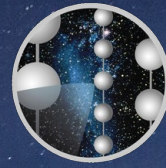


JOHANNES GUTENBERG  
UNIVERSITÄT MAINZ



WISCONSIN  
UNIVERSITY OF WISCONSIN-MADISON



ICECUBE  
SOUTH POLE NEUTRINO OBSERVATORY

# *The IceCube ice anisotropy*

Martin Rongen,  
Dmitry Chirkin

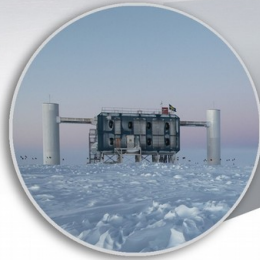
Second IceCube Polar Science Workshop 2021





# The IceCube Neutrino Observatory

*Design and construction*



## IceCube Laboratory

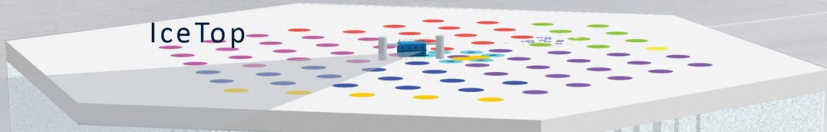
Data is collected here and sent by satellite to the data warehouse at UW–Madison



## Digital Optical Module (DOM)

5,160 DOMs deployed in the ice

50 m



IceTop

1450 m

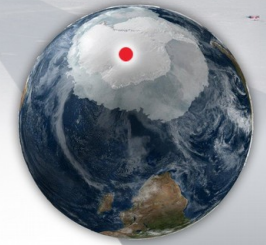
86 strings of DOMs, set 125 meters apart

2450 m

IceCube detector

DeepCore

Antarctic bedrock

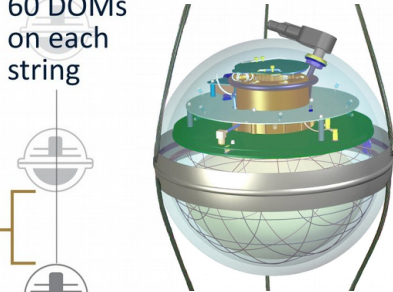


## Amundsen–Scott South Pole Station, Antarctica

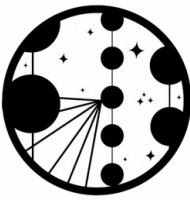
A National Science Foundation-managed research facility

60 DOMs on each string

DOMs are 17 meters apart

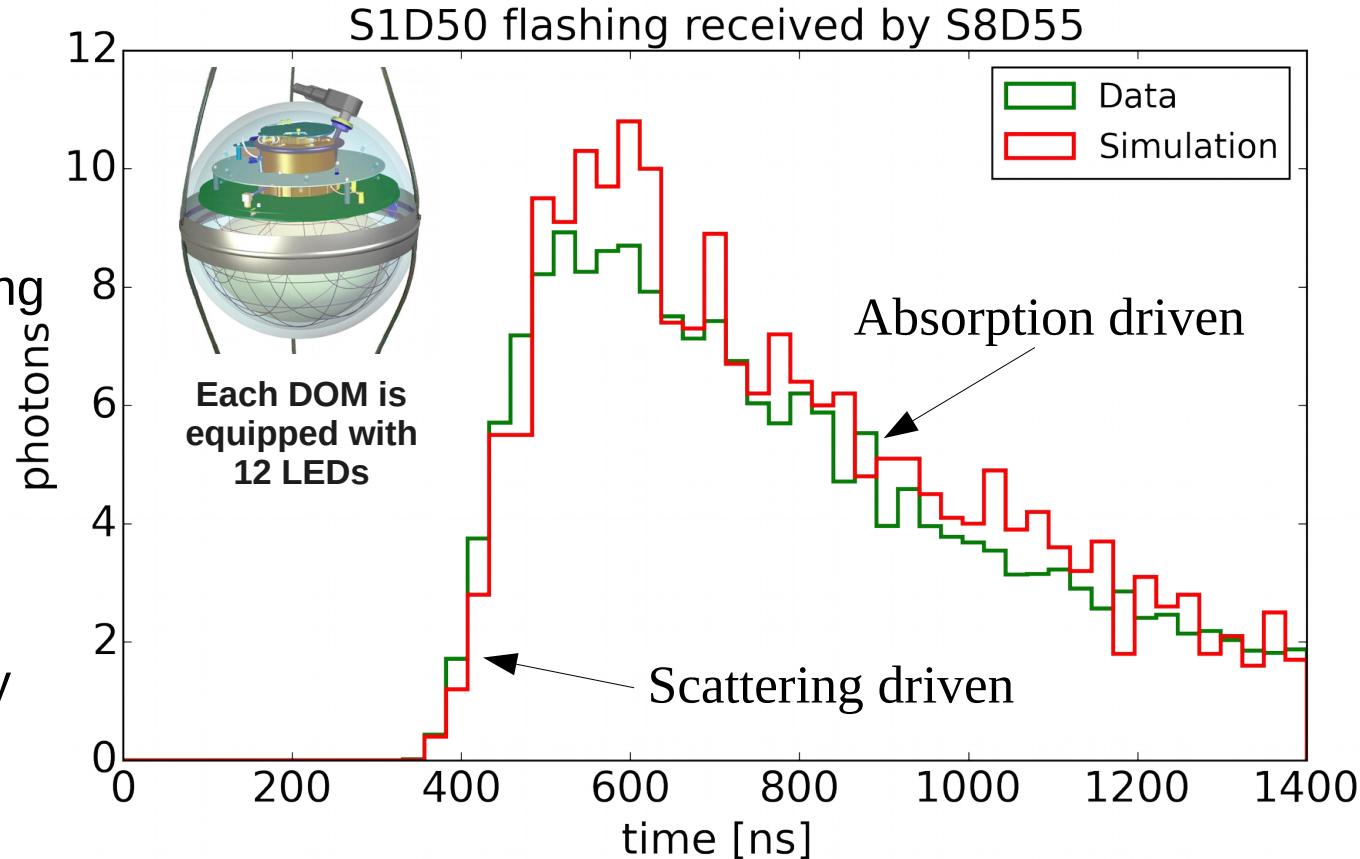


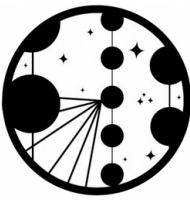
Each DOM is equipped with a 10" PMT and 12 LEDs



# Light curve sensitivity

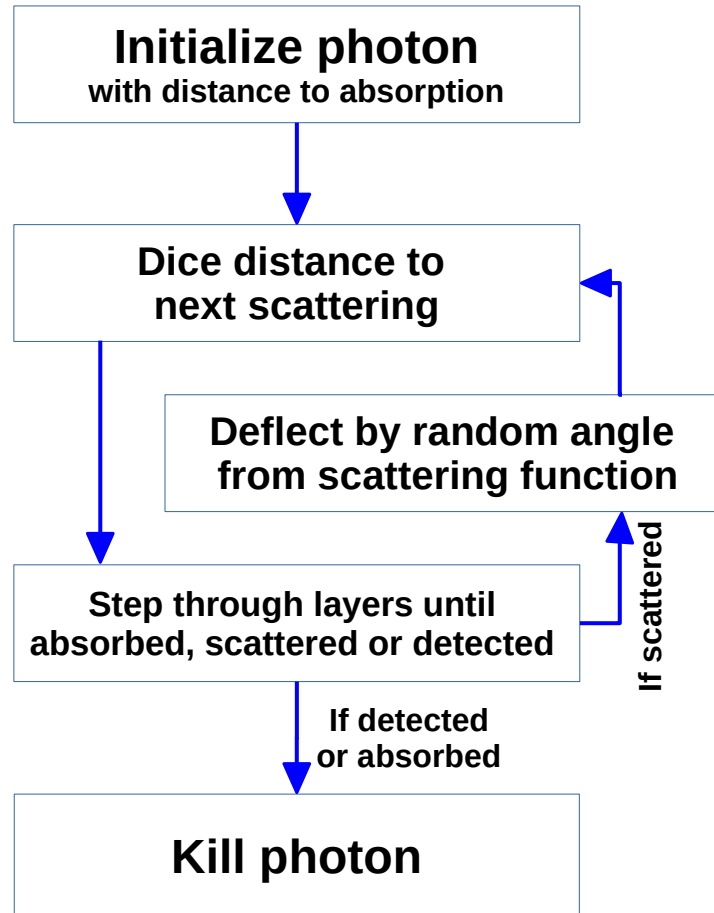
- Observation of photon arrival time distributions from pulsed light sources, allows for measurement of absolute absorption & scattering lengths
- Normalization independent, but observations at different distances help
- Distributions badly modeled by analytic random walk  
→ full simulation needed





# Light curve sensitivity

- Observation of photon arrival time distributions from pulsed light sources, allows for measurement of absolute absorption & scattering lengths
- Normalization independent, but observations at different distances help
- Distributions badly modeled by analytic random walk  
→ full simulation needed



Photon propagator code (PPC)



