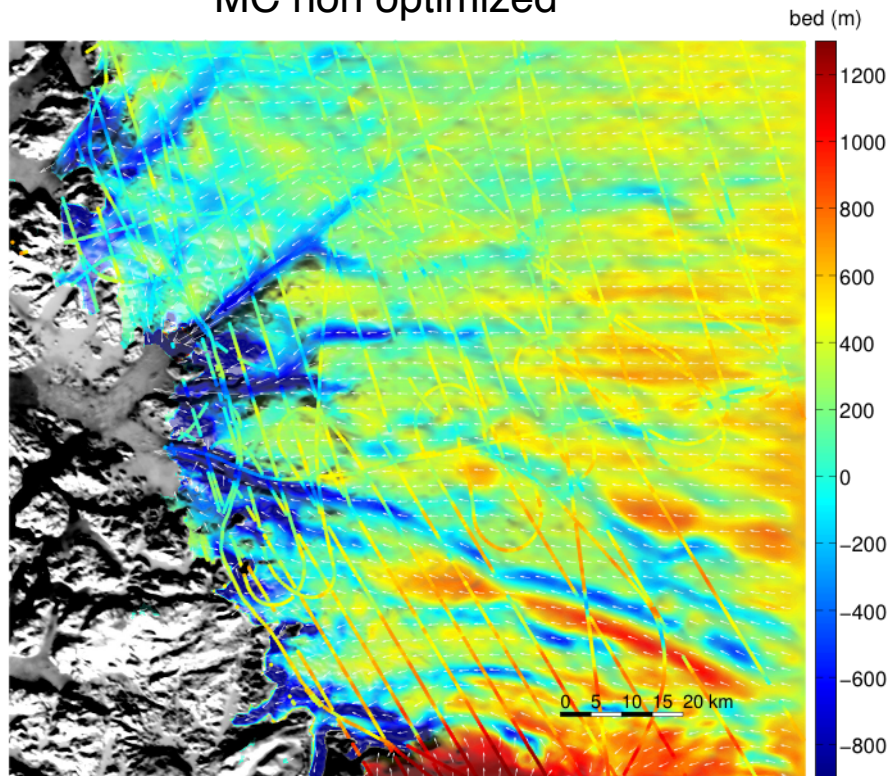
Gradients:

$$\mathcal{J}'(\bar{\mathbf{v}}) \simeq -H \nabla \lambda$$

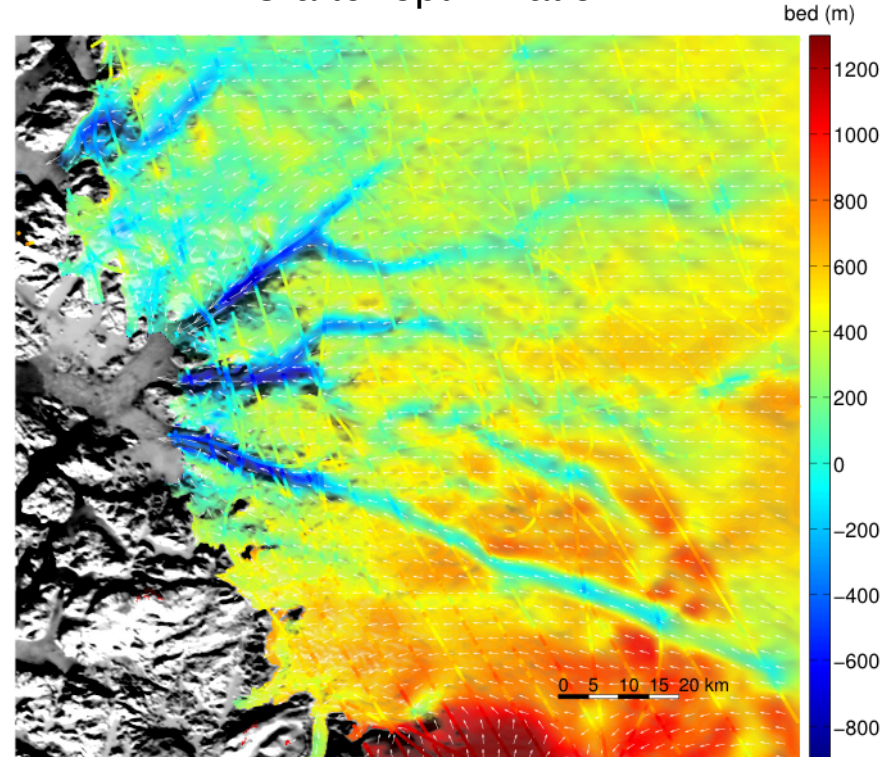
$$\mathcal{J}'(\dot{a}) \simeq \lambda$$

PDE-Constrained optimization

MC non optimized



MC after optimization



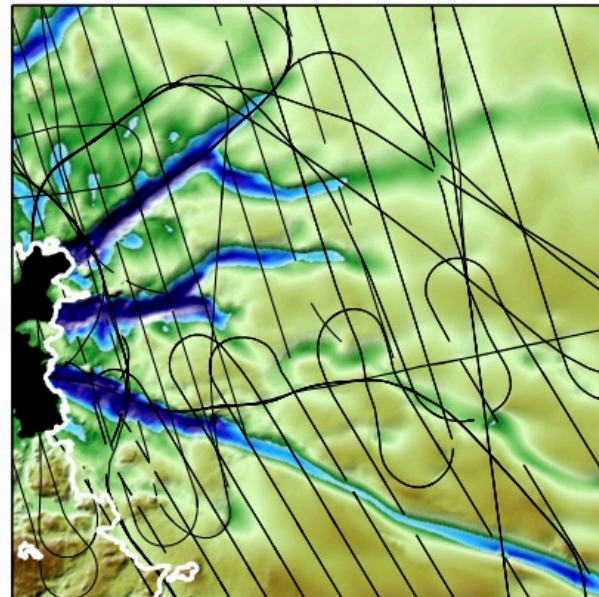
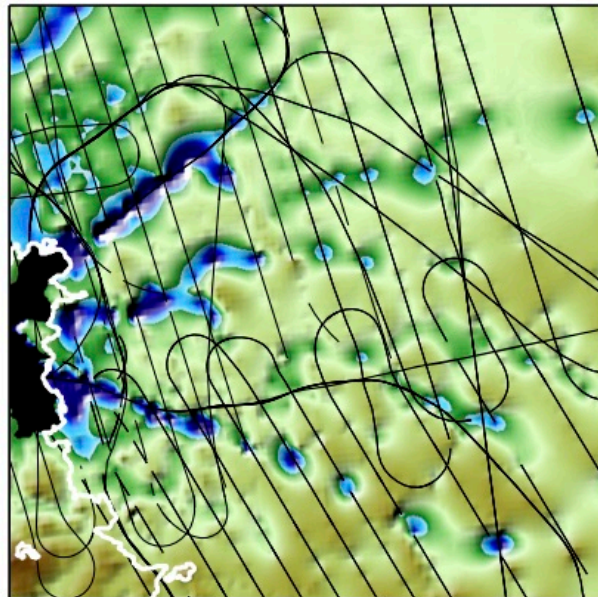
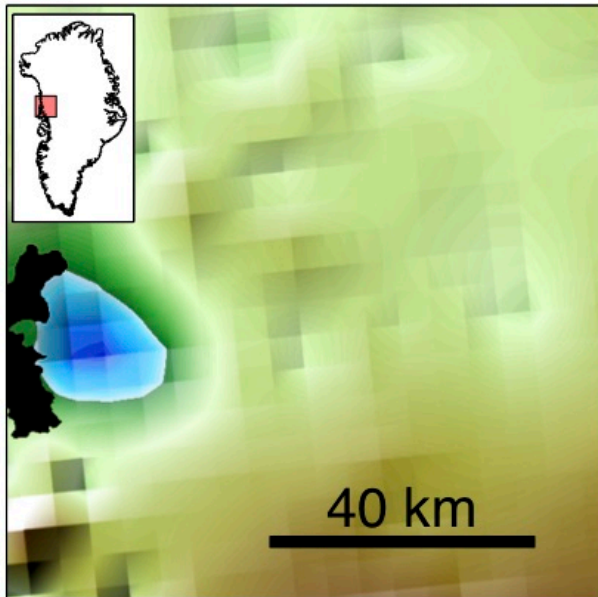
Comparison with other datasets

