

# Polarimetric radar sounding methods to characterise ice birefringence, fabric anisotropy, and flow history

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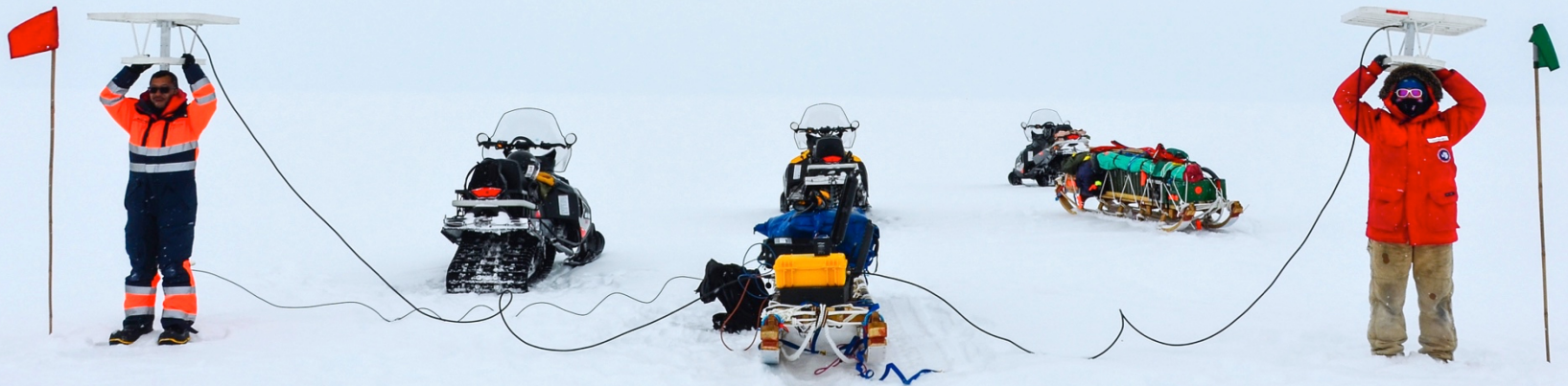
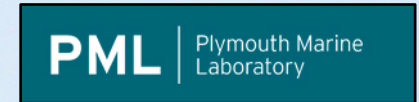
Scott Polar Research Institute

**Tom Jordan**

Plymouth Marine Laboratory

**Carlos Martín**

British Antarctic Survey



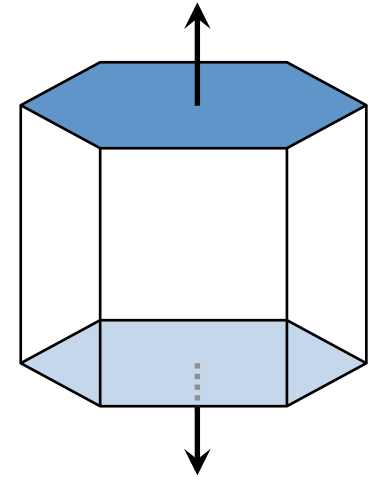
# What is ice fabric?

- Ice fabric represents the collective orientations of ice crystals
- Represented by a 2<sup>nd</sup> order orientation tensor with eigenvalues  $a$  along eigenvectors  $x$ :

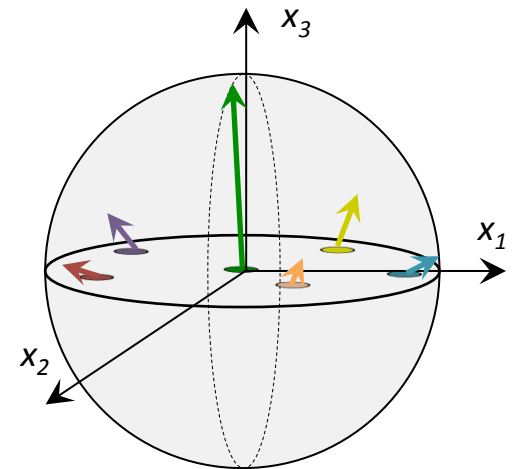
$$a_1 + a_2 + a_3 = 1$$

$$a_1 < a_2 < a_3 \text{ (radar, seismics)}$$

$$a_1 > a_2 > a_3 \text{ (ice cores)}$$



The  $c$ -axis of a single ice crystal

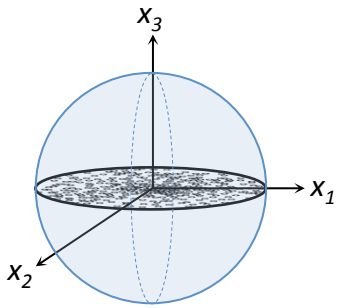


Ice fabric:  $c$ -axes of many ice crystals



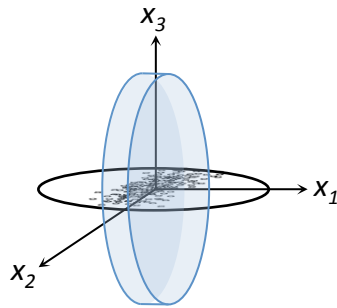
# Fabric anisotropy

- Fabric represents the deformation history of glacier ice
- “Complex history → complex fabric”



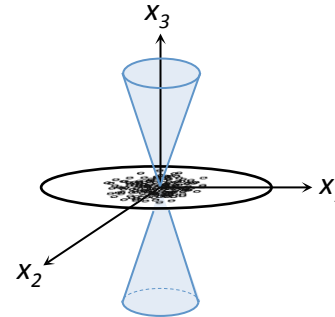
## Isotropic (near-random)

$a_1 \approx a_2 \approx a_3 \approx 1/3$   
 Near-surface firn/ice  
 Uniform deformation



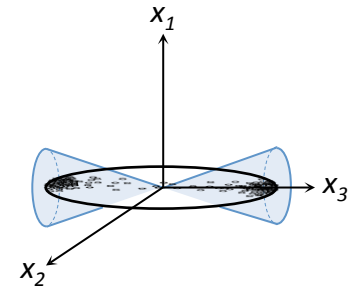
## Vertical girdle

$a_1 \ll a_2 \approx a_3$   
 Centre of glacier  
 Uniaxial compression,  
 longitudinal extension



## Vertical cluster

$a_1 \approx a_2 \ll a_3$   
 Near-bed ice  
 Planar simple shear



## Horizontal cluster

$a_1 \approx a_2 \ll a_3$   
 Glacier shear margin  
 Lateral simple shear

### Further reading

Azuma & Higashi (1985) AoG  
 Alley (1988) Nature

\* Shown are *general* examples (deviations occur!)