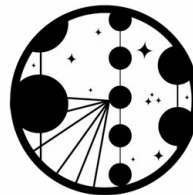
On GPUs ~250 times faster than CPUs

One full detector simulation takes 200-5000 GPU hours

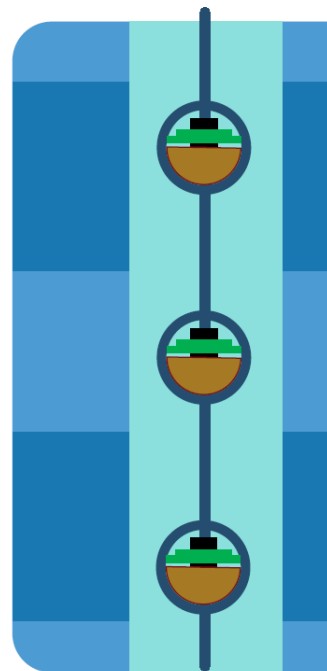
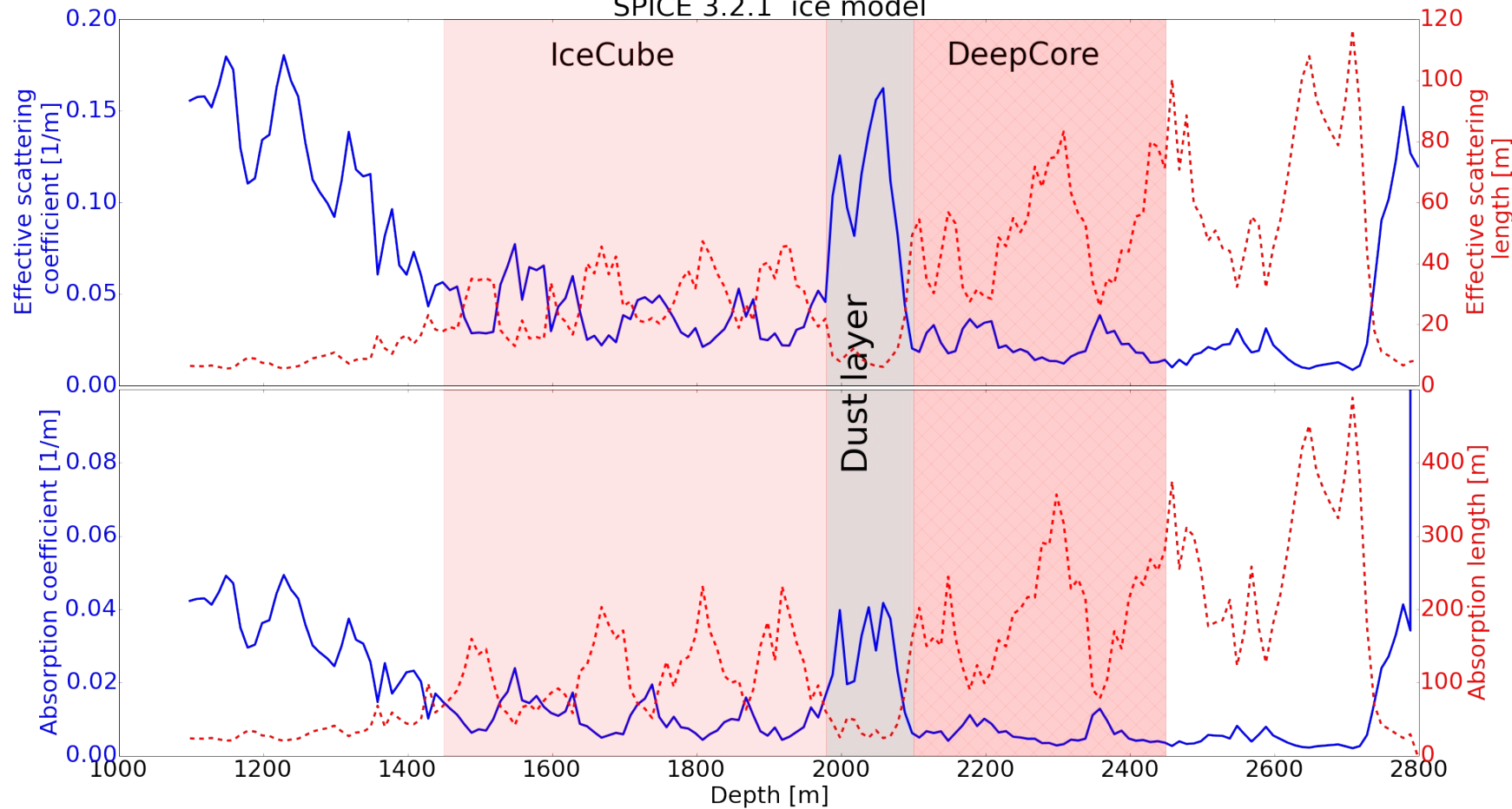


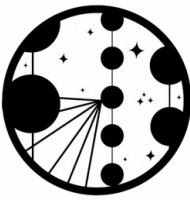
# Layered ice model (SPICE)

Rongen, Chirkin  
Polar Science 2021



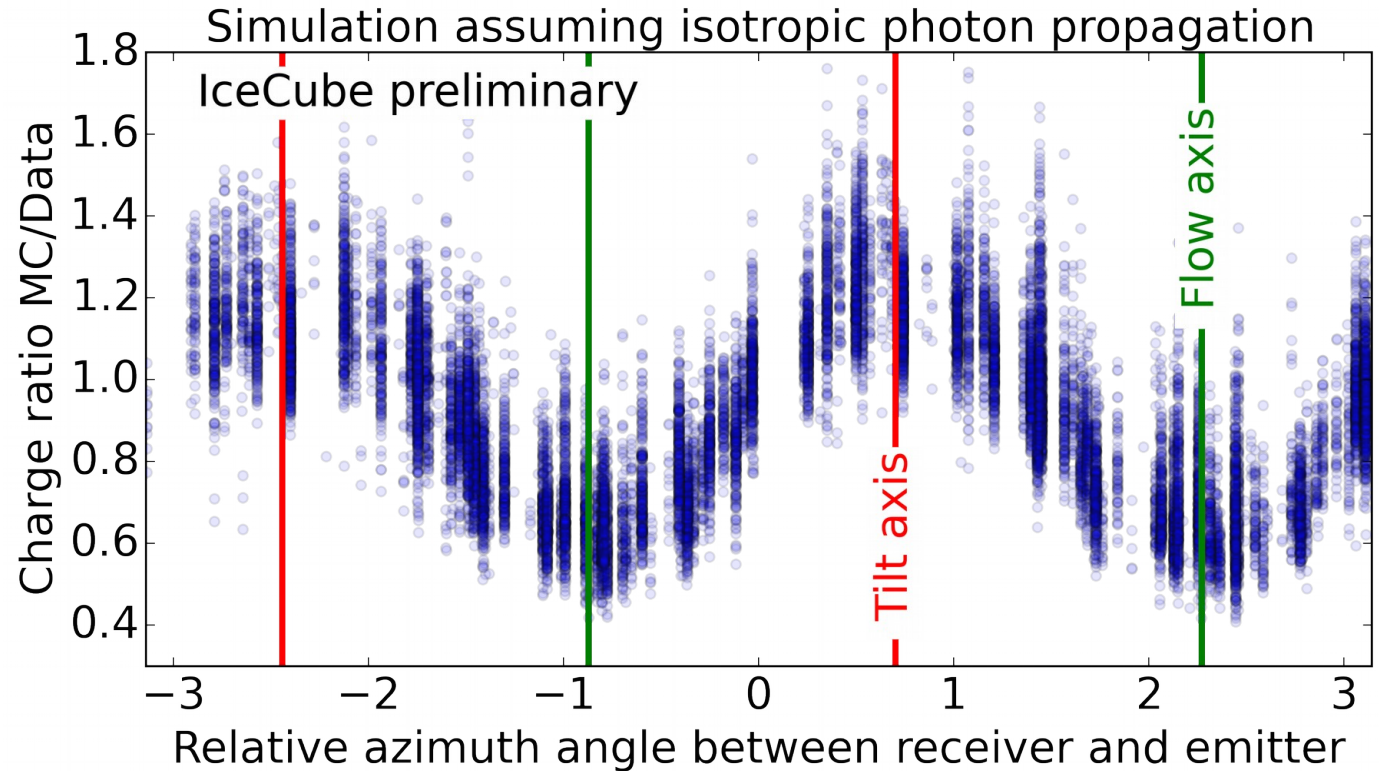
SPICE 3.2.1 ice model





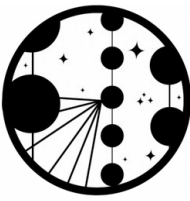
# The ice anisotropy effect

- Observed charge depends on orientation of receiver DOM with respect to emitter DOM
- Most charge seen along the flow axis, least orthogonal to it
- In 2013 implemented as a direction dependent modification of the scattering function