Bamber 2001

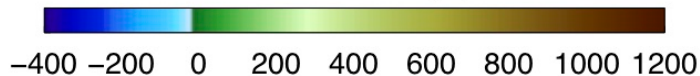
Bamber 2013

Mass Conservation

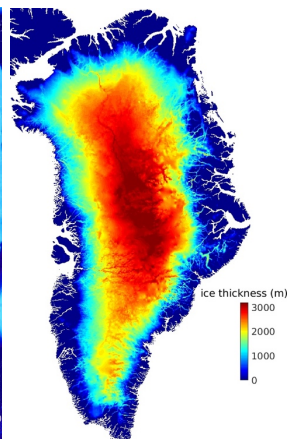
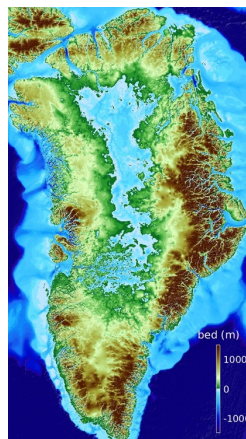
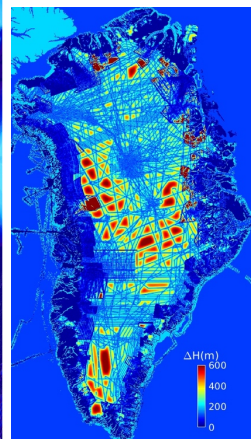
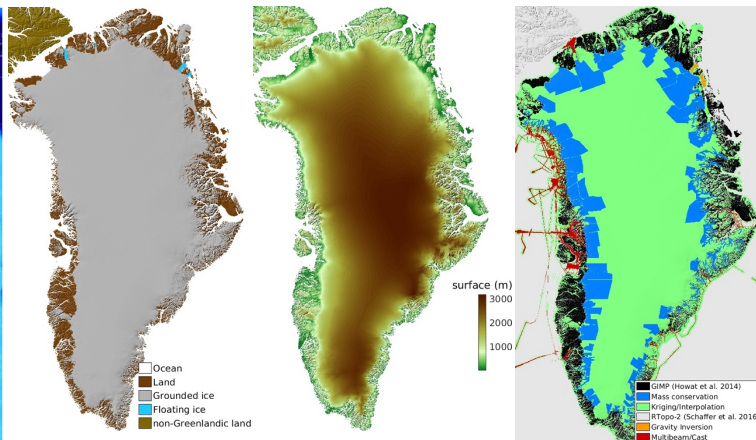
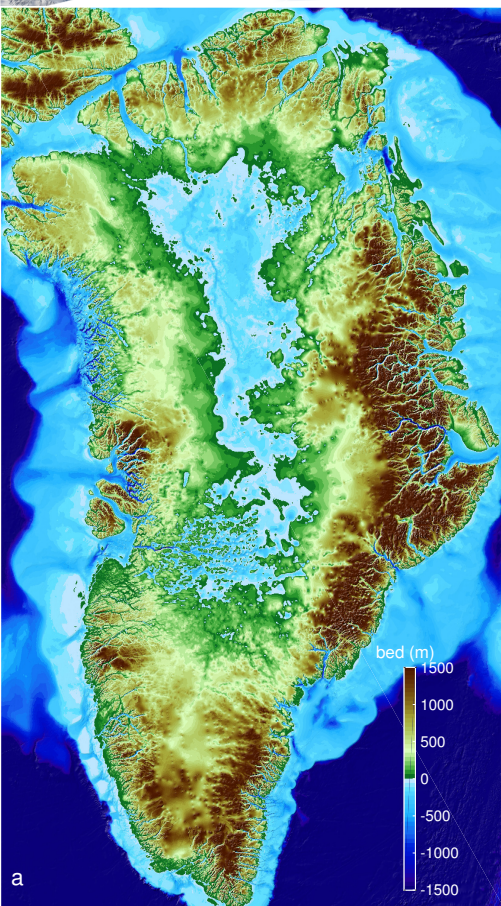
Upernavik Isstrøm



Bed elevation (m)