# Birefringence in ice

Radar is able to detect the horizontal components of fabric anisotropy due to the birefringence of ice as an effective medium

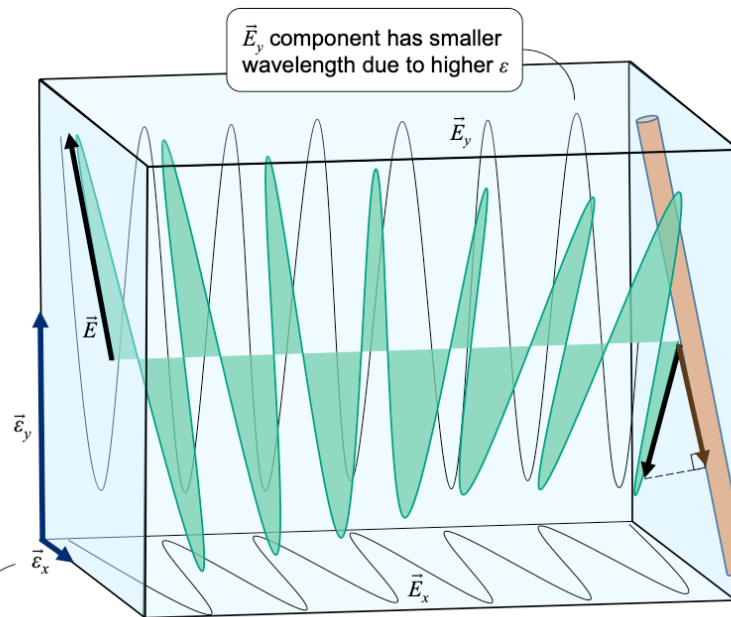
$$\epsilon = \begin{bmatrix} \epsilon_x & 0 & 0 \\ 0 & \epsilon_y & 0 \\ 0 & 0 & \epsilon_z \end{bmatrix}$$

$$= \begin{bmatrix} \epsilon_{\perp} + a_1 \Delta\epsilon' & 0 & 0 \\ 0 & \epsilon_{\perp} + a_2 \Delta\epsilon' & 0 \\ 0 & 0 & \epsilon_{\perp} + a_3 \Delta\epsilon' \end{bmatrix}$$



$$\Delta\epsilon = \epsilon_y - \epsilon_x = (a_2 - a_1) \Delta\epsilon'$$

This equation relates the bulk (macroscopic) birefringence  $\Delta\epsilon$  to the microscopic (crystal) birefringence  $\Delta\epsilon'$



$\vec{E}_y$  component has smaller wavelength due to higher  $\epsilon$

$\vec{E}_{measured}$   
Only the component of the field aligned with the antenna is captured, resulting in loss

Permittivity of ice is dependent on crystal orientation. Here, permittivity in the  $y$  direction is larger than the  $x$  direction

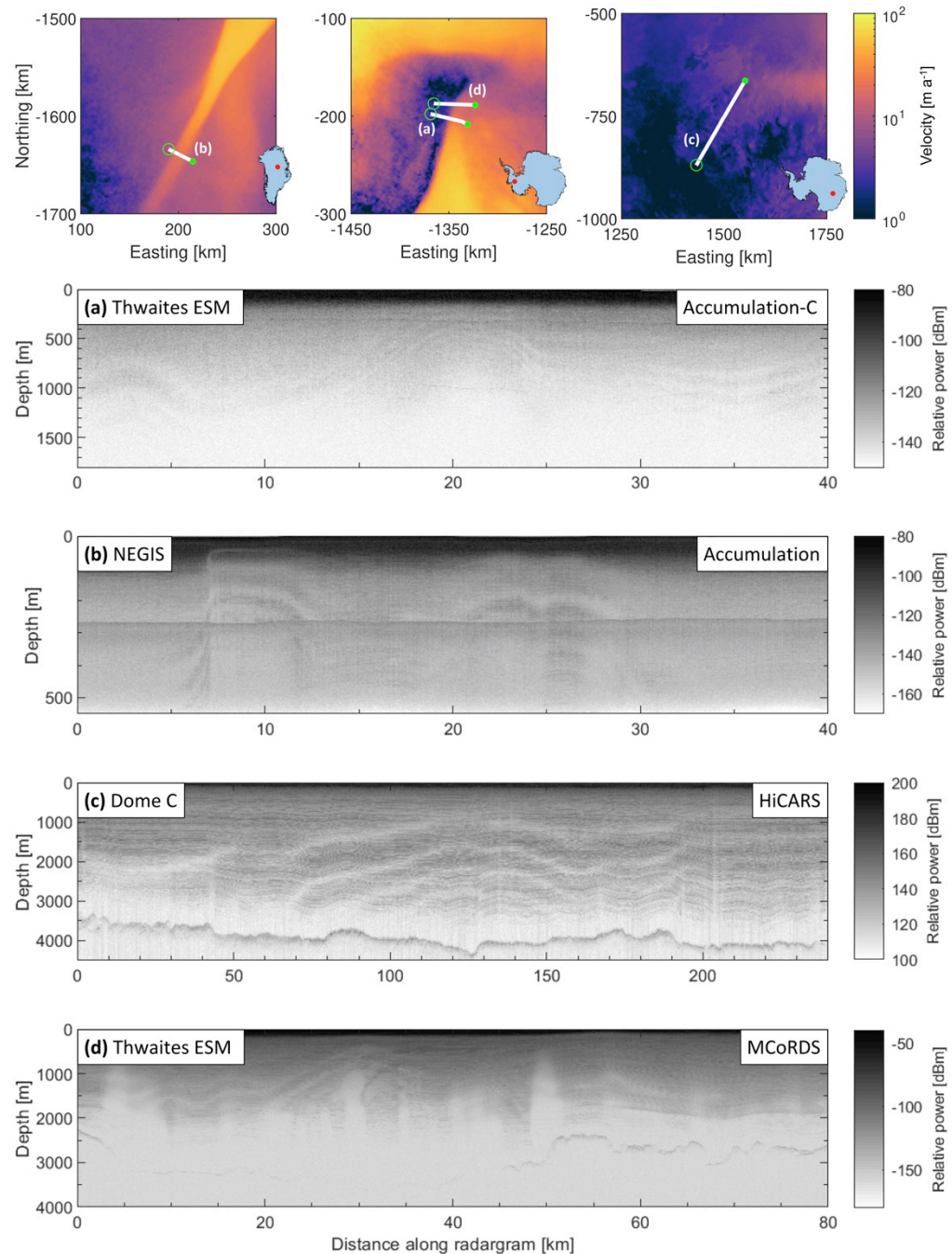
Difference in wavelength of  $x$  and  $y$  components causes rotation