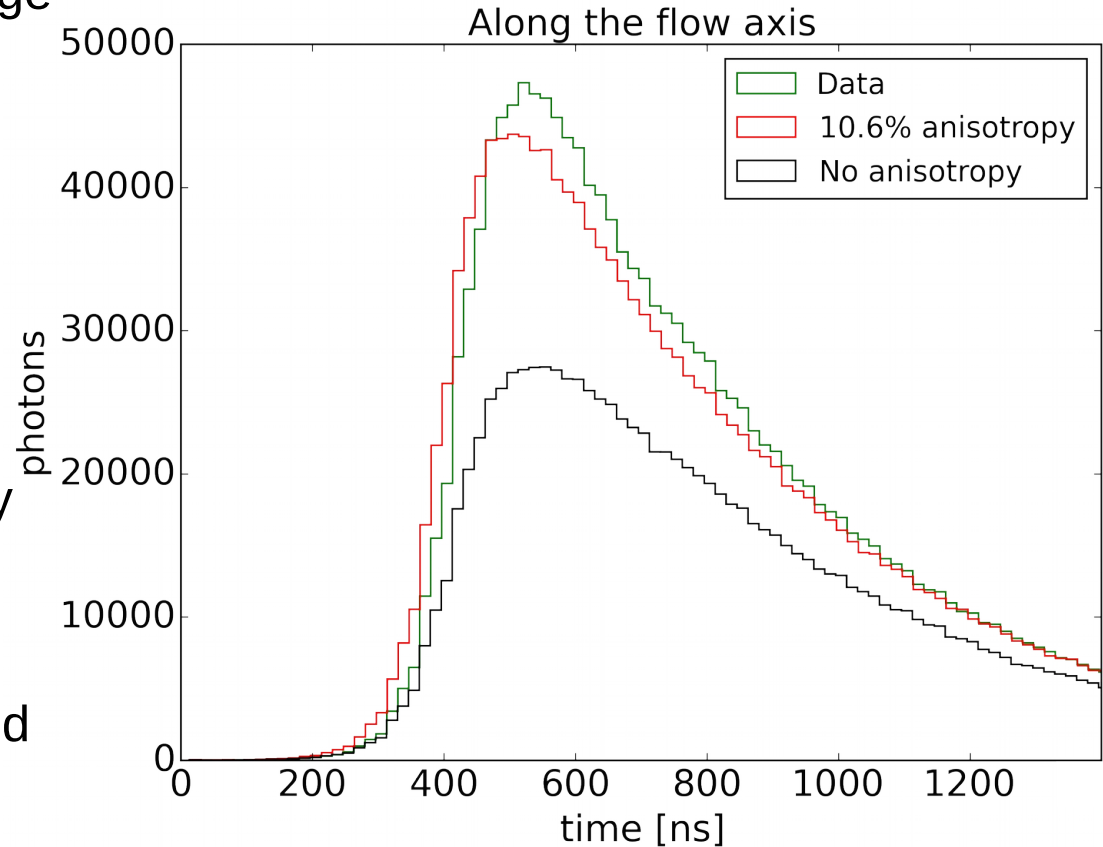
→ less scattering in the flow direction

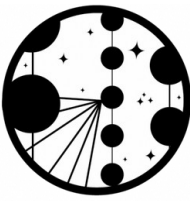




# Timing vs. intensity

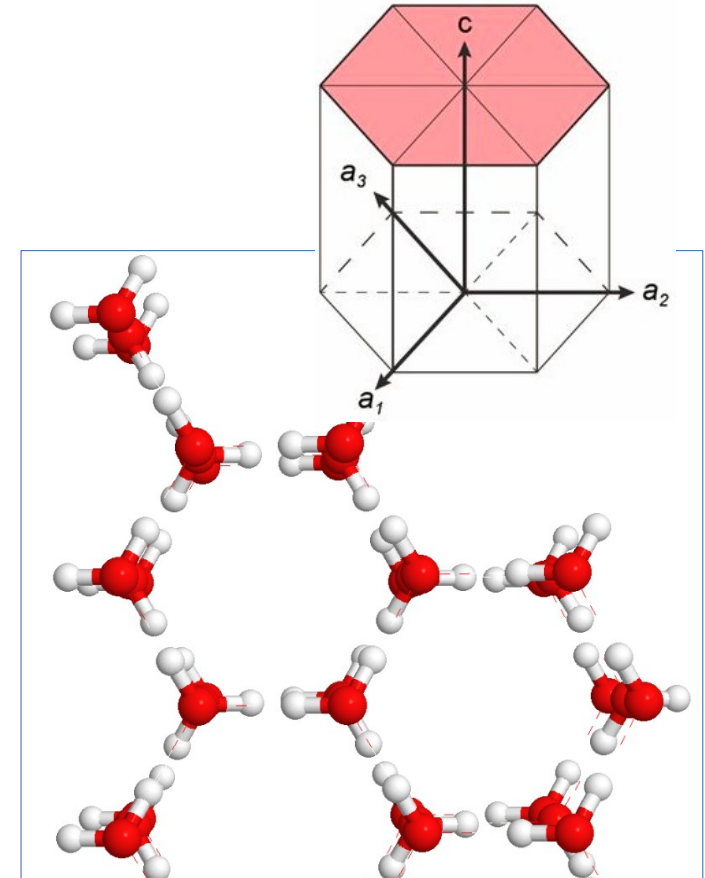
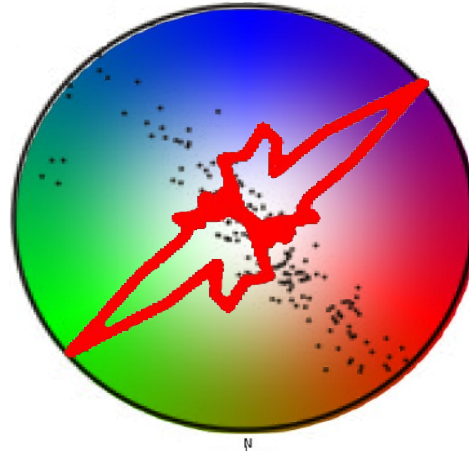
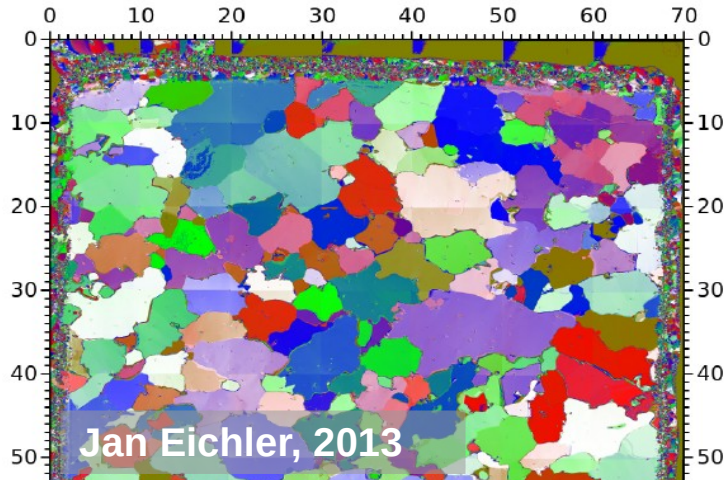
- Best fit does not fully compensate charge
- Compensating the intensity results in too-early rising edge
- Alternative model varying absorption instead of scattering was tested
- Better, but could also not fully match the light curves
- Also found that inhomogeneous impurity distributions can never lead to anisotropic Mie scattering
- Anisotropic absorption is possible, but a mechanism leading to the required strength is not understood



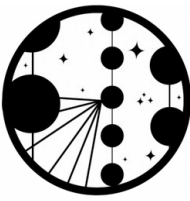


# c-axis and elongation

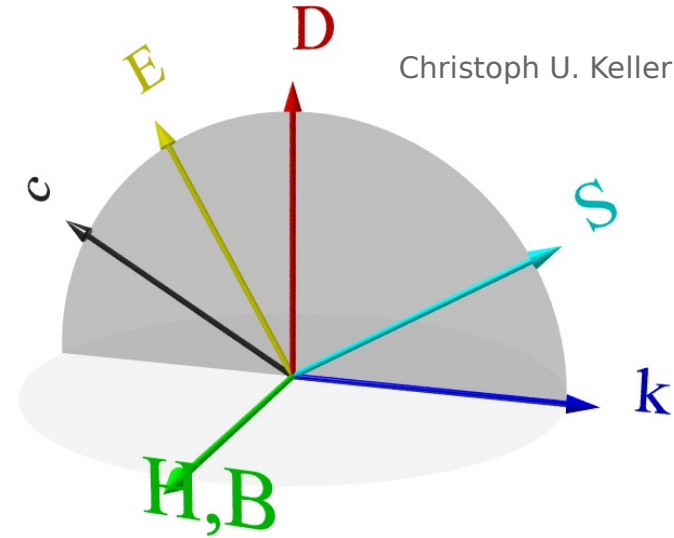
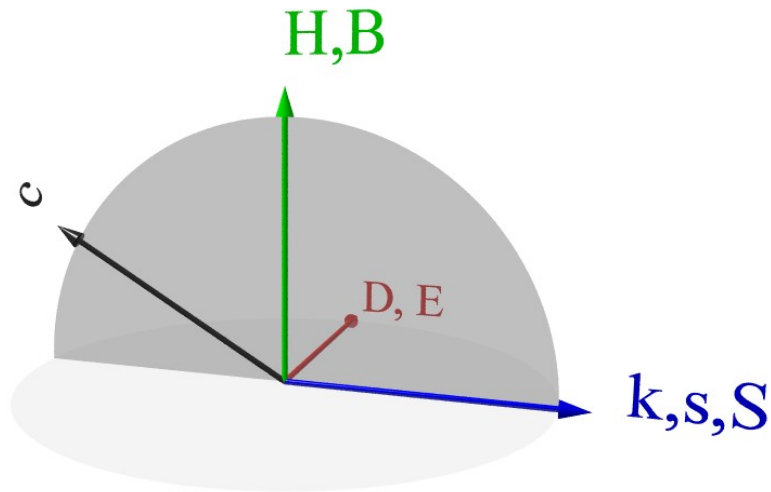
- Due to its hexagonal crystal lattice, each grain deforms essentially only by sliding of its basal planes
- An ensemble of grains under stress (such as flow) will elongate with the flow yielding a girdle of c-axes orthogonal to the flow





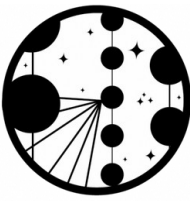


# Birefringence



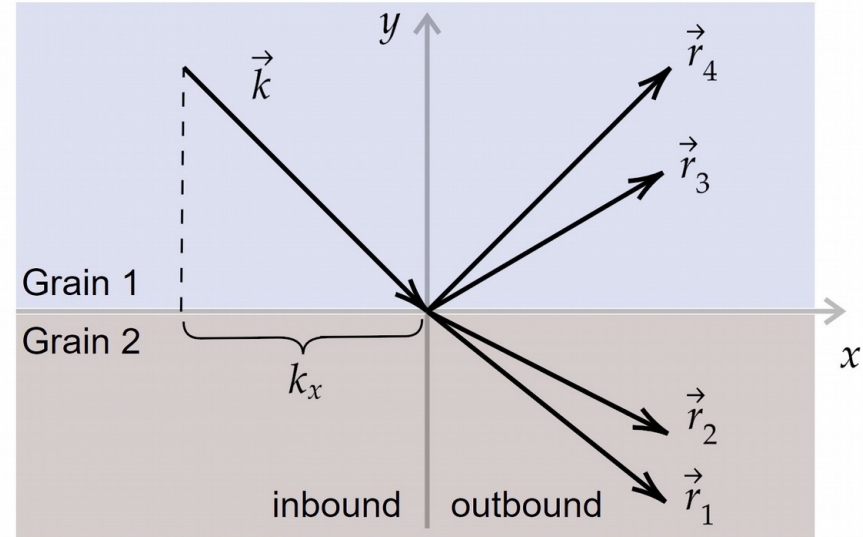
Christoph U. Keller

- Ice is a birefringent material:  
Light is split into an ordinary and an extraordinary ray with respect to the c-axis
- The refractive index of the extraordinary ray is direction dependent
- Extraordinary ray exhibits dispersion between the wave and Poynting vectors



# A single grain boundary

- At each grain boundary every ray is split into two reflected and two refracted rays one ordinary and one extraordinary ray each
- Wave vector component parallel to surface is conserved, norm is proportional to the refractive index
- Poynting vectors are derived from wave vectors and boundary conditions
- Outgoing ray is randomly sampled from Poynting vectors according to Poynting theorem (Poynting vector component through the boundary is conserved)



wavelength $\lambda$ (nm)	$n_o$	$n_e$
405	1.3185	1.3200
436	1.3161	1.3176
492	1.3128	1.3143
546	1.3105	1.3119
624	1.3091	1.3105
691	1.3067	1.3081

Physics of Ice,  
Victor F. Petrenko



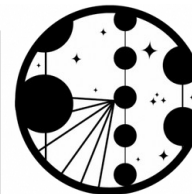
# c-axis sampling

Brief communication: Sampling c-axes distributions from the eigenvalues of ice fabric orientation tensors

Martin Rongen<sup>1</sup>

<sup>1</sup>RWTH Aachen University, Institute for Particle Physics III B, 52074 Aachen, Germany

Correspondence: Martin Rongen (rongen@physik.rwth-aachen.de)



Abstract. For simulation purposes, it can be necessary to generate an arbitrarily large sample of c-axes based on the commonly used descriptive statistics provided in publications of ice core analyses. This paper describes a technique to