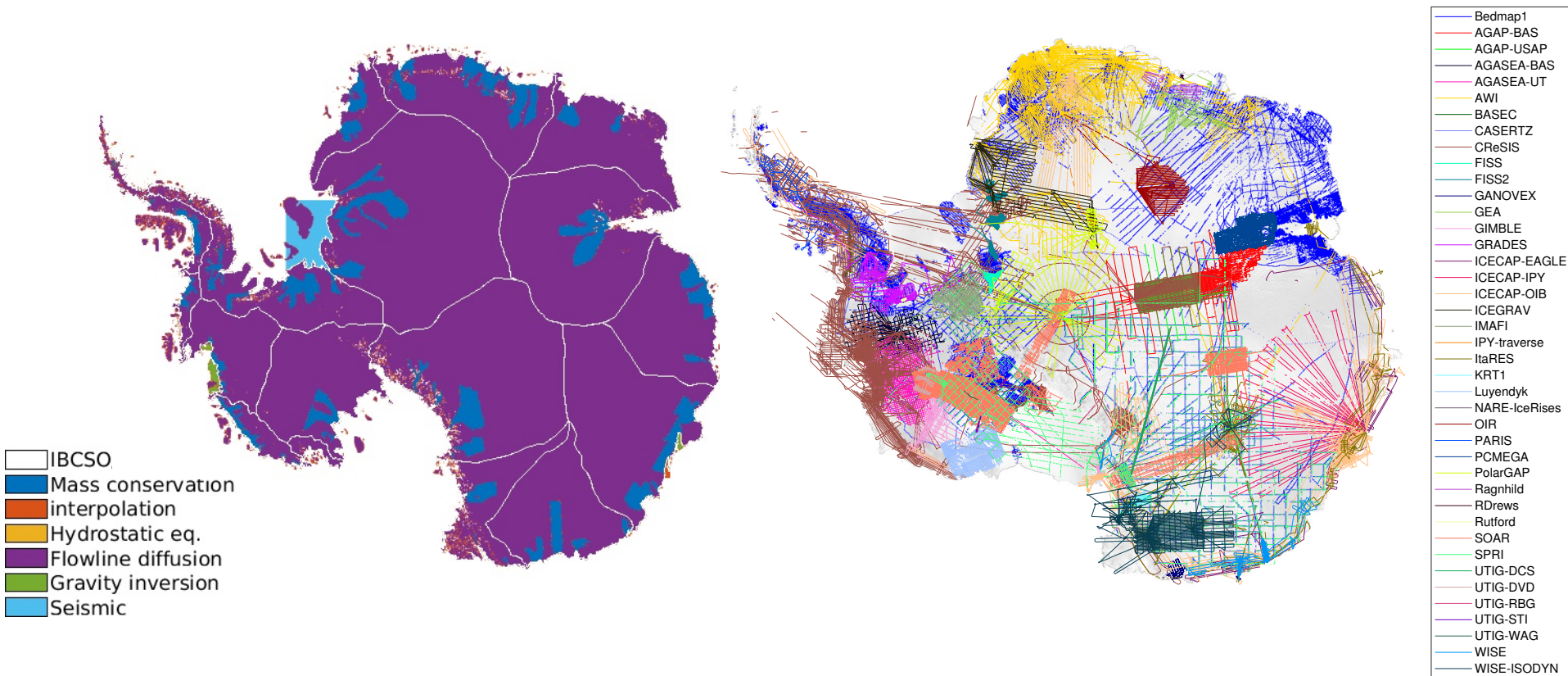


BedMachine Greenland v3



- Self-consistent dataset of the topography of Greenland
 - mask
 - surface elevation
 - ice thickness
 - bed elevation
 - error map
 - source map
 - Geoid height (WGS84)
- 150-m grids (true resolution depends on the region)
- V1: Posted at NSIDC since 2014
- V3: Available since Sept. 2017
- V4: under development...

BedMachine Antarctica





Credit: P Fretwell (BAS)

Flowline diffusion

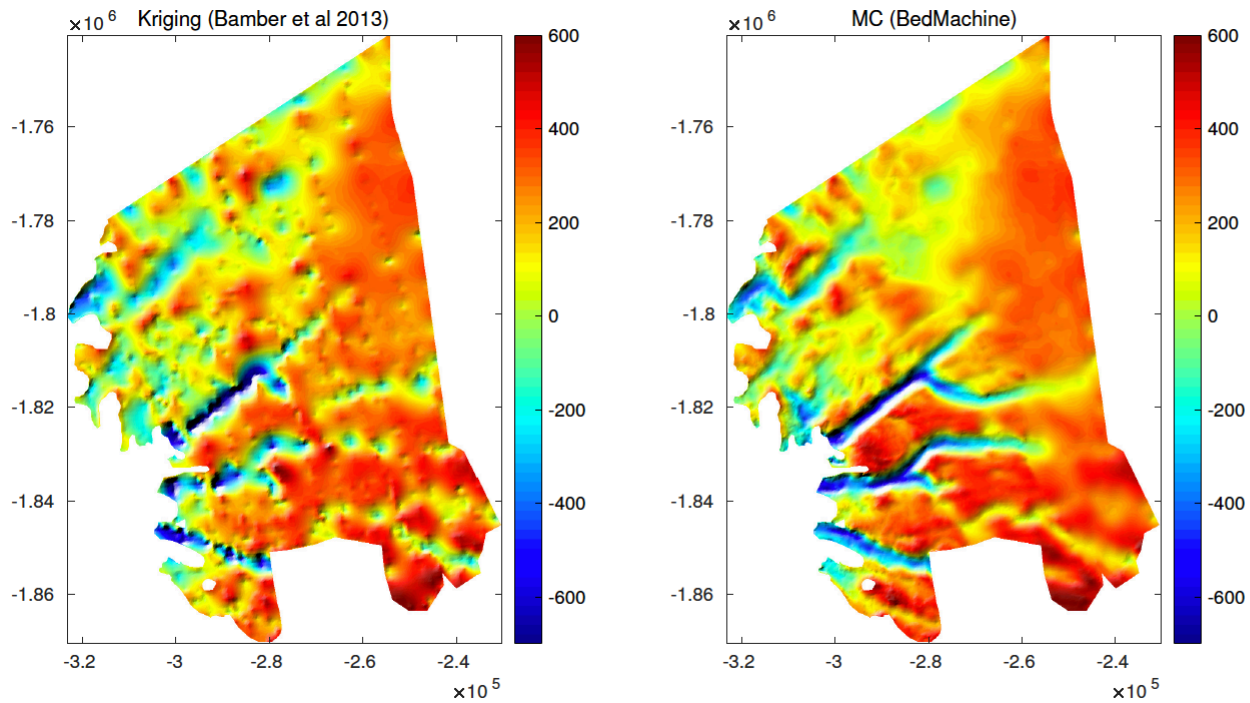
Find the ice thickness, H , such that

$$\begin{cases} \nabla \cdot \mathbf{D} \nabla H = 0 & \text{in } \Omega \\ H = H_{obs} & \text{on } T \end{cases}$$