## Further reading

Hargreaves (1978) JGlac  
Doake (1981) GJI  
Fujita et al. (2006) JGlac  
Young et al. (in review) JGR

# Birefringence in ice

- In radargrams, periodic patterns appear as a result of birefringence
- Radar must be angled off-parallel and off-perpendicular to fabric axes!



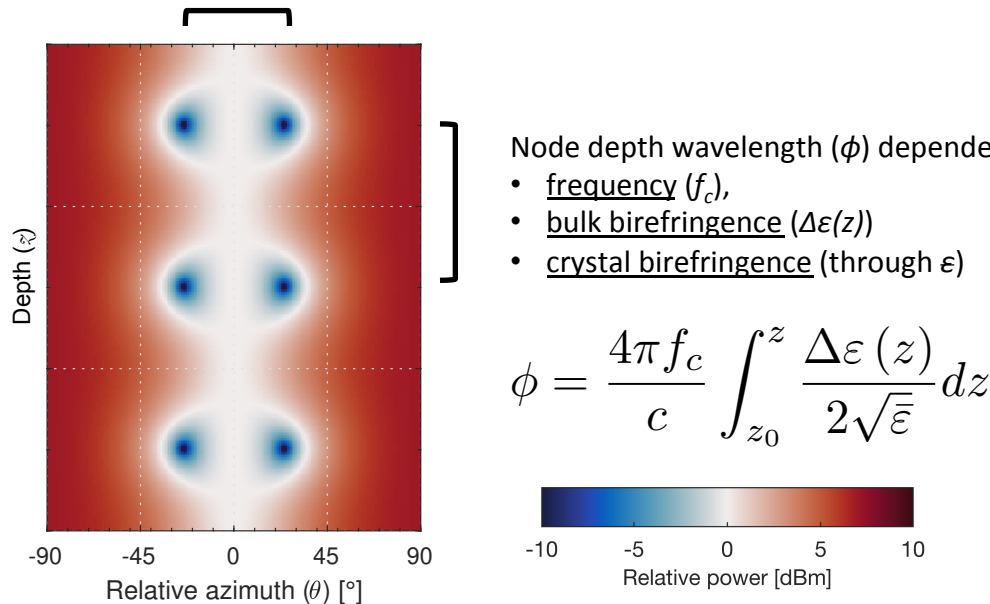
# Polarimetric backscatter model

$$S(\theta) = \underbrace{\left[ \frac{\exp(jk_0 z)}{4\pi z} \right]^2}_{\text{free-space propagation}} \cdot \underbrace{\left[ \prod_{i=1}^N (RT R')_{\theta, \varepsilon, N+1-i} \right]}_{\text{received (upward) propagation}} \cdot \underbrace{[R \Gamma R']_{\theta, \beta, N}}_{\text{boundary scattering}} \cdot \underbrace{\left[ \prod_{i=1}^N (RT R')_{\theta, \varepsilon, i} \right]}_{\text{transmitted (downward) propagation}}$$

$$R = \begin{bmatrix} T_x & 0 \\ 0 & T_y \end{bmatrix}$$

$$\Gamma = \begin{bmatrix} r & 0 \\ 0 & 1 \end{bmatrix}$$

Node azimuth separation dependent on anisotropic scattering ( $r$ )