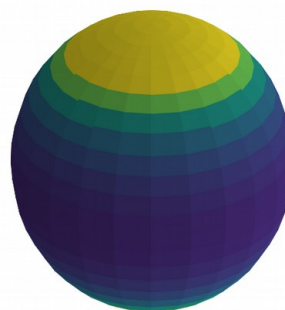
$$a = \begin{bmatrix} \sum n_{ix}^2 & \sum n_{ix} \cdot n_{iy} & \sum n_{ix} \cdot n_{iz} \\ \sum n_{iy} \cdot n_{ix} & \sum n_{iy}^2 & \sum n_{iy} \cdot n_{iz} \\ \sum n_{iz} \cdot n_{ix} & \sum n_{iz} \cdot n_{iy} & \sum n_{iz}^2 \end{bmatrix}$$

$$f(\theta, \phi) = C_w \exp(k \cdot \cos^2 \theta) \sin \theta$$

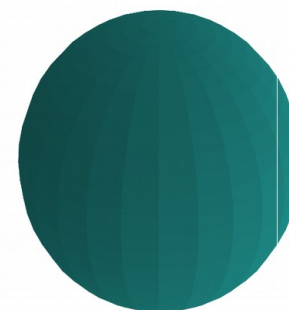
Watson distribution

$$C_w = 1 / \left( 4\pi \int_0^1 \exp(k \cdot u^2) du \right).$$

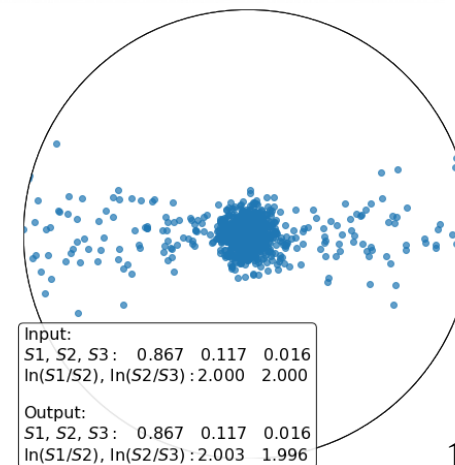
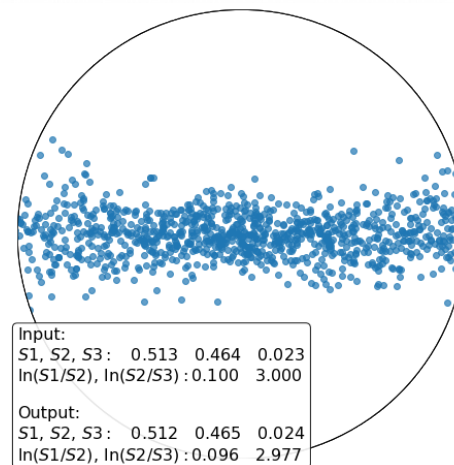
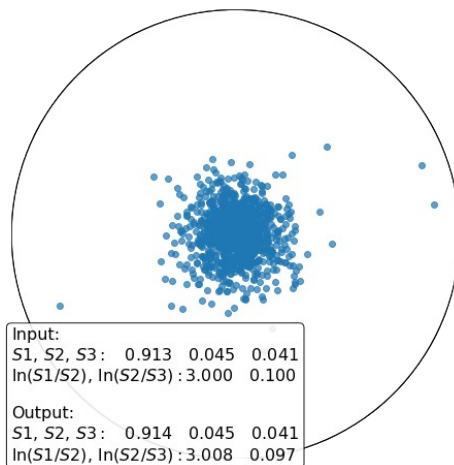
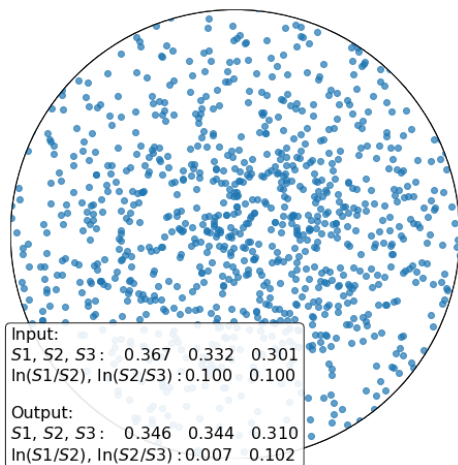
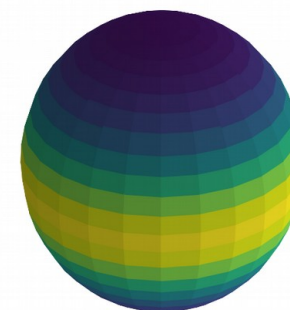
$k < 0$

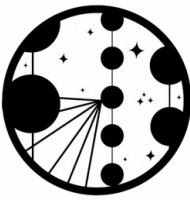


$k = 0$



$k > 0$





# Chord and surface orientations

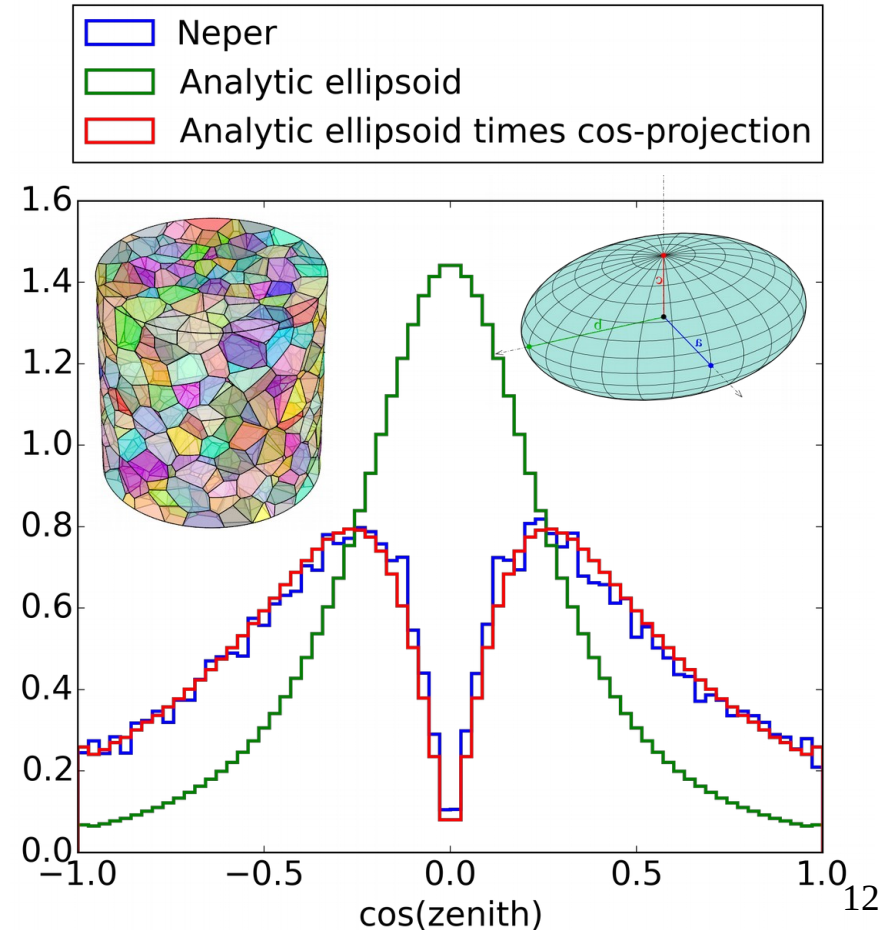
To be able to simulate propagating through a polycrystal we need to further know:

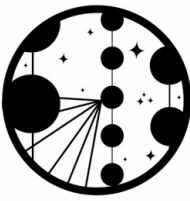
1. The average distance spend in a grain
2. The distribution of boundary orientations

For elongated grains these quantities are direction dependent.

Quantities computed on an average grain shape, assumed to be a triaxial ellipsoid.

Verified by comparison to crystal tessellation software. (Neper: [doi:10.1016/j.cma.2011.01.002](https://doi.org/10.1016/j.cma.2011.01.002))



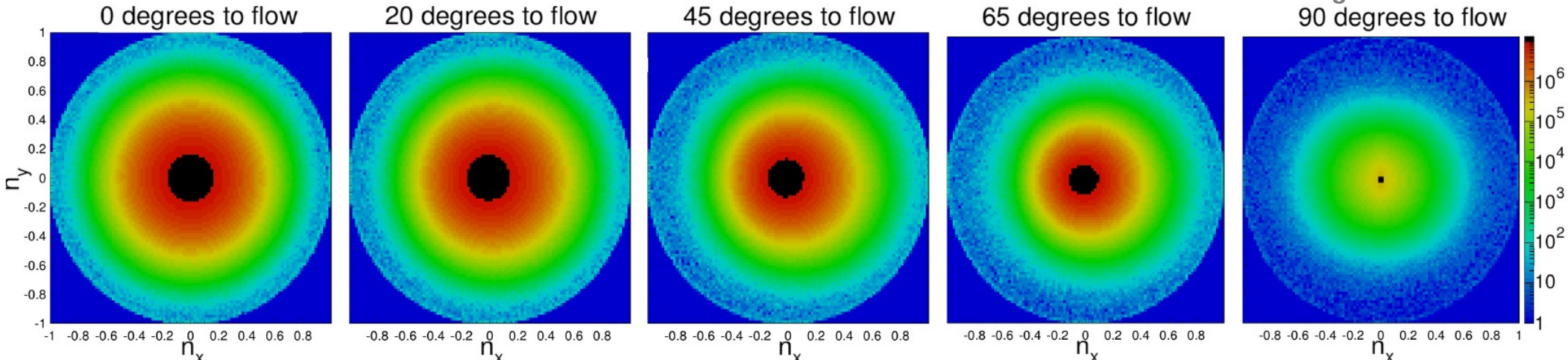


# Birefringence Monte Carlo

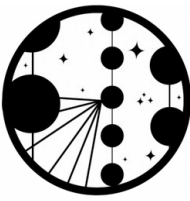
Running MC simulation with many photons through 1000 grains with ideal girdle fabric shows two effects:

- I. Diffusion is largest on flow axis and smallest orthogonal to it
- II. Photons on average get deflected towards the flow axis