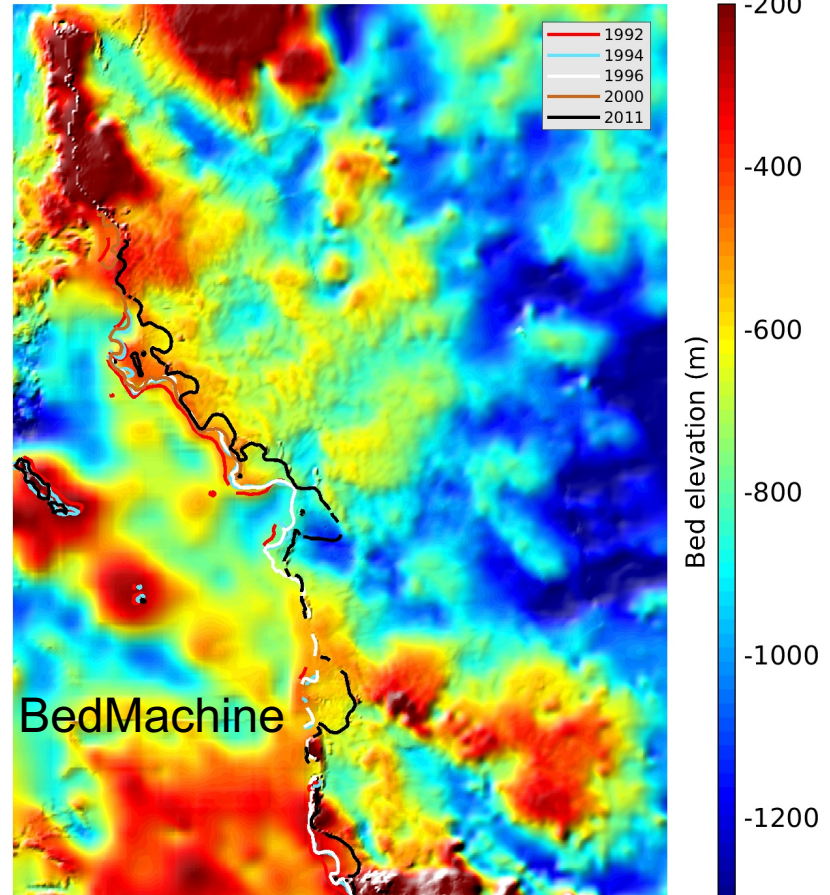
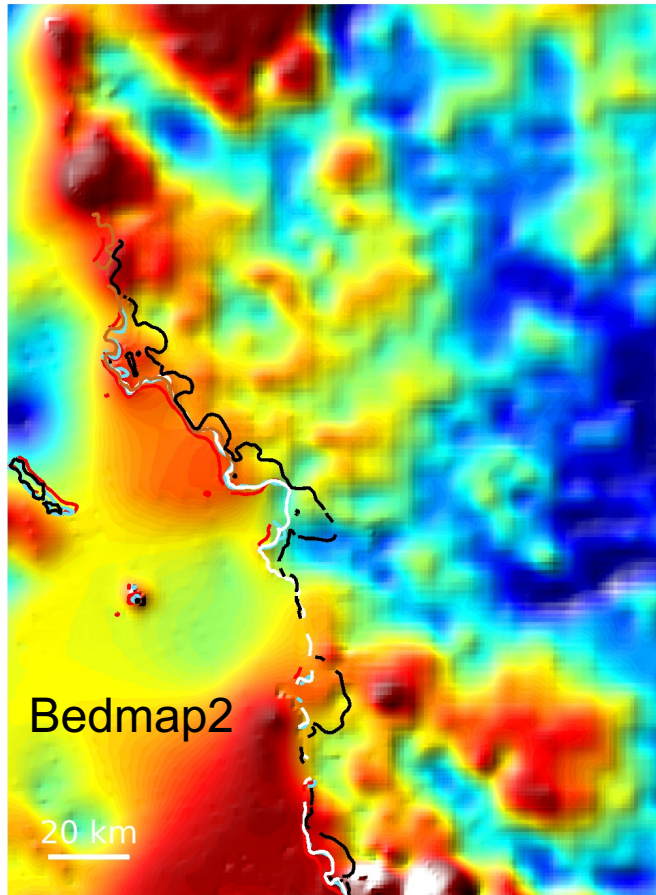
Anisotropic diffusion tensor:

$$\mathbf{D} = \mathbf{v}_s \otimes \mathbf{v}_s + \tau \mathbf{I}$$

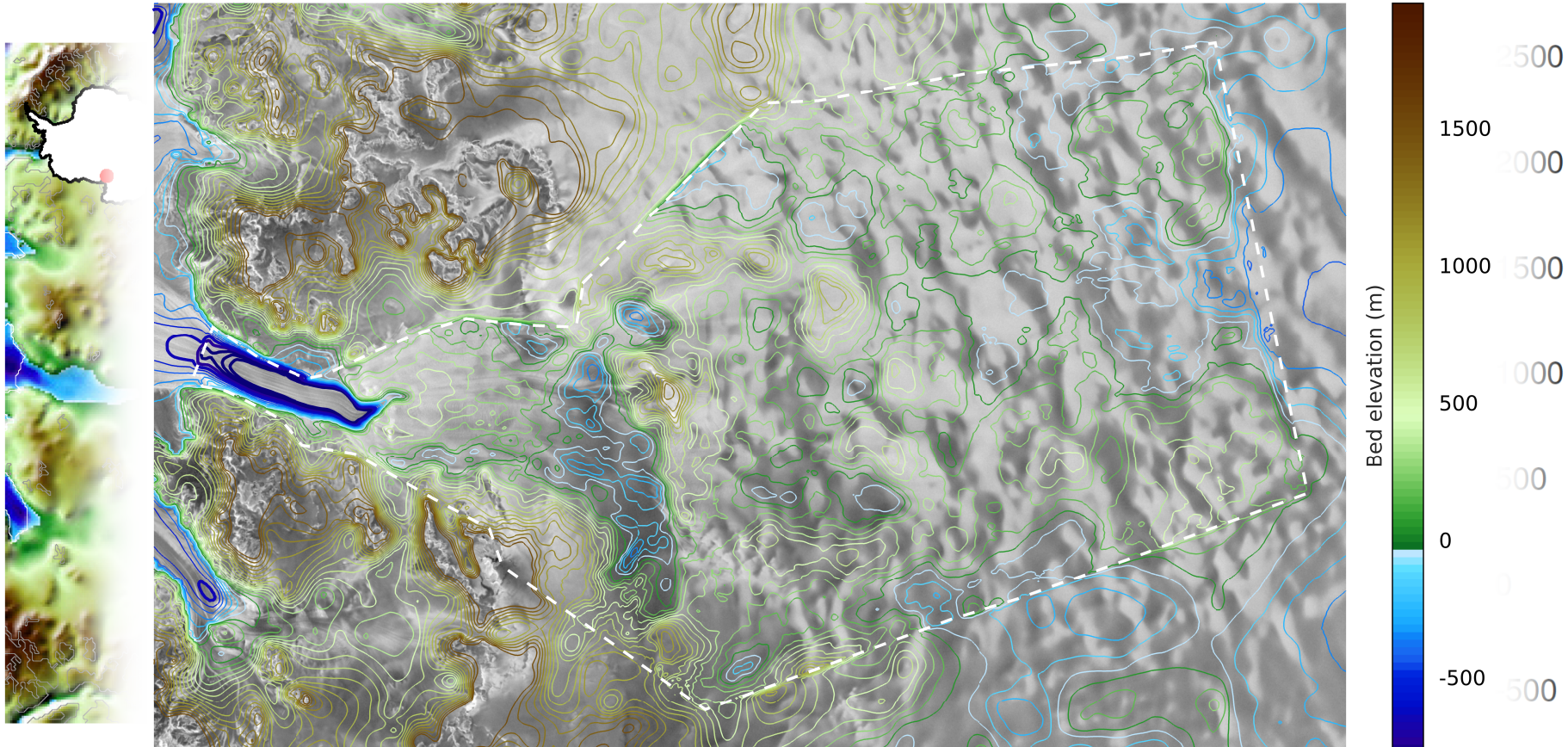
Flowline diffusion



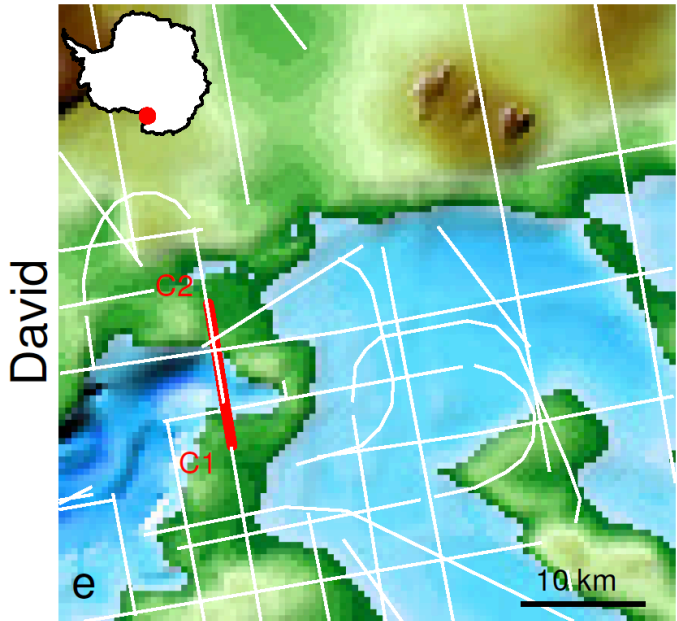
Thwaites Glacier



Mulock Glacier

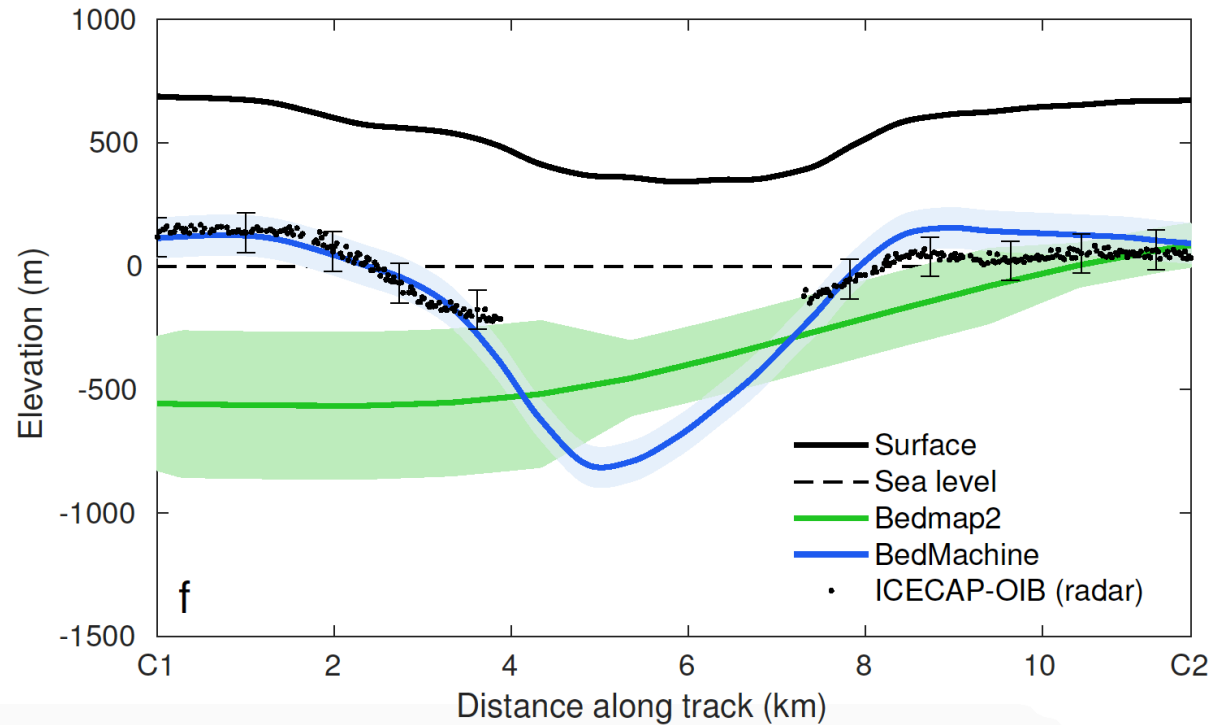


David Glacier

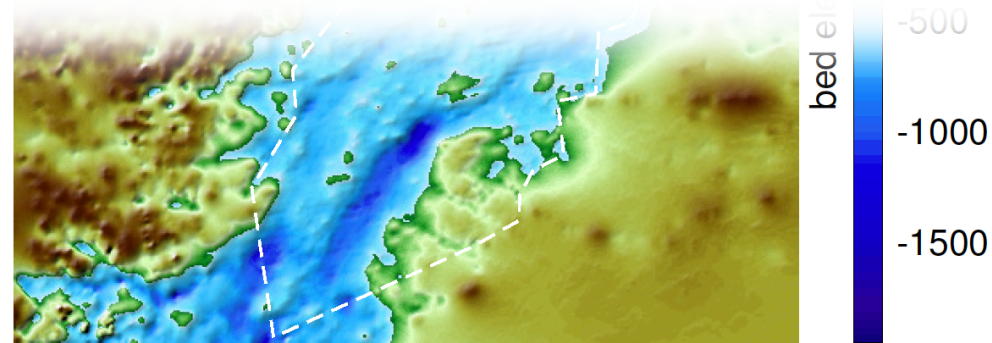