**Further reading**

Fujita et al. (2006) JGLac  
Young et al. (*in review*) JGR

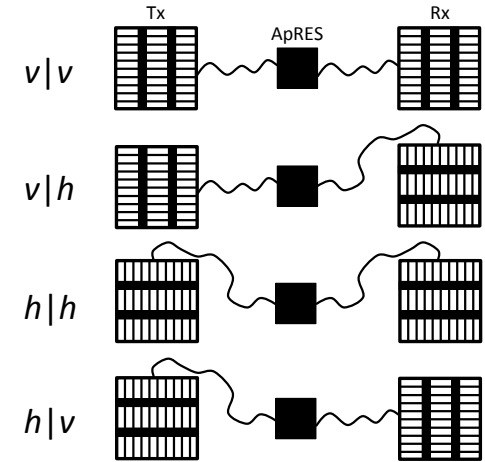


# Polarimetric radar sounding

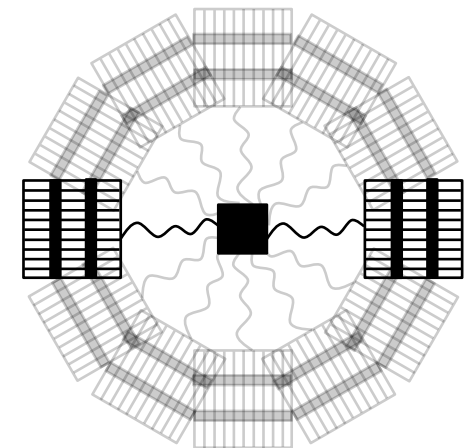
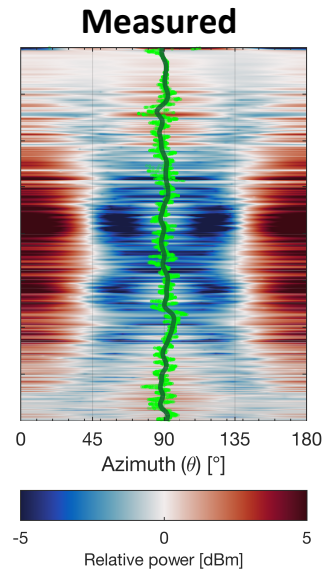
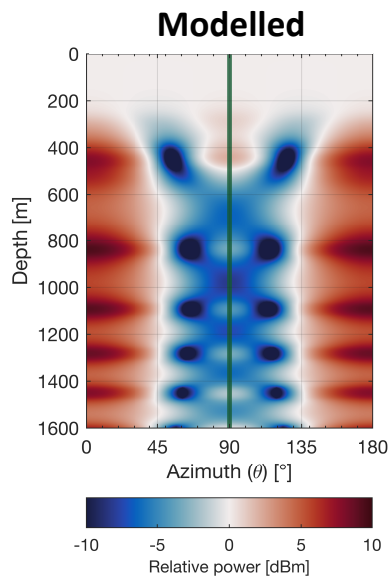
Application to radar sounding using linearly-polarised antennas can detect azimuthal (x-y) fabric asymmetry

$$S(\theta) = \begin{bmatrix} \cos\theta & \sin\theta \\ -\sin\theta & \cos\theta \end{bmatrix} \begin{bmatrix} s_{hh} & s_{vh} \\ s_{hv} & s_{vv} \end{bmatrix} \begin{bmatrix} \cos\theta & -\sin\theta \\ \sin\theta & \cos\theta \end{bmatrix}$$

$$= \begin{bmatrix} s_{hh} \cos^2\theta + (s_{vh} + s_{hv}) \sin\theta \cos\theta + s_{vv} \sin^2\theta & s_{hv} \cos^2\theta + (s_{vv} - s_{hh}) \sin\theta \cos\theta - s_{vh} \sin^2\theta \\ s_{vh} \cos^2\theta + (s_{vv} - s_{hh}) \sin\theta \cos\theta - s_{hv} \sin^2\theta & s_{vv} \cos^2\theta - (s_{vh} + s_{hv}) \sin\theta \cos\theta + s_{hh} \sin^2\theta \end{bmatrix}$$



Quad-pol setup



Azimuthal rotation setup

## Further reading

Brisbourne et al. (2019) JGR  
Young et al. (2020) TCD

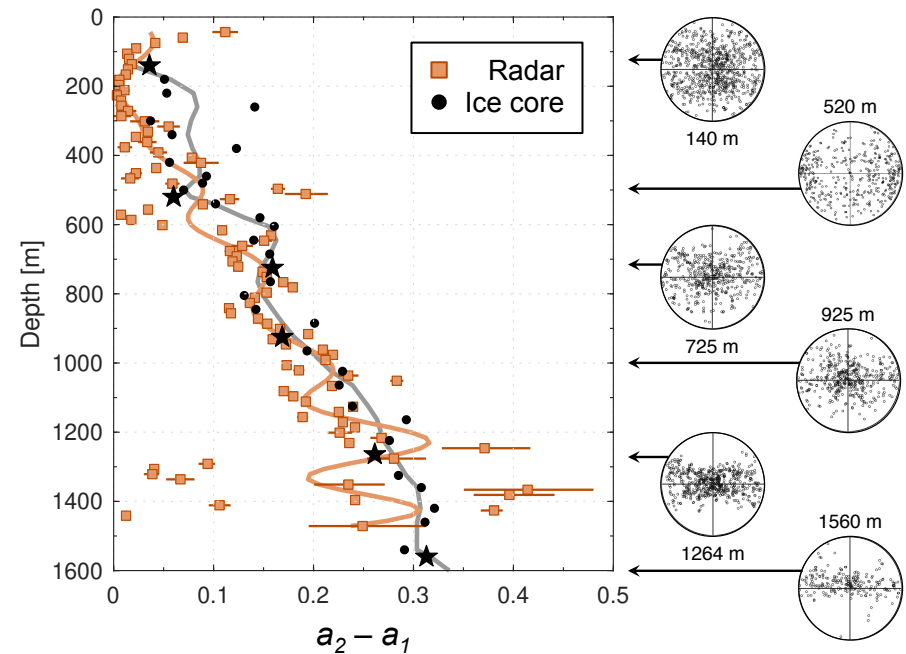
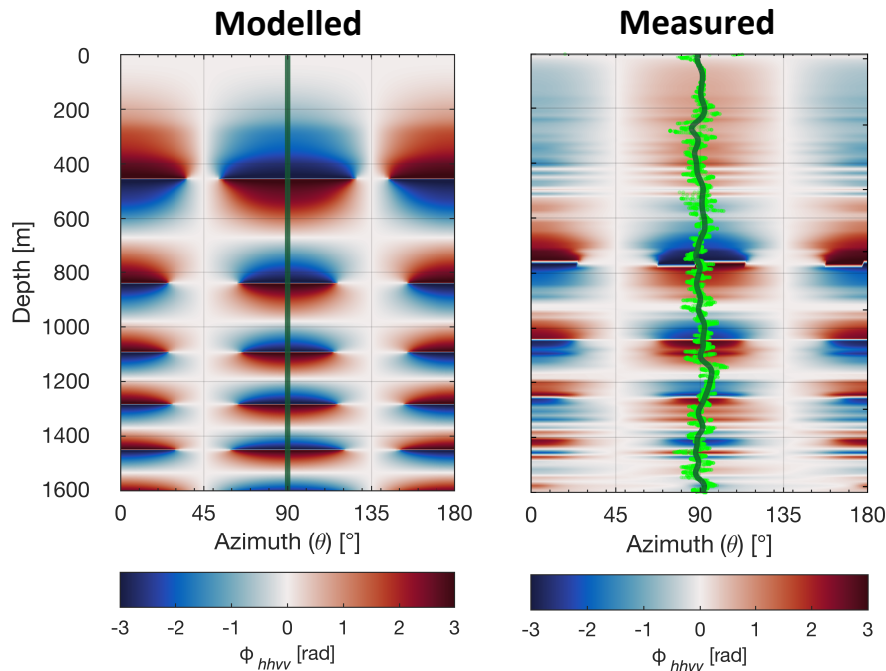
# Polarimetric coherence