→ **photons effectively fly a curve towards the flow axis**

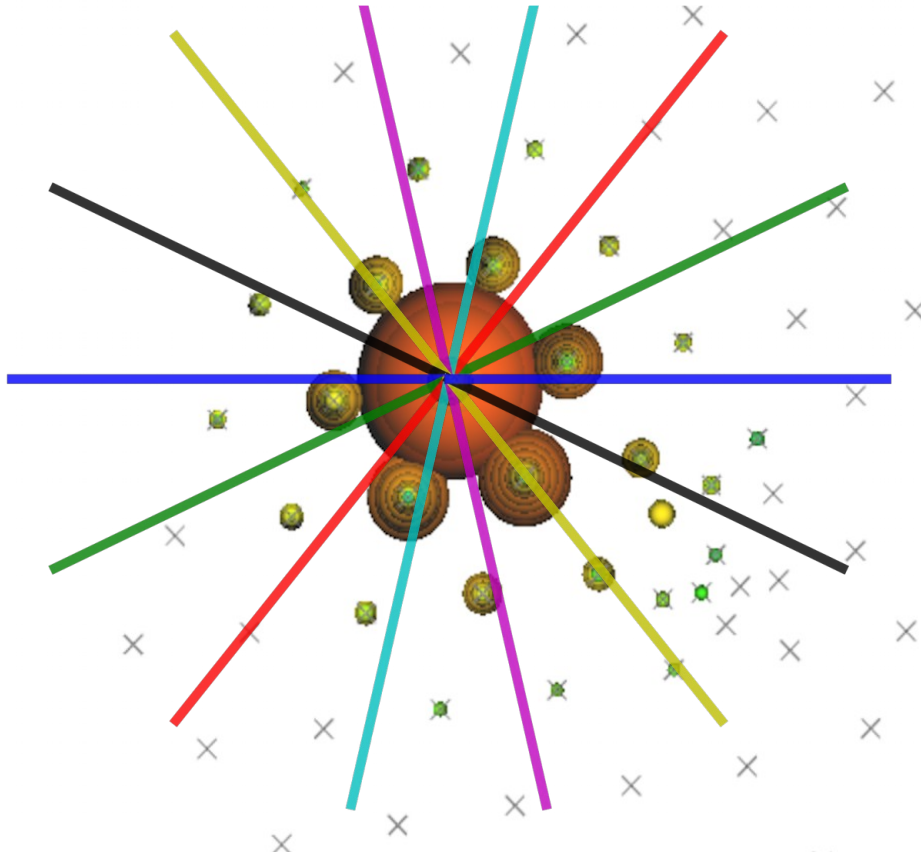




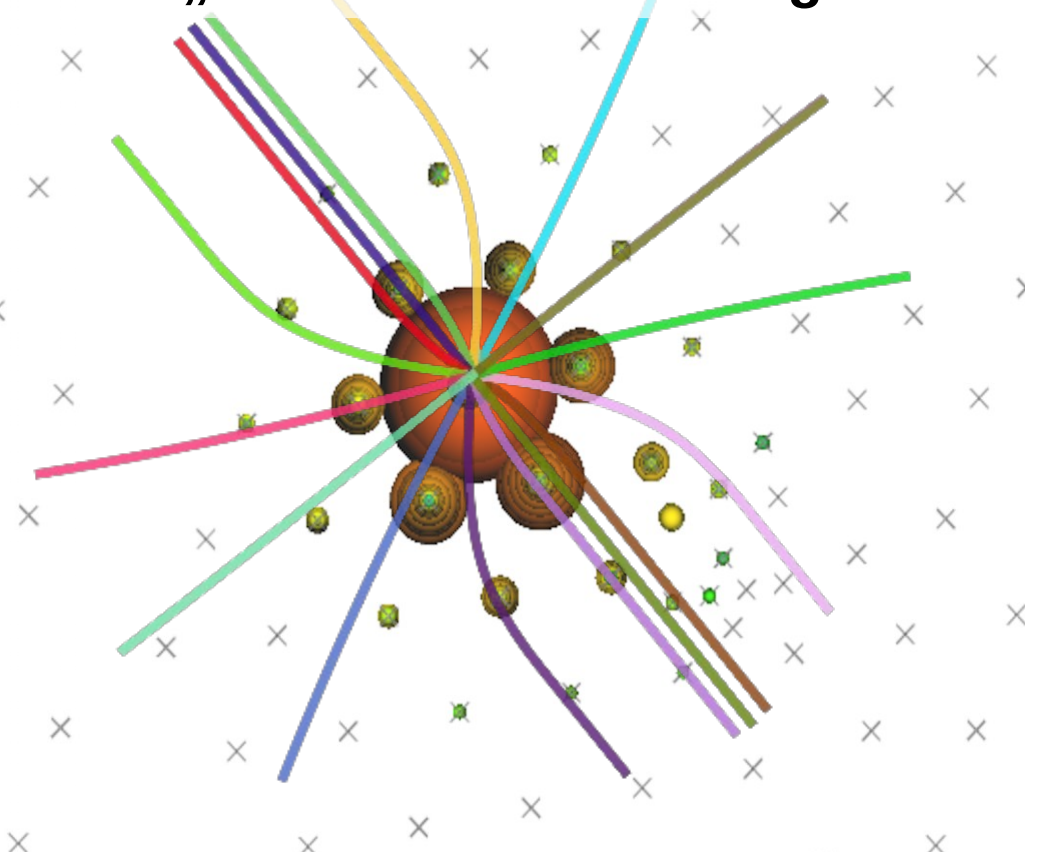


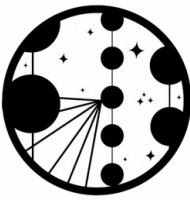
# Average photon trajectories

No birefringence



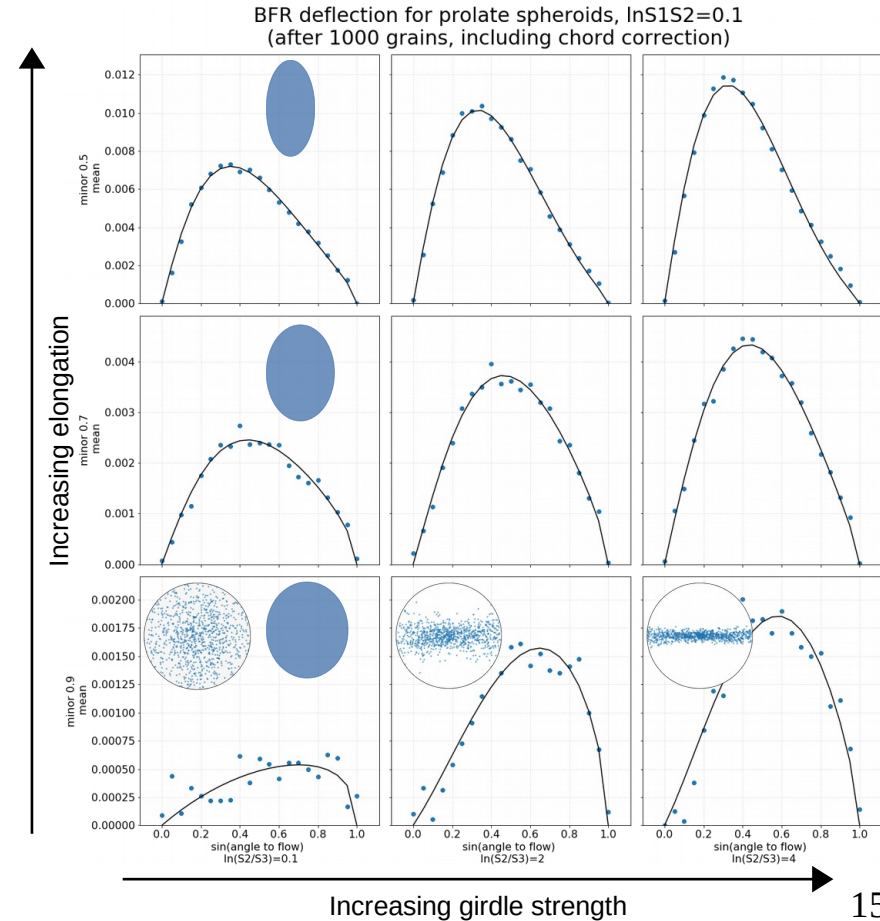
„Reasonable“ birefringence

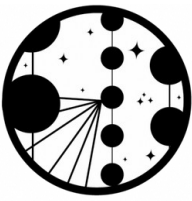




# Parametrization

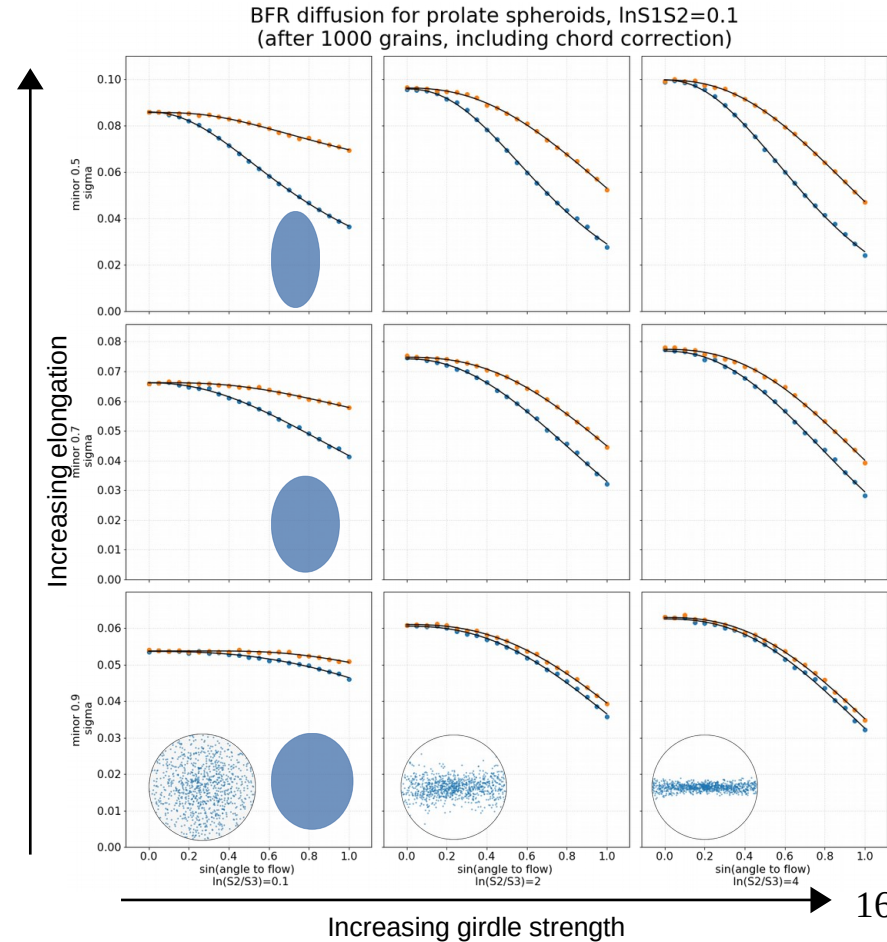
- Can not perform per crystal simulation for full-scale photon propagation
- Parametrize complex patterns using a displaced 2D-Gaussian
- In PPC photons are displaced and deflected at each scattering site (after proper distance scaling) depending on their current angle to the flow
- Elongation has a stronger effect compared to fabric (these in reality are of cause related)
- Free parameters:  
 $\ln(S1/S2)$ ,  $\ln(S2/S3)$ , elongation, size  
 (+ absorption, scattering & absorption anisotropy)