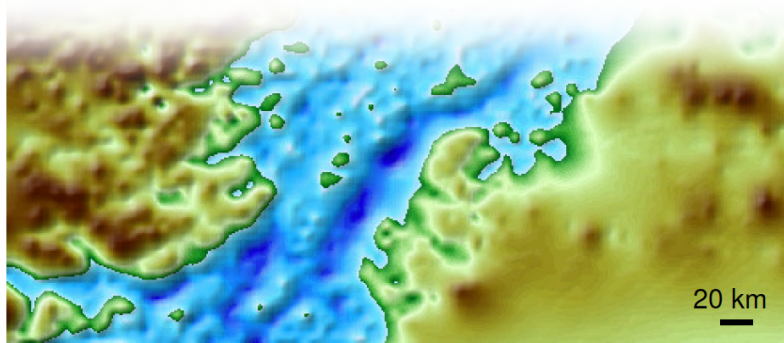
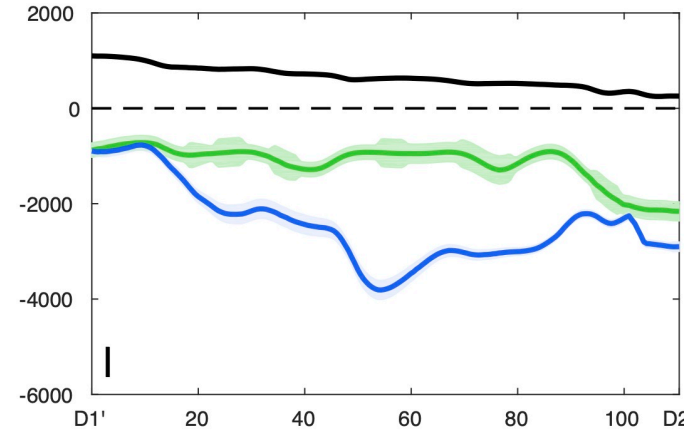
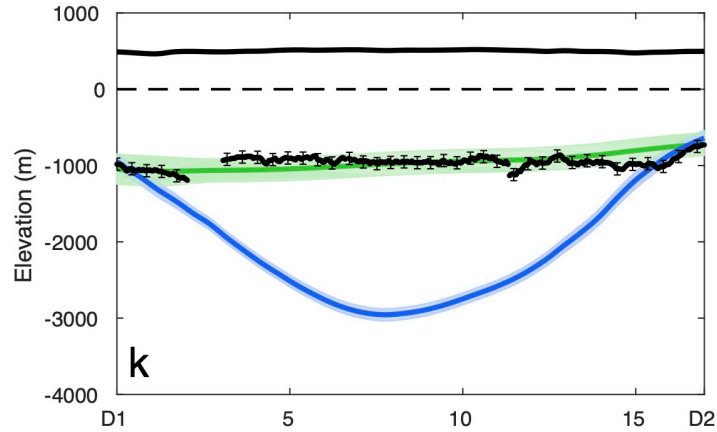
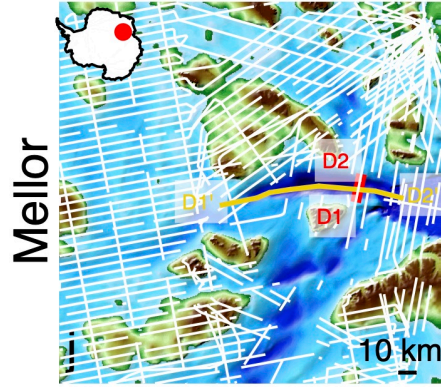


bed elevation (m)

-3000 -2000 -1000 0 1000

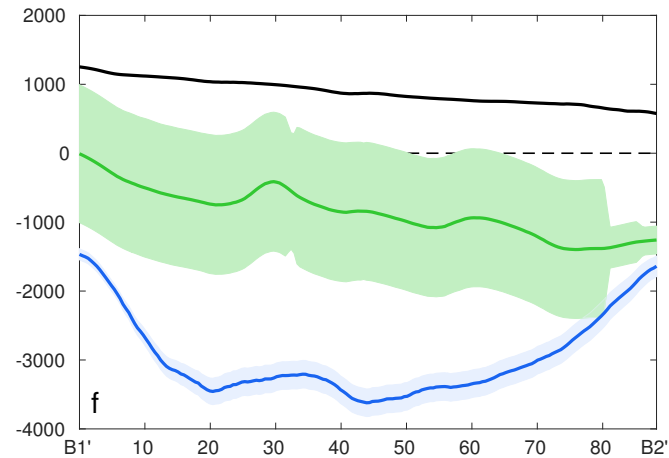
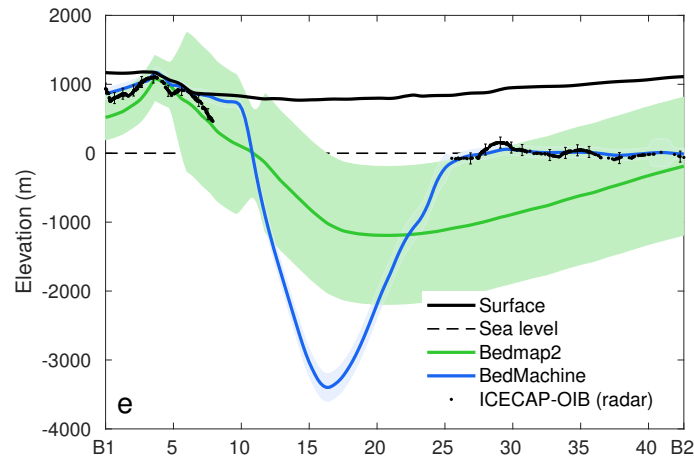
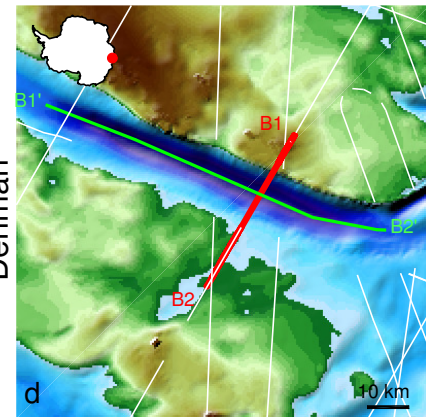