Application of polarimetric coherence to the effective medium model quantifies the azimuthal fabric asymmetry

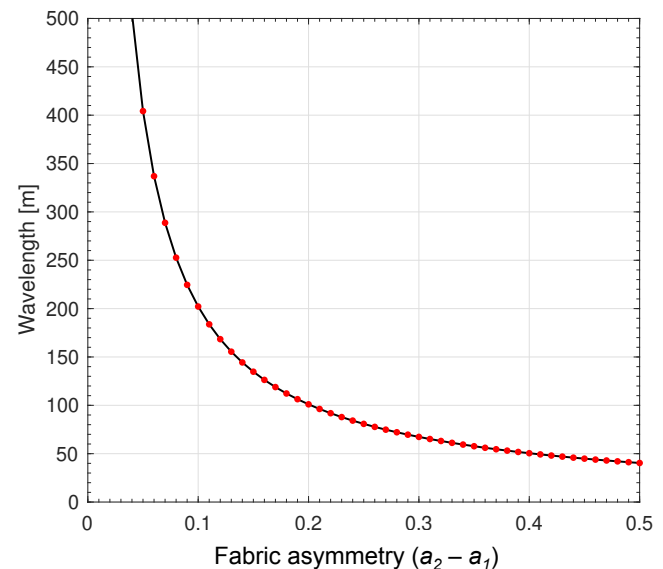
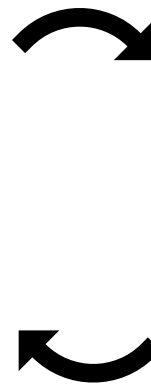
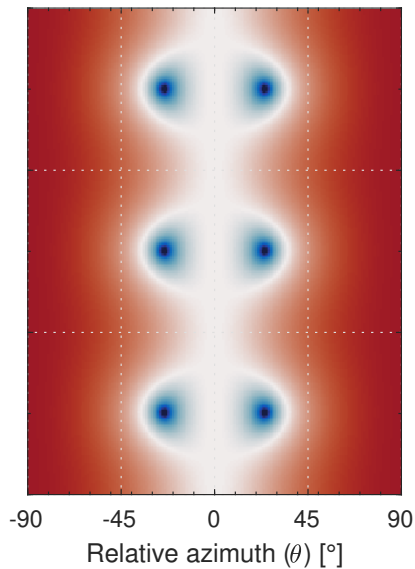
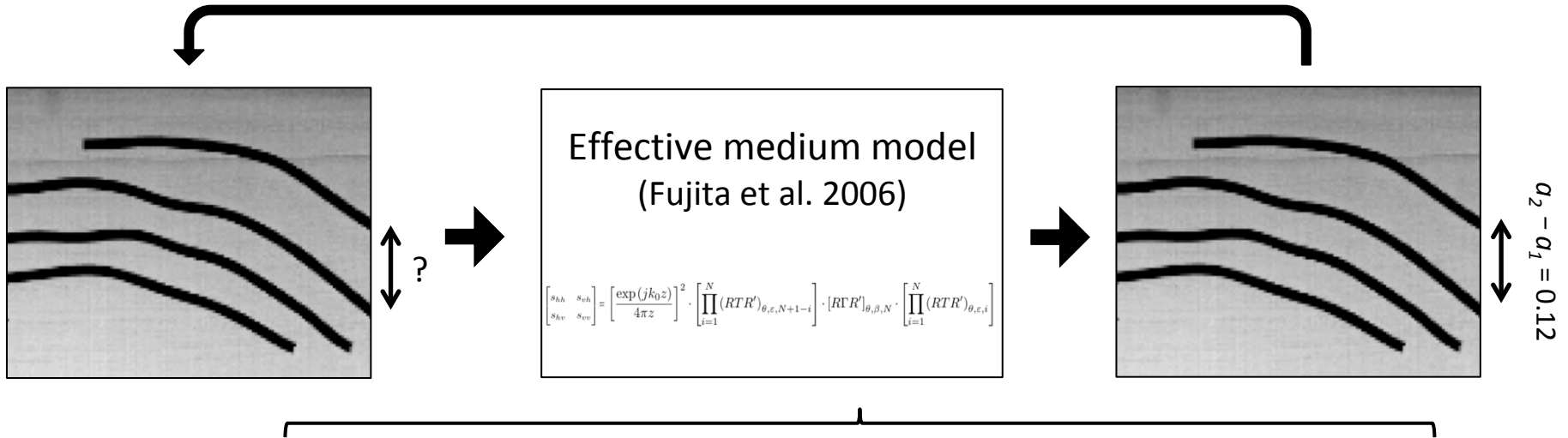
$$c_{hhvv}^* = \frac{\sum_{i=1}^N s_{hh,i} \cdot s_{vv,i}^*}{\sqrt{\sum_{i=1}^N |s_{hh,i}|^2} \sqrt{\sum_{i=1}^N |s_{vv,i}|^2}}$$