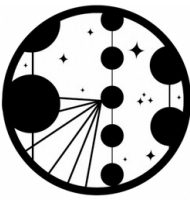


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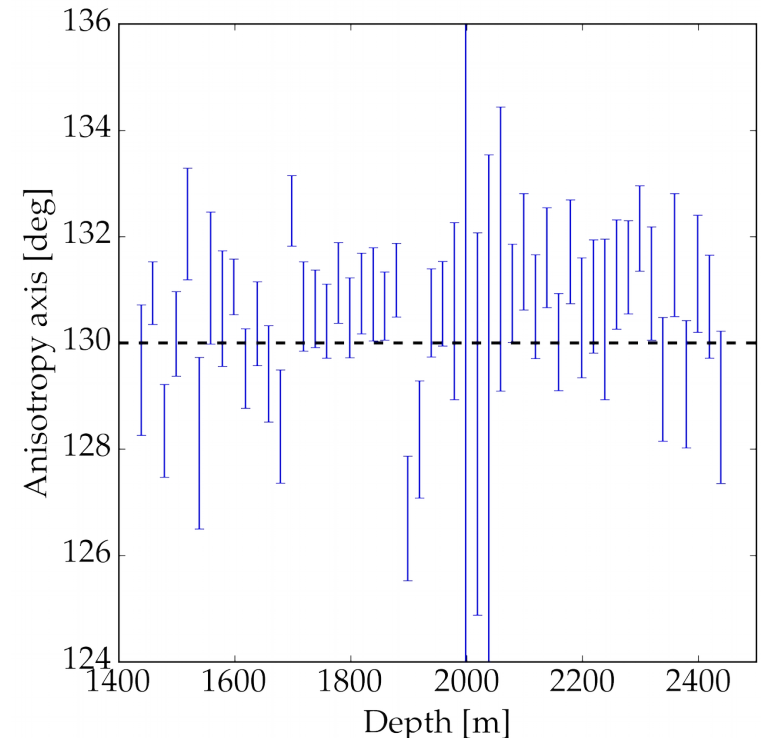


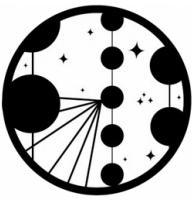




# Pre-fits

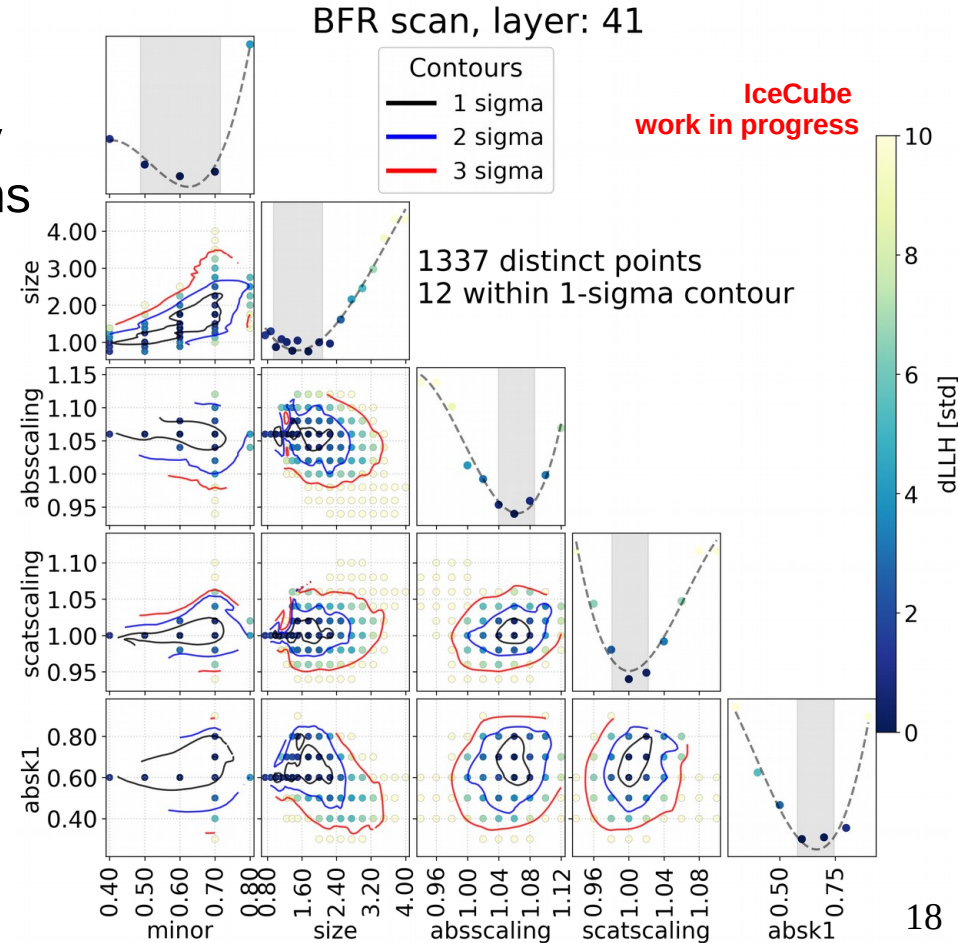
- We can't optimize 9+ dimensions at once, especially as a function of depth...
  - need to find parameters which are less relevant or near constant at all depths
- Spheroids (horizontal = vertical minor axis) strongly preferred over arbitrary ellipsoid
- The anisotropy angle is indistinguishable from  $130^\circ$  at all depths and over the entire surface footprint
- Given a girdle ( $\ln(S2/S3) \gg \ln(S1/S2)$ ) the actual fabric values are near indistinguishable
  - fixed to  $\ln(S1/S2)=0.1$  &  $\ln(S2/S3)=4$  at all depth (may be problematic in the very deep ice where at some point a switch to uni-modal is expected)

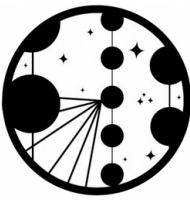




# Single layer fit example

- Size and elongation are strongly correlated  
→ data just as well described by small, nearly spherical or large, strongly elongated grains
- Fix minor axes at 0.7 (elongation  $\sim 1.4$ )
- Birefringence only model describes the data better than any previous model
- BUT crystal sizes are unphysically small
- Mixing in additional absorption anisotropy strongly improves data-MC agreement and yields sensible crystal sizes
- Including the scattering based model does not improve the fit (not shown here)

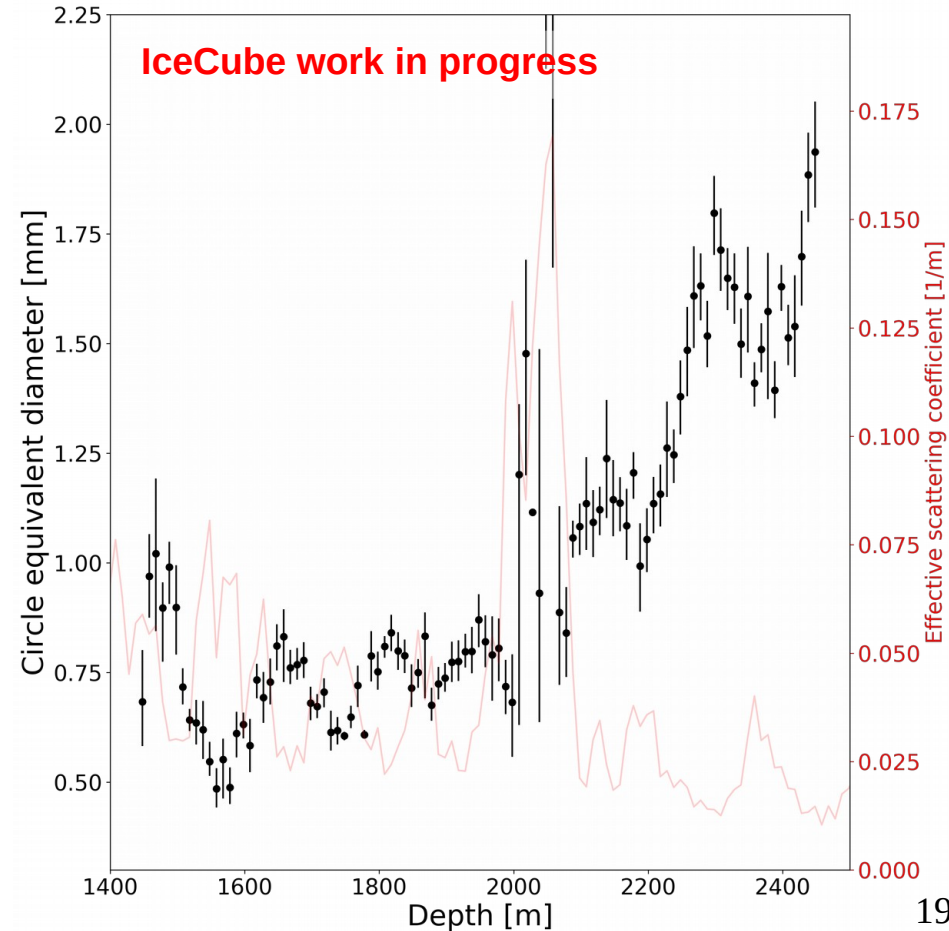




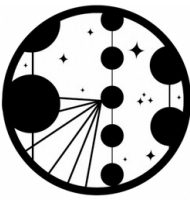
# Grain size vs. depth

- Due to the assumptions on unconstrained parameters (elongation, absorption anisotropy...) the size has an overall scaling uncertainty
- The fixed parameters (such as fabric) may also bias the result
- Still we see that grain size increases with depth and in particular below the dust layer
- Grain size seems to be anticorrelated to the scattering coefficient → crystals are smaller where scattering is stronger, there is more dust