Mellor/Amery

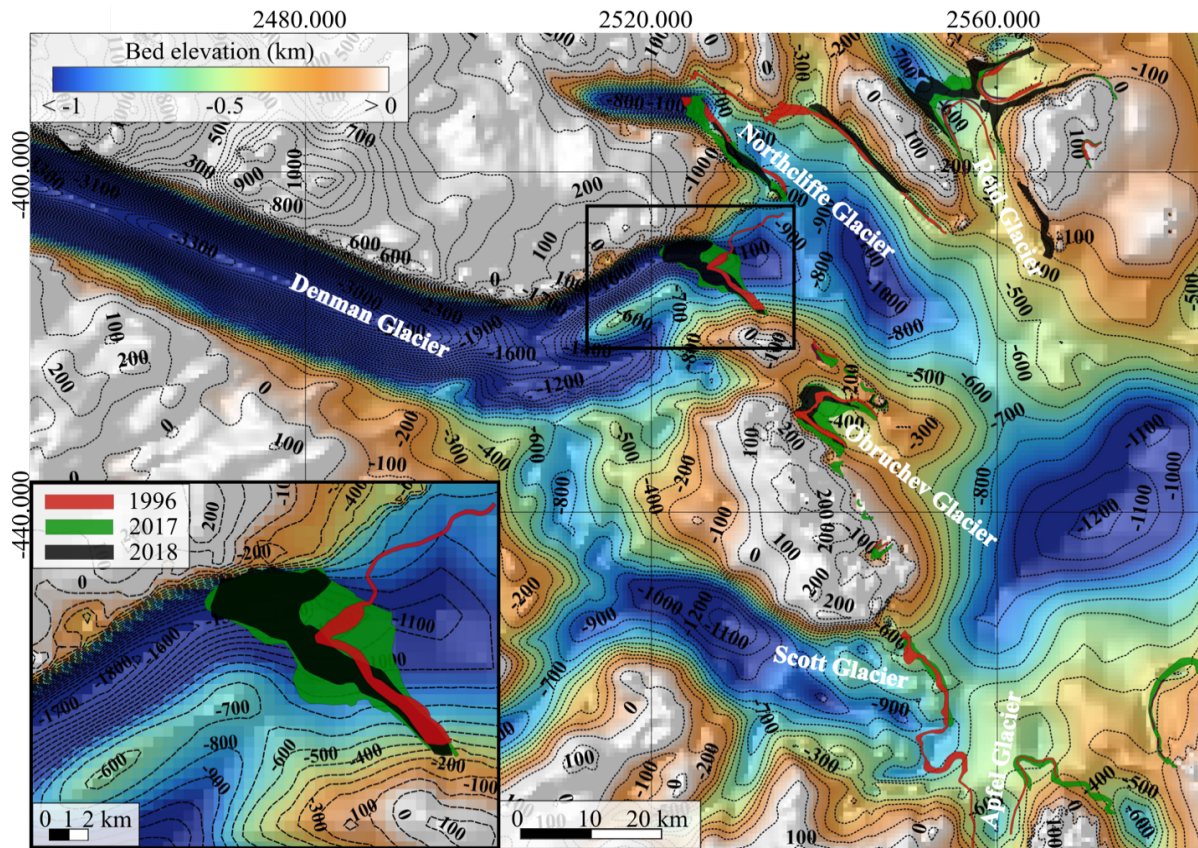


Denman

Bedmap2

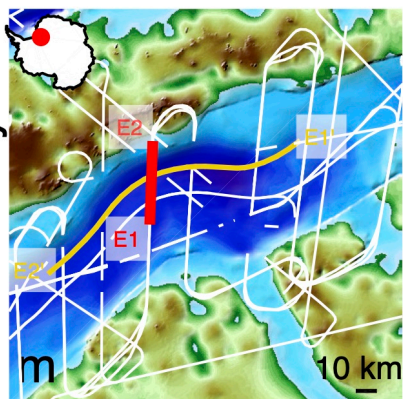


Denman

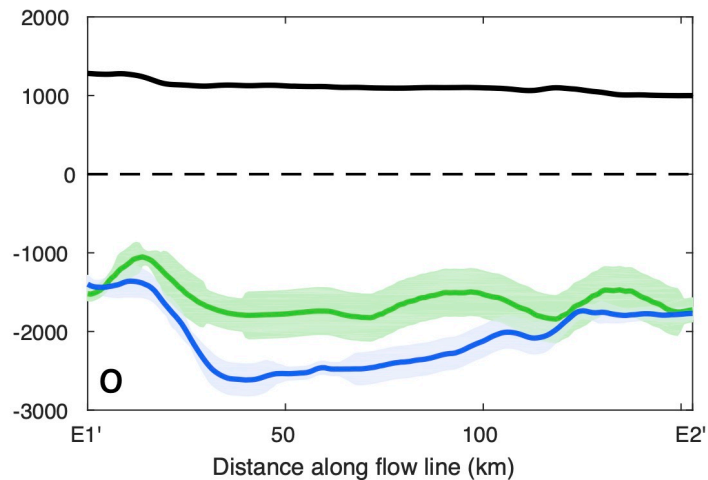
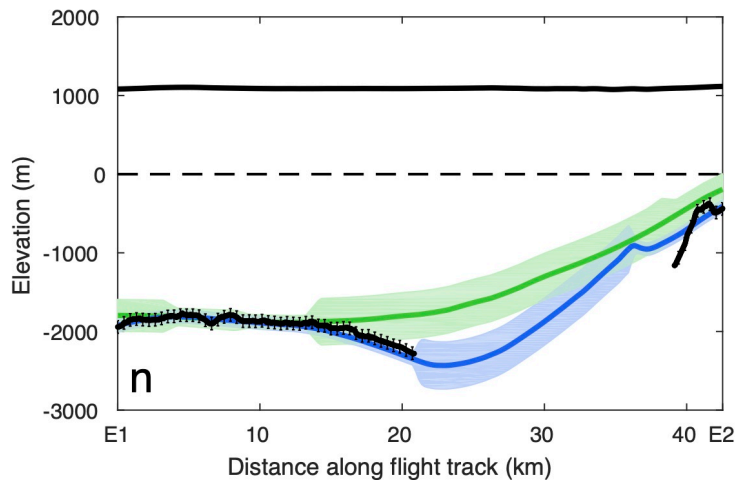