$$\phi_{hhvv}^* = \arg(c_{hhvv})$$

$$a_2 - a_1 = \frac{c}{4\pi f_c} \frac{2\sqrt{\epsilon}}{f(\nu)\Delta\epsilon'} \left| \frac{d\phi_{hhvv}}{dz} \right|$$



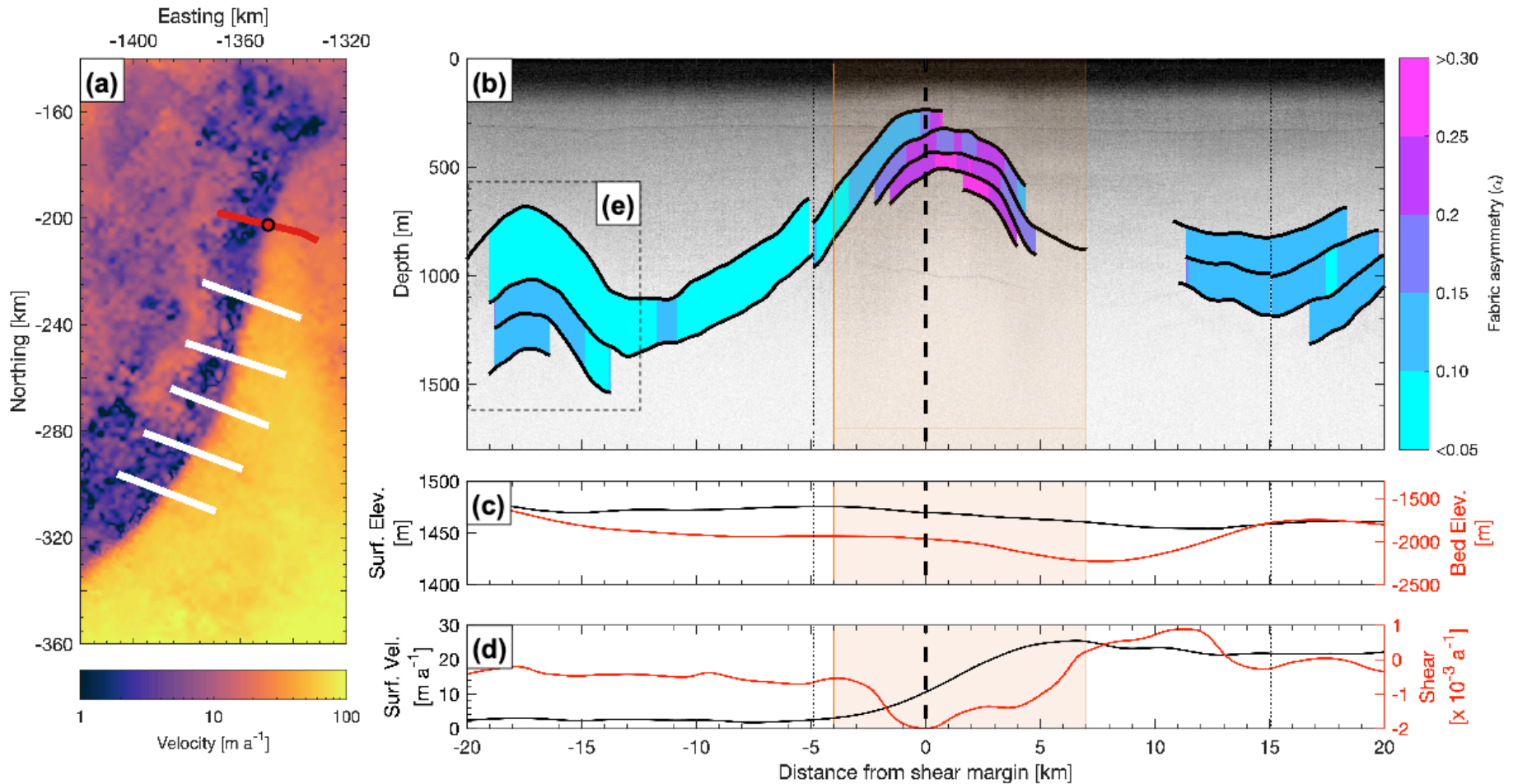
# Application: Thwaites Glacier



Further reading

Young et al. (*in review*) JGR

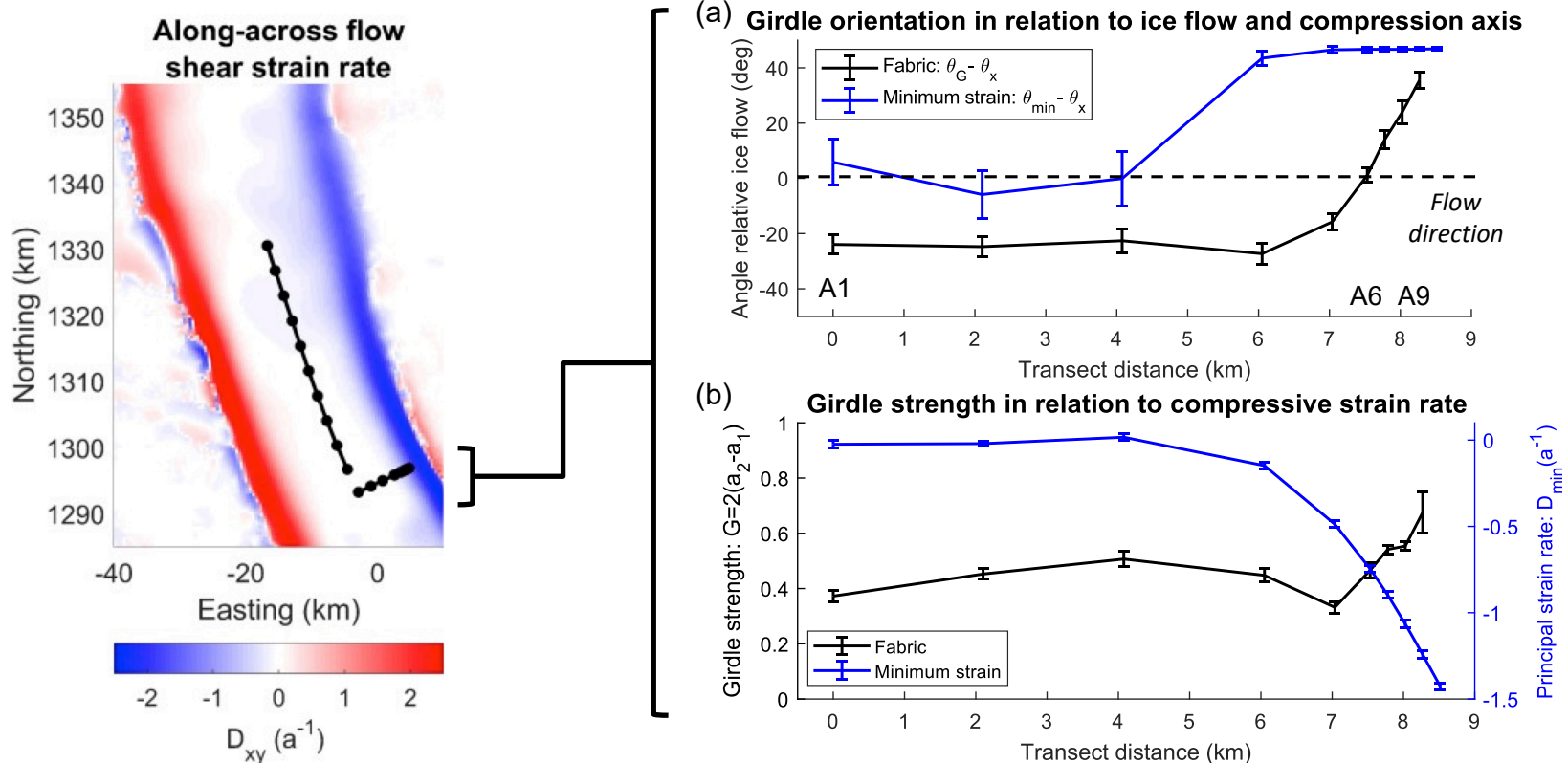
# Application: Thwaites Glacier





# Application: Rutford Ice Stream

pRES observations