**Disclaimer: This fit is still ongoing! Best fit values and statistical errors will change.**

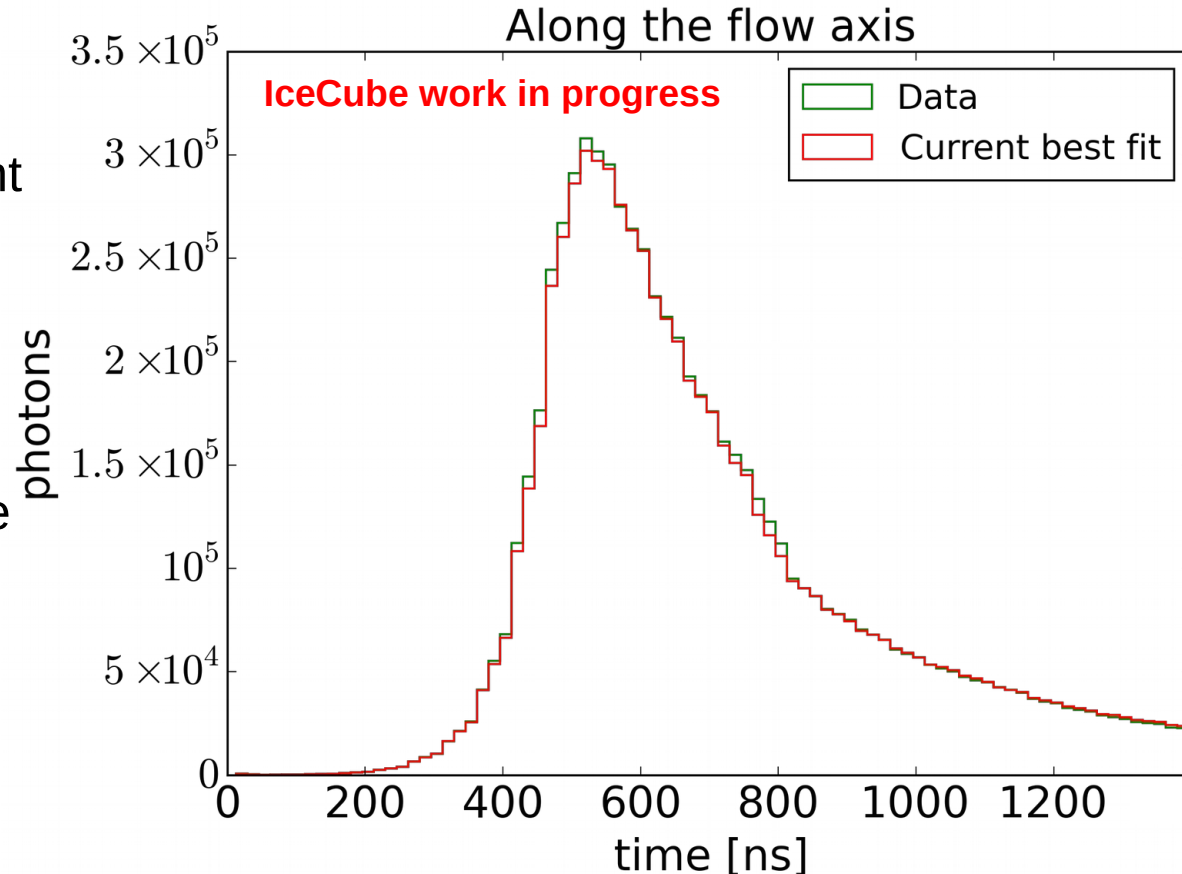






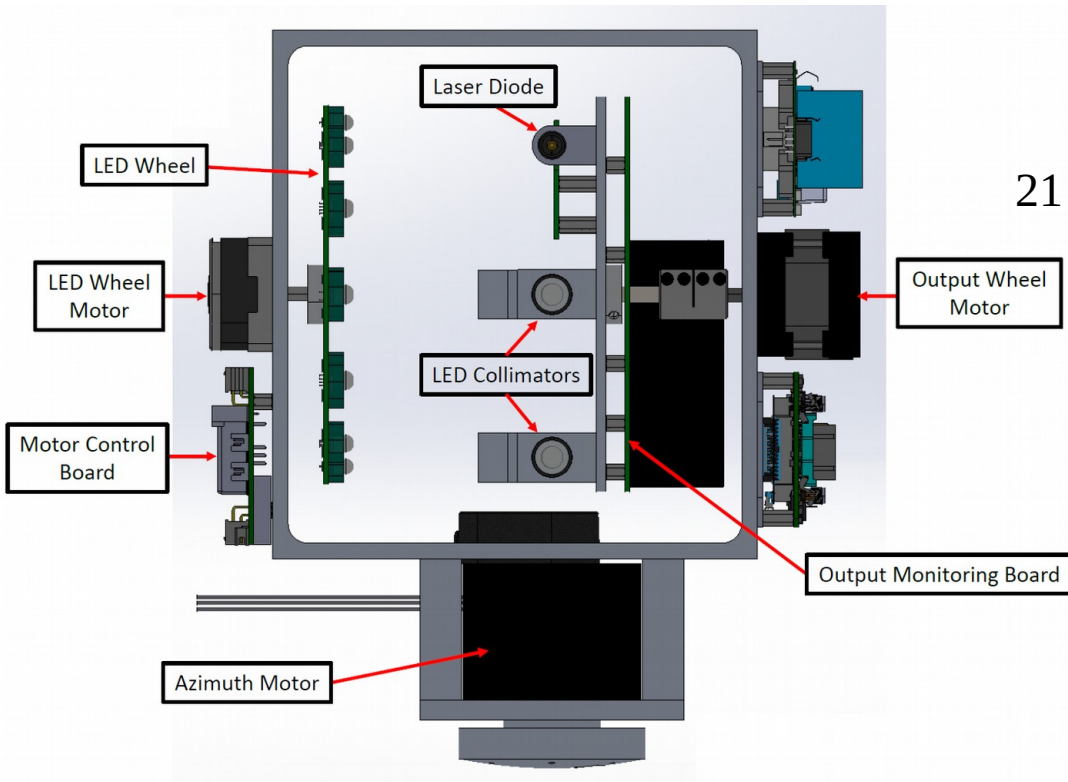
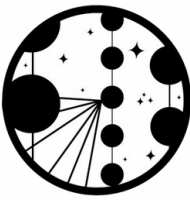
# Quality & open issues

- Using the full-model described above, we achieve an unprecedented data-MC agreement
- This allows us to probe crystal properties using only PMT data sampled at 125m intervals
- BUT the LED data is insufficient to unambiguously determine the large number of parameters involved
- And we still require a first principle explanation for what appears to be an absorption anisotropy (which may in turn change the birefringence understanding)

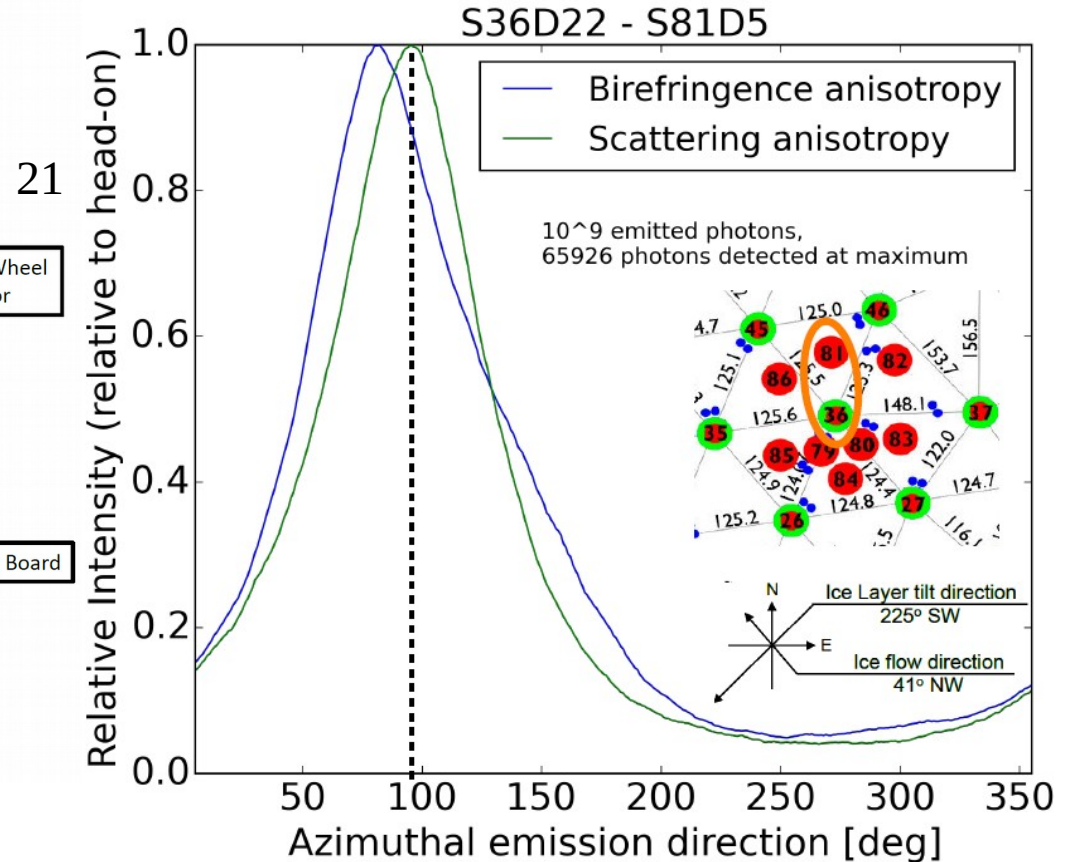


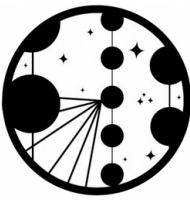
# Pencil beam in the Upgrade

Rongen, Chirkin  
Polar Science 2021



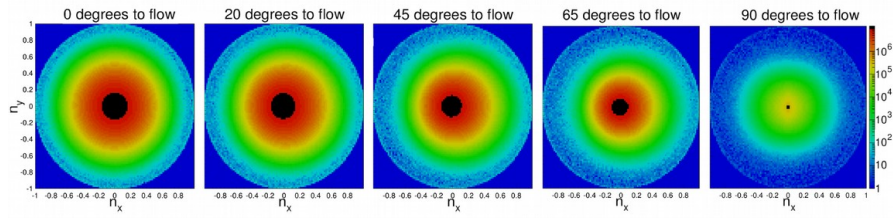
$4\pi$  steerable, ns, collimated light source





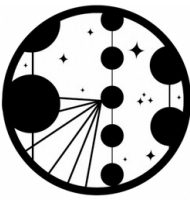
# Summary

- IceCube observes anisotropic propagation aligned with the ice flow
- Best best modeling of the effect is achieved through anisotropic absorption (still to be understood) combined with diffusion in birefringent polycrystals with preferential c-axis distribution which results in an average photon deflection
- The resulting ice model achieves an unprecedented data-MC agreement
- In the process IceCube is able to deduce crystal properties based on single photon arrival times sampled at 125m distances