Brancato et al. 2019, submitted

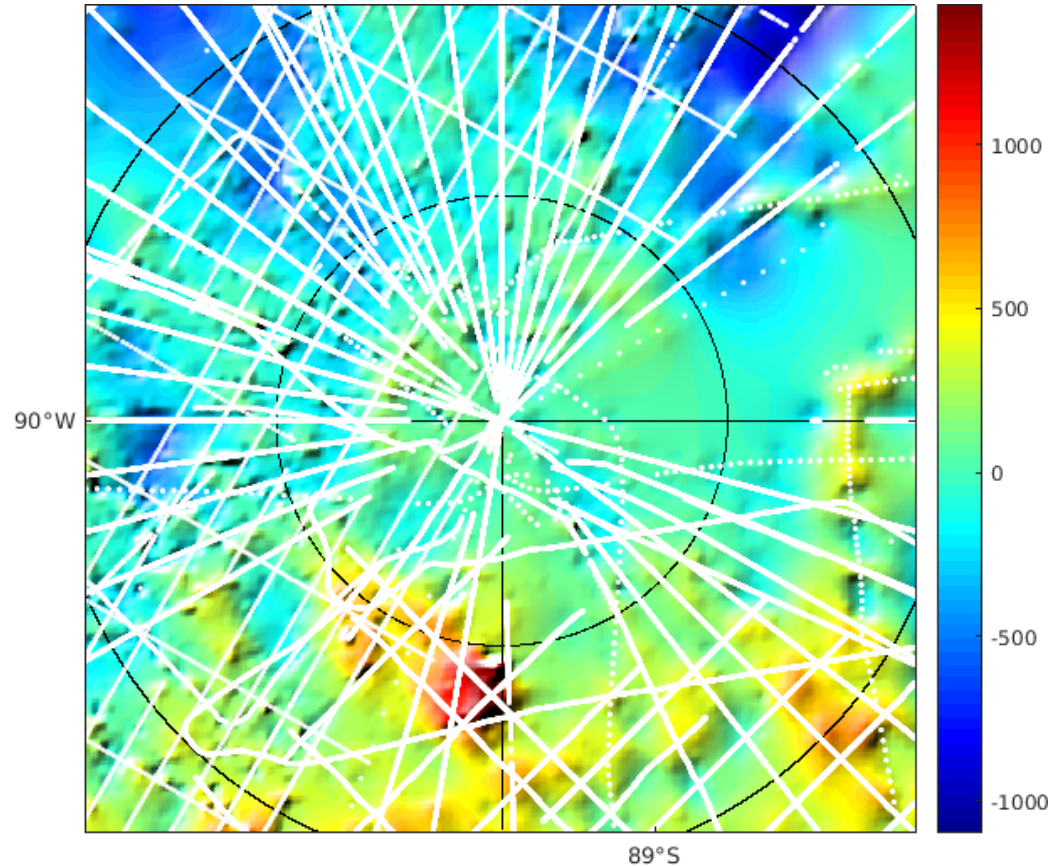
Recovery Ice Stream



bed elevation (m)

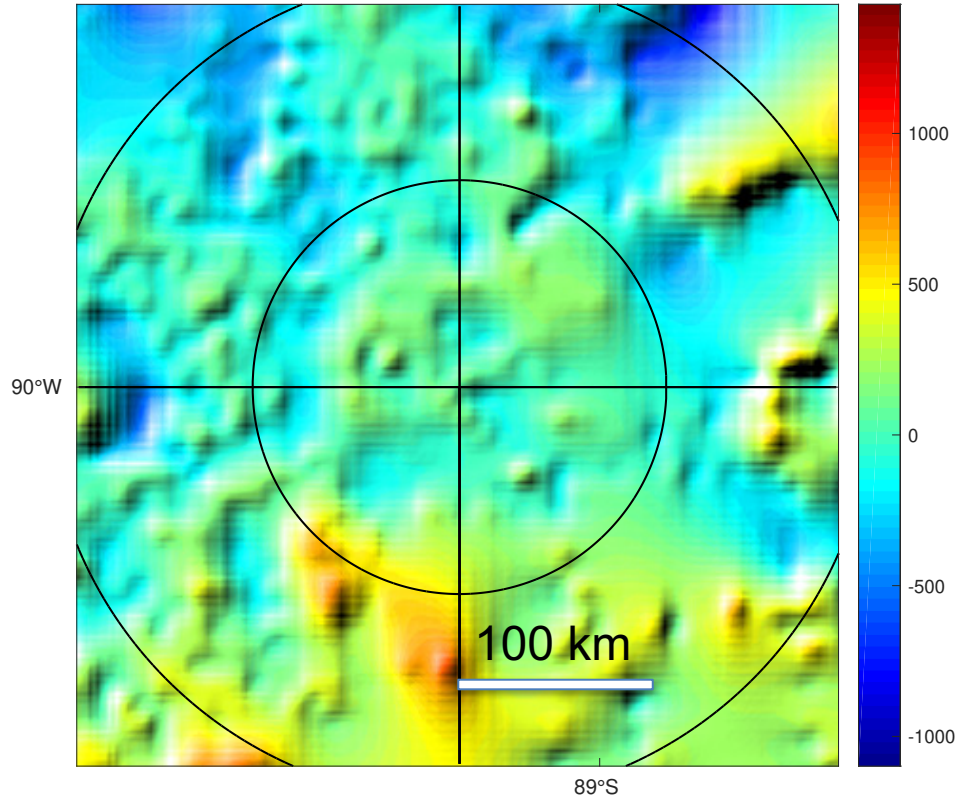


Bed topography around South Pole

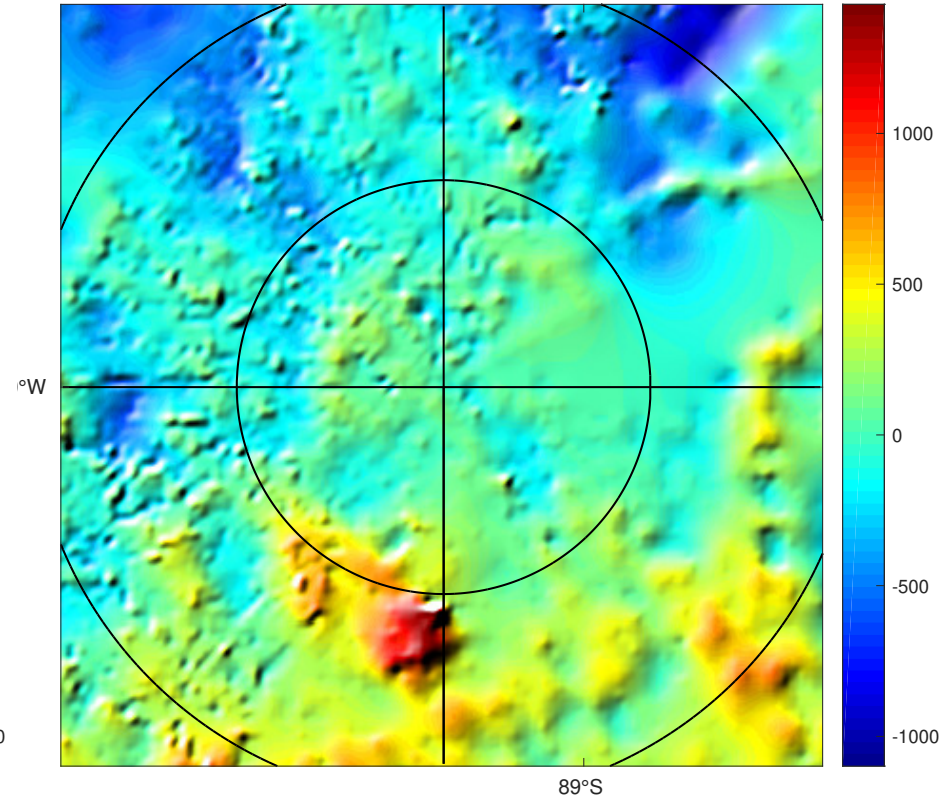


Bed topography around South Pole

Bedmap2



BedMachine



Thank you !
Questions ?