approaching the shear margin of Rutford Ice Stream reveals increasing fabric asymmetry and axis rotation

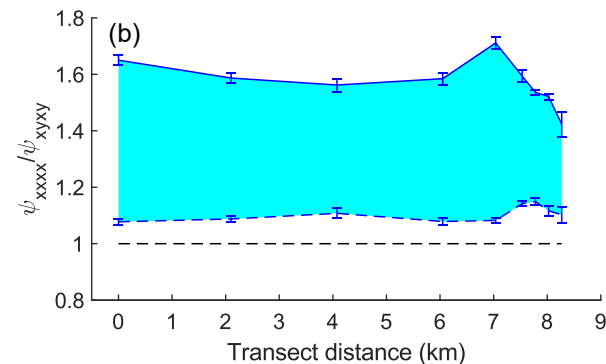
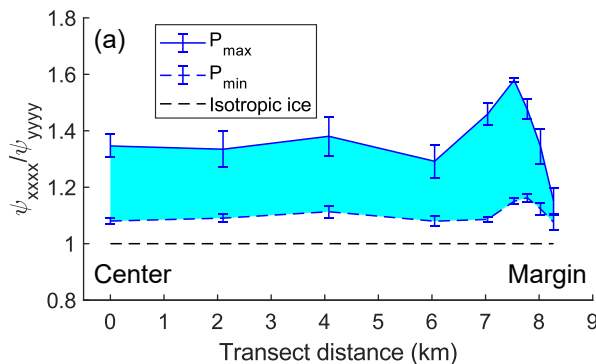


# Application: Rutford Ice Stream

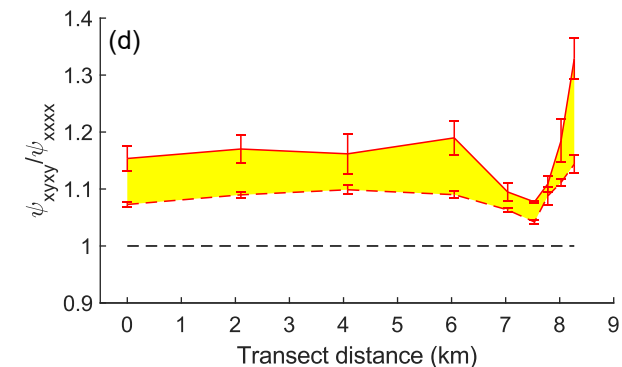
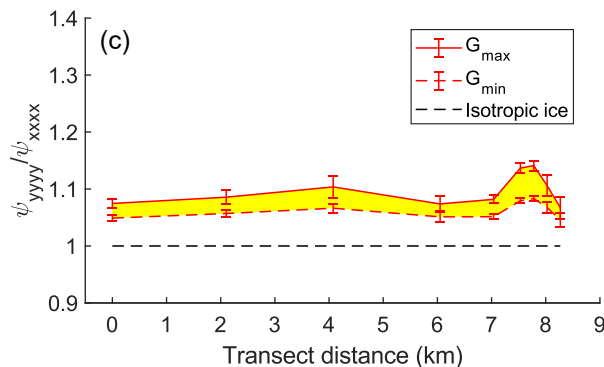
Radar fabric measurements can be used to parameterise an anisotropic flow law via the fluidity tensor  $\psi$