Thank you for your attention!  
Questions are welcome



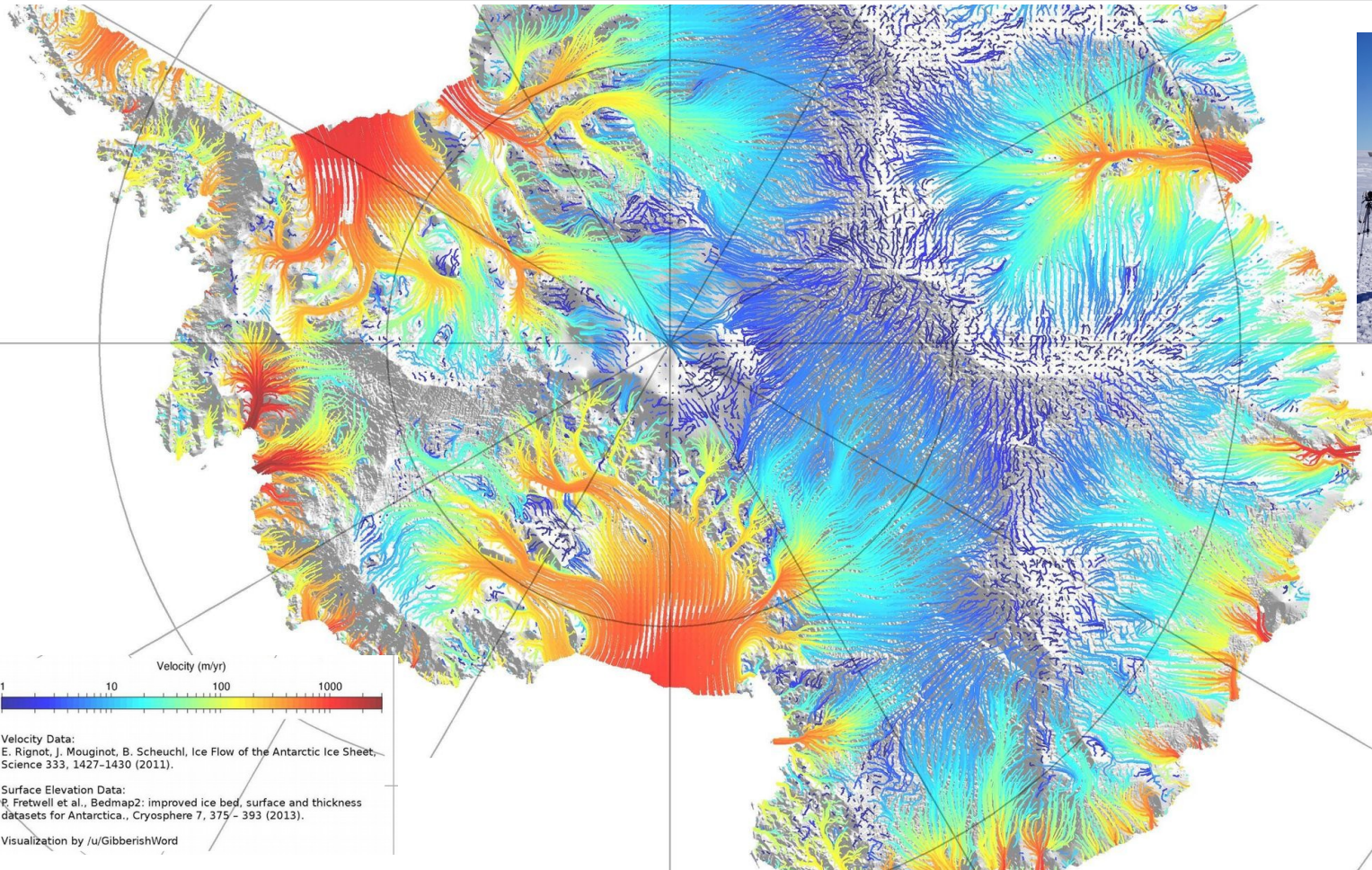
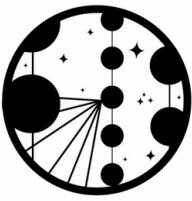


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# Backup....



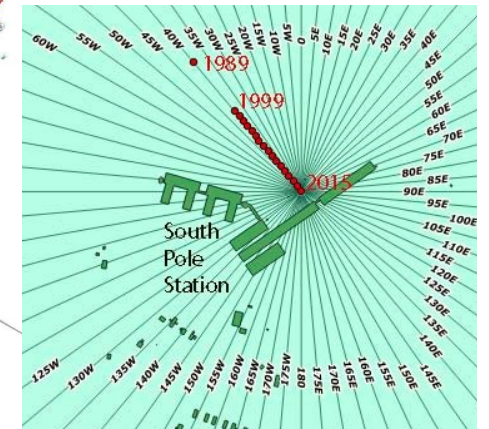
# Rivers of ice



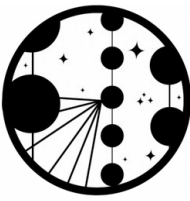
Velocity Data:  
E. Rignot, J. Mouginot, B. Scheuchl, Ice Flow of the Antarctic Ice Sheet, Science 333, 1427-1430 (2011).

Surface Elevation Data:  
P. Fretwell et al., Bedmap2: improved ice bed, surface and thickness datasets for Antarctica., Cryosphere 7, 375 - 393 (2013).

Visualization by /u/GibberishWord







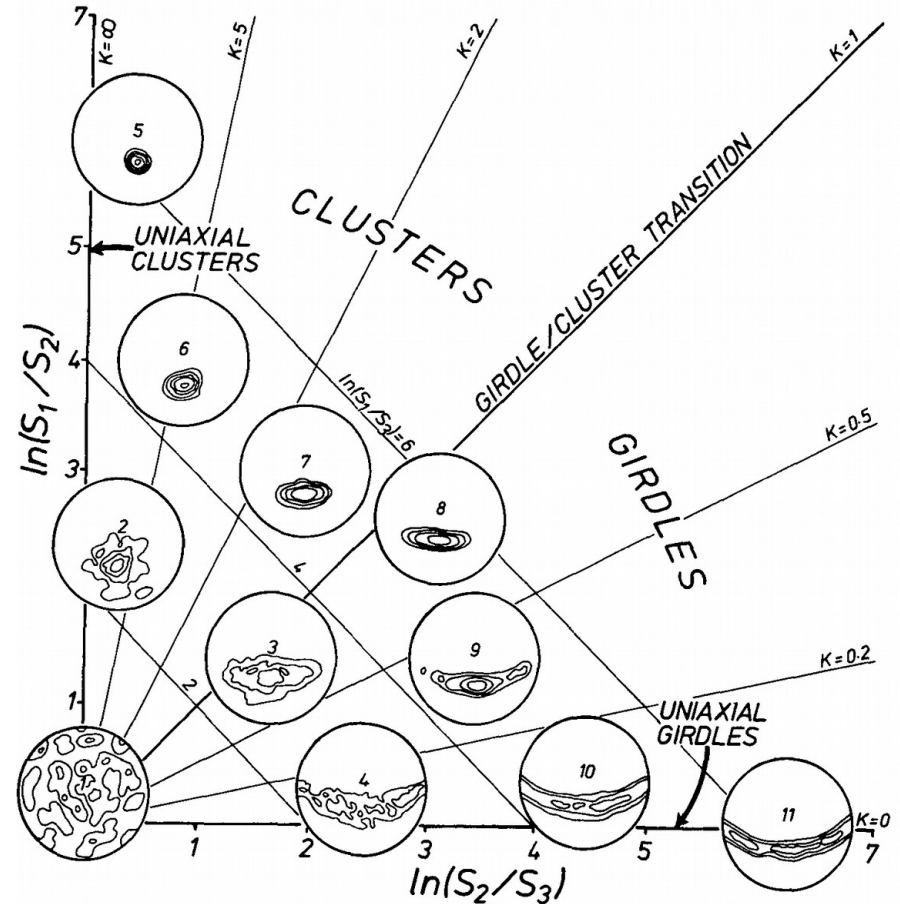
# c-axis sampling

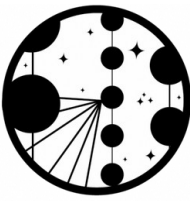
We can describe an ensemble of c-axis by their direction vectors  $n$  and construct the orientation tensor:

$$a = \begin{bmatrix} \sum n_{ix}^2 & \sum n_{ix} \cdot n_{iy} & \sum n_{ix} \cdot n_{iz} \\ \sum n_{iy} \cdot n_{ix} & \sum n_{iy}^2 & \sum n_{iy} \cdot n_{iz} \\ \sum n_{iz} \cdot n_{ix} & \sum n_{iz} \cdot n_{iy} & \sum n_{iz}^2 \end{bmatrix}$$

This matrix can be diagonalized.  
This is equivalently to choosing the coordinate system such that the x-axis is along the axes mean and the other two are orthogonal.

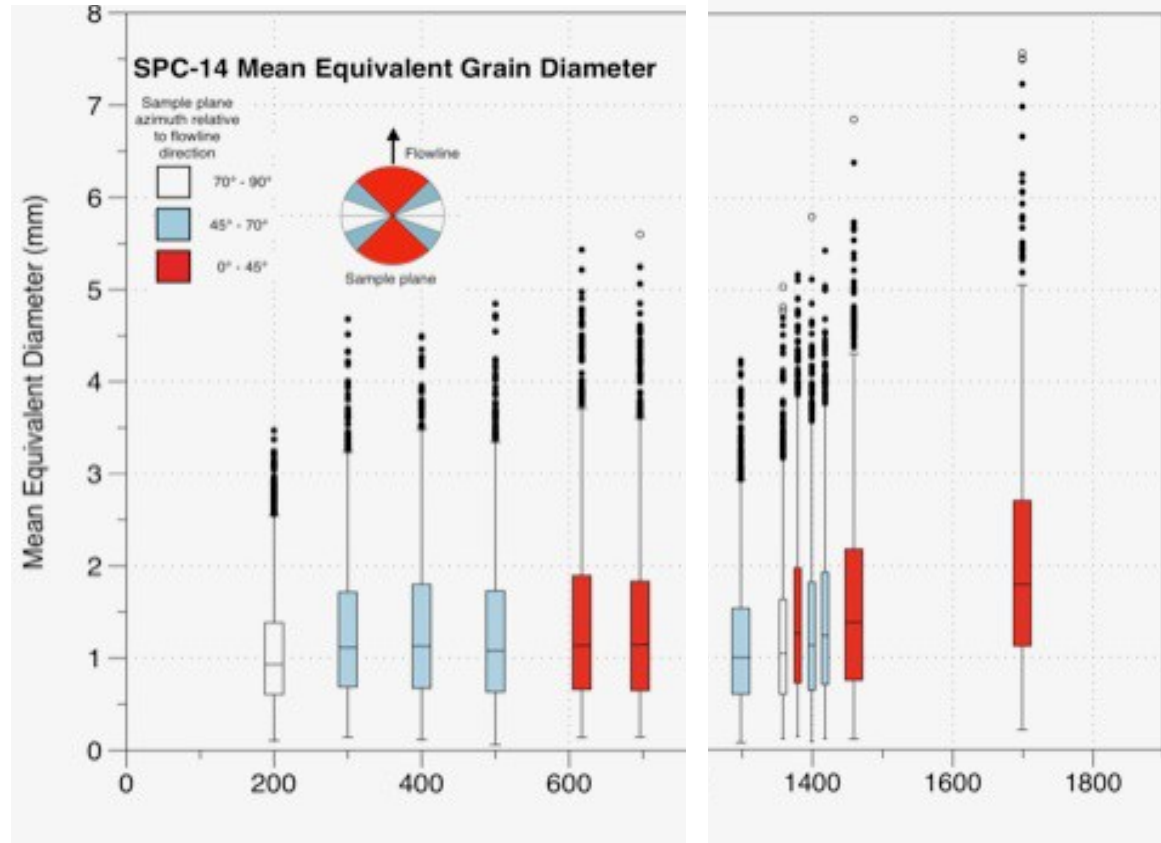
It's eigenvectors are then the axes and it has three eigenvalues with:  $S_1 > S_2 > S_3$





# Circle equivalent diameter

- SPICEcore measures grain size as the circle equivalent diameter
- That is the diameter of a circle with the same area as the grain
- It obtains 1-2mm  $\varnothing$  in deep ice
- For an ellipsoid the circle equivalent diameter depends on the cut plane
- In our parametrization the range is between:
  - $d \text{ [mm]} = 2 \sqrt{\text{minor} / \text{size}}$
  - &
  - $d \text{ [mm]} = 2 \text{ minor} / \text{size}$



Juan Fitzpatrick, private communication