$$\begin{pmatrix} D_{11} \\ D_{22} \\ D_{33} \\ D_{12} \\ D_{13} \\ D_{23} \end{pmatrix} = \psi_0 \begin{pmatrix} \psi_{1111} & \psi_{1122} & \psi_{1133} & 0 & 0 & 0 \\ \psi_{1122} & \psi_{2222} & \psi_{2233} & 0 & 0 & 0 \\ \psi_{1133} & \psi_{2233} & \psi_{3333} & 0 & 0 & 0 \\ 0 & 0 & 0 & \psi_{1212} & 0 & 0 \\ 0 & 0 & 0 & 0 & \psi_{1313} & 0 \\ 0 & 0 & 0 & 0 & 0 & \psi_{2323} \end{pmatrix} \begin{pmatrix} \bar{S}_{11} \\ \bar{S}_{22} \\ \bar{S}_{33} \\ \bar{S}_{12} \\ \bar{S}_{13} \\ \bar{S}_{23} \end{pmatrix}$$

Anisotropy of ice rheology for a vertical girdle: Transect A, U1

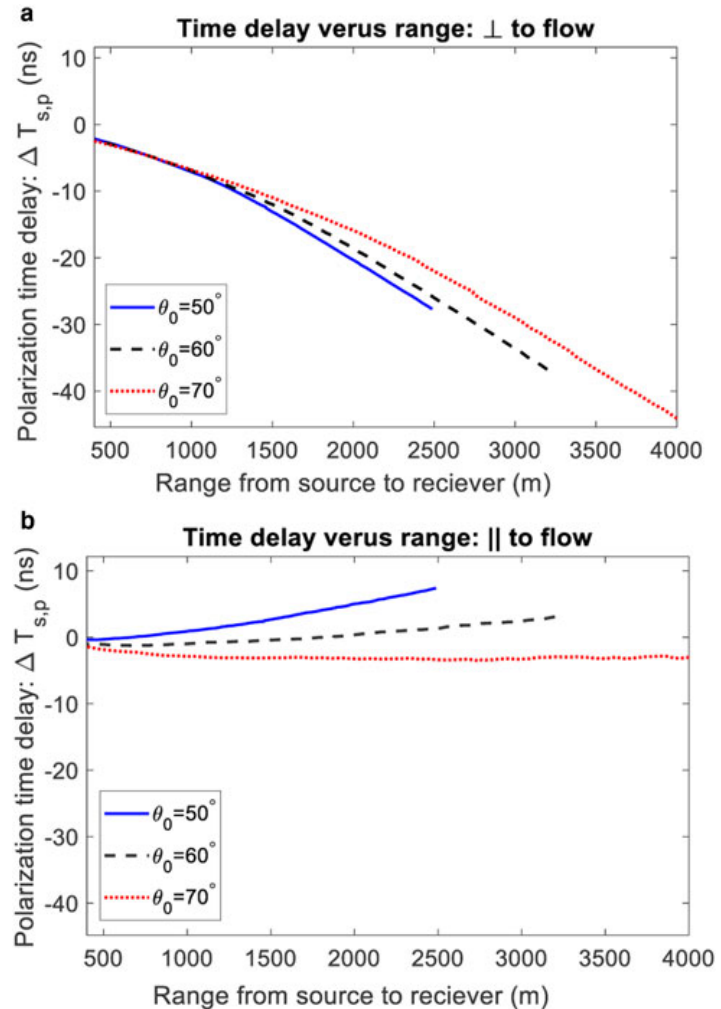
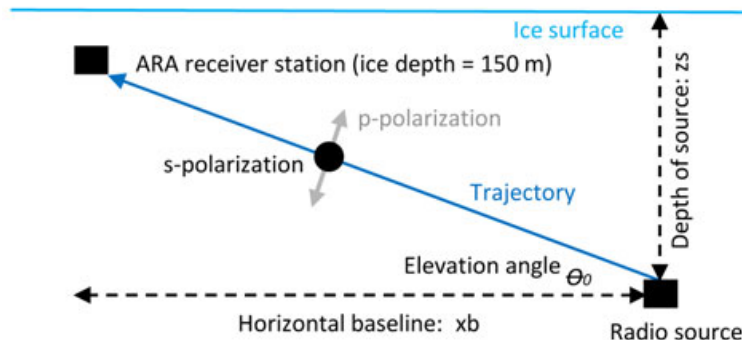


Anisotropy of ice rheology for a horizontal pole: Transect A, U1



# Application: ARA Neutrino Detection

Because Cherenkov radiation occurs within applicable radar frequencies ( $\sim 150 - 800$  MHz), the effective medium model can be repurposed to model oblique propagation delay and aid neutrino energy reconstruction



# Proposals for future work at South Pole

- Quantifying depth-space variations in fabric strength and orientation across IceCube domain
- Generalising Jordan et al. (2020)'s model framework for neutrino detection for off-axis alignment
- Bistatic radar surveys to resolve  $a_3$
- Anisotropic flow parameterisation and modelling of South Pole domain